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# 北美科迪勒拉山系 中生代大陆拼合

[美]Robert S.Hildebrand 原著

邱瑞照 李小伟 周 肃 孙 凯 陈玉明 等译



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## 北美科迪勒拉山系中生代大陆拼合

[美] Robert S. Hildebrand 原著

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地资出版社

・北京・

本书为国家国际科技合作专项"中美环太平洋成矿带成矿规律对比与资源评价技术研究"(2012-2015)项目成 果之一。作者 Robert S. Hildebrand 通过广泛资料收集和综合研究,系统论述了晚泥盆世-早密西西比世安特勒造山运 动以来,科迪勒拉山系形成过程的新模型,提出卢比亚带状大陆是由西雅里塔和原卢比亚在 160 Ma 时碰撞拼合组成的 复合体,北美大陆西缘部分向西俯冲到卢比亚带状大陆之下,碰撞分为两个阶段,最初的碰撞导致形成局部性的塞维 尔(Sevier)褶皱-逆冲带,后来的碰撞导致更广泛的拉勒米变形事件;新模型认为,约 125 Ma 时,卢比亚到达北美 大陆附近;随后卢比亚带状大陆与北美的大盆地分区(GBS)碰撞;北美向西俯冲板块在 100~96 Ma 时发生断离,导 致了板块断离岩浆作用活动,并且这时北美和卢比亚一起向北移动,但是因为卢比亚带状大陆的移动更慢,形成它们 之间左旋剪切运动;在大约 80 Ma,北美开始向南移动并与整个卢比亚碰撞,大洋俯冲板块断离产生板片断离岩浆作 用,同时卢比亚地体开始和太平洋板块一起向北运动;提出科迪勒拉山系造山带是一个以多个弧-陆和弧-弧碰撞为 特征的典型碰撞带,没有必要用科迪勒拉型模式(弧后模式)来解释其成因;指出北美大陆没有通过渐进式加积方式 向西逐渐增生。

本书可供从事大地构造、造山带过程以及境外地质矿产研究相关人员参考使用。

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### 前 言

《北美科迪勒拉山系中生代大陆拼合》是继 B. C. Burchfiel 编著的《北美大地构造》(1992) 以来,最为系统地论述北美西部形成演化的论著。非常令人敬佩的是 Robert S. Hildebrand 博士查 阅了海量资料、得到众多专家学者的帮助、自费完成该书,并以此献给 Cliff Hopson 先生,理由 是他最早向作者揭示了加利福尼亚州中生代岩石的奇妙和问题。该书的主要观点包括:

1)对于不同区段的科迪勒拉山系造山带,一个广泛被接受的假说是认为它发育在一个长期向东倾斜的俯冲带之上,然而,这个假说与许多重要的观察事实却不相符。因此,作者探索了一种替代的碰撞模型,即北美大陆西缘部分向西俯冲到卢比亚带状大陆 (Rubian ribbon continent) 下面,碰撞分为两个阶段,最初的碰撞导致形成局部性的塞维尔 (Sevier)褶皱-逆冲带,后来的碰撞导致更广泛的拉勒米变形事件。

2) 卢比亚带状大陆是一个逐渐拼合组成的复合体,但在 160 Ma 时,两个以前的复合地块发 生了碰撞,分别称为西雅里塔 (Sierrita)和原卢比亚 (Proto-Rubia)。原卢比亚 (Proto-Rubia) 在密西西比纪时期由罗伯茨 (Roberts)山外来体与安特勒 (Antler)边缘碰撞形成,安特勒为一 个来源不详的新元古代和古生代的被动陆缘,在 260~250 Ma 时,添加了包括育空 - 塔纳纳 - 斯 莱德山 (Yukon-Tanan-Slide Mt.)地体和宝山 (Golconda)外来体。西雅里塔形成于中侏罗世时 期约 170 Ma 和 160 Ma 之间,当包括斯马特维尔 (Smartiville)、斯莱特溪 - 结合湖 (Slate creek-Lake Combie)和海福克 (Hayfork)等几个面朝东的弧被合并到内华达山区 (Sierran) - 黑石 (Black Rock)弧的西侧之前,大约在同一时间,其东部又与原卢比亚 (Proto-Rubia)的西部边 缘碰撞。随后发生的俯冲板片断离形成一个与弧平行的后碰撞侵入岩套,包括独立的岩墙群和双 峰式、碱性柯维亚 (Ko Voya)岩套。在卢比亚西边缘之下,新的向东俯冲开始于 169 Ma 到约 130 Ma 之间的某一时期。

3)科迪勒拉山系造山带的塞维尔(Sevier)阶段开始于约125 Ma,即当位于北美克拉通的 大盆地区段一个隆起被拖入到向西倾斜的俯冲带时,这个俯冲带处于卢比亚超级地体的古大洋一 侧,进入海沟的边缘形成了塞维尔褶皱-逆冲断层带,并导致外来的大型逆冲席增生到北美大陆 西部。在碰撞期间(?)或之后,卢比亚超级地体的大部分相对于北美大陆向南迁移。约100 Ma 时的西部,一个右旋转换挤压碰撞导致了加拿大和阿拉斯加的格拉维纳(Gravina)-Nutzotin-Dezadeash-甘比尔(Gambier)等盆地闭合、下加利福尼亚(墨西哥境内)的Alistis 弧增生和内 华达山脉内一个现今的神秘盆地闭合;诸如拉波斯塔(La Posta)和内华达山脊后碰撞深成岩 套,可能是由于俯冲板片断离所导致。

4) 在 80 Ma 左右,北美大陆开始向南迁移,这导致了与整个卢比亚带状大陆碰撞,该带状大陆展布自阿拉斯加至少延伸到南美洲北部,在科迪勒拉山系造山带的拉勒米阶段时期,才成为北美大陆的外侧边缘。随着沿(带状大陆)边界两边俯冲的停止,导致:①科迪勒型岩浆作用终结;②增生杂岩内的弗朗西斯科(Franciscan)蓝片岩剥露,该增生杂岩沿卢比亚大陆西边界分布;③在索诺拉-莫哈韦沙漠地区,有一个俯冲板片断离岩浆作用形成的线状岩浆带与当时邻近海岸的深成岩带侵位;④卢比亚大陆向北斜向迁移。卢比亚和北美大陆之间的斜向汇聚导致在

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奧罗菲诺 (Orofino) 断裂以南的大盆地段出现厚皮构造变形,而在阿拉斯加、加拿大、索诺兰 沙漠等地则出现薄皮褶皱和逆冲断层。在向北迁移前,原先拼合在一起的地体 (现在位于加拿 大科迪勒拉山脉) 位于几千千米远的南方,它的南端和索诺兰沙漠分区的北端拼接。

5)碰撞依次为线性的、平行造山带区域的伸展崩塌和剝露,最后,加厚碰撞带的崩塌形成 盆-岭省,出现在大盆地和索诺兰沙漠分区,那里的北美克拉通已被牵引到卢比亚超级地体之下。新的向东俯冲在约53 Ma 左右开始,俯冲到混合的碰撞带之下,并持续至今。

本书的初衷是我们自 2008 年以来承担科技部的国际科技合作项目,与美国地质调查局西部 中心的专家合作开展中美环太平洋成矿带对比研究。2014 年 4 月,国家国际科技合作专项"中 美环太平洋成矿带成矿规律对比与资源评价技术研究"(2012~2015)项目合作者、美国地质调 查局 Steve Ludington 博士应邀来华工作访问,带来了 Robert S. Hildebrand 的《Mesozoic Assembly of the North American Cordillera》一书,并兴冲冲地告诉我们"你们对美国西部所有疑问,答案都 在这里!"。众所周知,非洲大陆是白垩纪时期伴随大西洋张开与南美大陆分离的,因此,环太平 洋成矿带是具有全球意义的中新生代巨型构造-岩浆成矿带,该带涉及的国家多、资源丰富,而 中国东部、美国西部是开展对比研究的最佳地区,从中总结的规律,对于"有的放矢"地在南 美、东北亚、东南亚等地实施矿产资源"走出去"战略具有重要的启示和示范作用,可望为有 重点地在相关国家开展"资源外交"提供理论和技术支撑;同时,符合我国矿产资源"走出去" 战略总体布局。

国家国际科技合作专项"中美环太平洋成矿带成矿规律对比与资源评价技术研究"(2012~ 2015),由中国地质调查局发展研究中心承担,中国地质大学(北京)、中金集团参加,与美国 地质调查局西部中心合作完成。项目采用"走出去"与"请进来"相结合的方式,分别赴美国 西南部、西北部、中国东北部、东南沿海、三江等地考察:通过系统收集、整理,建立了中国东 部、美国西部基础地质矿产数据库 (火成岩成分、矿床和火成岩年龄、矿产地),结合实地考 察,建立了中国东部、美国西部中新生代构造—岩浆—成矿事件序列,对比了大规模成矿作用差 异;结合南美、墨西哥、加拿大、中南半岛、俄罗斯远东等地已有资料,全面对比了环太平洋成 矿带成矿规律:在吸收其他方法优点基础上,总结了"地质过程解析的资源评价与选区"知识 结构、有关术语解释、工作步骤,丰富了我国资源评价方法和内容;提出中国东部(晚中生 代)、美国西部(新生代)大规模成矿作用与太平洋板块并无直接关系,即为"非俯冲性质", 大规模成矿作用发生与加厚的造山带垮塌有关,受控于造山岩石圈去根、软流圈物质和热上涌引 发的大规模岩浆活动;首次将造山带垮塌过程分为三个阶段,以青藏高原为代表的造山带属于初 始垮塌阶段,以美国西部科迪勒拉为代表的属于不完全垮塌阶段,以中国东部为代表的属于完全 垮塌阶段,对于重新认识大规模成矿作用背景、成矿规律具有重要参考意义。2015年5月28日, 项目成果通过科技部组织的专家验收 (北京香山饭店). 获专家好评。《北美科迪勒拉山系中生 代大陆拼合》一书对于项目任务的完成提供了重要参考。"中美合作"项目,得到有关单位和领 导的关心和大力支持。中国地质调查局发展研究中心徐勇、严光生 (现任中国地质调查局总工 程师和中国地质科学院书记)、邓志奇、齐亚彬(现任中国地质科学院测试研究所主任)、蔡刚、 施俊法 (现任中国地质调查局办公室主任) 等领导十分重视此项工作,从多方面提供保障和业 务指导:中国地质调查局发展研究中心科技处、财务处、办公室等职能处人员在许多具体事务上 给予了很多帮助。

在项目实施过程中,陆续组织项目组成员及有关专家等对《北美科迪勒拉山系中生代大陆拼合》一书进行翻译,中国地质大学(北京)周肃、李小伟、李小鹏等,中金集团公司刘彦兵、

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中国地质调查局天津地质调查中心孙凯、张琳琳等分段进行了翻译:要点段落,中国地质调查局 发展研究中心总工程师谭永杰研究员、技术指导中心副总吕志成研究员、《地质通报》主编刘志 刚研究员、广西自治区地质矿产勘查开发局副局长战明国研究员、四川鑫顺矿业股份有限公司董 事长黄朝刚教授级高工中国地质调查局沈阳地质调查中心李永飞高级工程师等进行了详细讨论和 译校:中国地质大学(北京)硕士生陆贵龙、张国庆、凌丹、朱泉龙、赵立克、任晓栋等翻译 了有关图表,蔡琍玲绘制了部分图件;肖庆辉、王保良、李廷栋、叶天竺、邓晋福、潘桂棠、彭 聪、毛建仁等专家十分关心本书工作并提供了许多有益帮助:中国地质调查局发展研究中心境外 部向运川、叶锦华、陈秀法、陈玉明、赵宏军、张振芳、崔敏利、李娜、陈喜峰、王杨刚、何学 洲、王靓靓、张潮等提供了诸多帮助和鼓励。在本书译校过程中,有关术语的翻译与武警黄金指 挥部地质处徐年生研究员、中国地质科学院地质研究所曾令森研究员、美国爱达荷大学王达博士 等进行了有益讨论:美国地质调查局地质、矿产、能源和地球物理学(GMEG)科学中心 Niki E. Wintzer 博士帮助联系美国地质学会执行主编 April Leo 博士和 Robert S. Hildebrand 博士,承蒙授 权出版该书。全书由邱瑞照统稿、由于首次外文出版物、加上书中涉及范围大、专业术语和地名 多,进行了反复核对、翻译和校对,耗时良多:为便于读者理解北美地质,在书后列出了附录, 内容包括:有关术语、地名英汉对照;美国州名 (州名缩写) --首府名;美国西部重要地质事 件:北美地壳运动等。

本书是国家国际科技合作专项"中美环太平洋成矿带成矿规律对比与资源评价技术研究" (项目编号:2011DFA22460)成果之一。感谢蔡刚副主任(现任中国地质图书馆书记)积极协 调出版经费,获得中国地质调查局"海上丝绸之路重点地区有色金属资源潜力评价"二级项目 (项目编码:DD20160118)资助。

在此对以上单位和个人给予的指导、支持和帮助一并致以衷心的感谢!

本书涉及的范围大、专业术语和地名多,其中翻译难免存在错漏和偏颇, 敬请广大读者批评 指正。

译者

#### 2017年8月8日

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### 第1章 引 言

科迪勒拉山系造山带的独特性在于它是两个性质截然不同的岩石圈发生碰撞所形成的:其中一个 是致密、低位、下伏大洋岩石圈,而另一个则为低密度、有浮力的上覆大陆岩石圈。关于这一观点的 形成源于 40 年前 Hamilton (1969a, 1969b) 提出的现代安第斯山火山岩带可以作为北美大陆大岩基 形成的现实模型。很快,新的认识蜂拥而至,Hsu (1971)认识到弗拉西斯科混杂岩代表了一个古老 的俯冲杂岩: Moores (1969, 1970) 认为蛇绿岩是大洋地壳和下伏地幔的残片: Ernst (1970) 提出 大谷地(Great Valley)-弗拉西斯科杂岩接触带是一条中生代毕鸟夫带(Benioff zone)的标志; Dickinson(1970)将弧岩浆作用、岩基和弧前沉积与俯冲作用联系起来; Dewey和 Bird(1970)创造了 科迪勒拉山系造山带这个词,用来描述这些地质分带。从那时起,大多数地质学家认同这样的假设, 即美洲大陆西缘构成向东倾的大洋板块的上覆板块, 起始于晚泥盆世 (Burchfiel and Davis, 1972, 1975; Price, 1981; Monger and Price, 2002; Dickinson, 2000, 2004, 2006; Colpron et al., 2007) 或 三叠纪 (Schweickert and Cowan, 1975; Ingersoll, 2008)。以上两种解释均认为下伏大洋板块和上覆 大陆板块的汇聚和相互作用塑造了科迪勒拉山系造山带,伴随巨量的科迪勒拉型大岩基、局部区域的 双倍地壳厚度、内陆带变质作用和弧后褶皱-逆冲带(Armstrong, 1974; Coney and Reynolds, 1977; Dickinson and Snyder, 1978; Saleeby, 2003; DeCelles, 2004; DeCelles et al., 2009)。以上是科迪勒拉模 型整体概况,我宽泛地将其称之为弧后模型,因为绝大多数的变形被认为形成于一个弧后背景,除了 轻微的地区差异,该模型被极为广泛地应用于现代和古老大陆边缘的成因解释。

在现有的地质格架中,加利福尼亚州地质扮演了一个决定性的角色,因为该区弗朗西斯科 (Franciscan) 混杂岩-大谷地弧前-内华达山脉岩基构成的三位一体格架, 被认为分别代表了"中生 代俯冲杂岩-弧前盆地-大陆岛弧" (Dickinson, 1981), 这些岩石构造在加利福尼亚州出露良好, 并 且大多数地质出露点都可以轻易到达,因此它们成为科迪勒拉模型中岩石-构造的核心要素 (McPhee, 1993; Moores et al., 1999)。尽管经历了长期调查和大量的野外研究,但在许多问题上仍 然没有达成共识,所以主要问题仍然悬而未决。例如,在弗朗西斯科杂岩内近 123 Ma 时突然出现的 强烈增生阶段,在现有模型中难以得到满意解释,因为该模型无法提供一个供给巨量沉积物的合理源 区,也无法解释为何在那个时间发生这一变化。对于 100 Ma 在内华达山区和半岛山脉的科迪勒拉岩 基内发生的变形事件,也是知之甚少,难以合理解释,因为它们是碰撞后深成岩套。同样,对于一些 问题,诸如弗朗西斯科杂岩和海岸山脉蛇绿岩-蒂黑马-科卢萨(Tehama-Colusa)蛇纹化混杂岩-大 谷地群之间的接触关系是如何形成的,弗朗西斯科高级变质岩是如何被剥露出来的,或者这种剥露作 用是何时启动的等,这些问题在学界都还没有达成共识。为什么在内华达山脉和白印优山脉 (White -Inyo Mt.) 明显不存在拉勒米 (Laramide) 变形, 然而这一时期变形却在南美北部到阿拉斯加广泛存 在?为什么塞维尔(Sevier)逆冲作用仅仅局限于大盆地地区,以及为什么那里的岩浆作用在 80 Ma 的时候停止,而在其他地区却得以持续?上述这些问题,以及其他数不尽的长期存在的疑难问题可能 需要一个不同的、更有活力的模型才能更好地解释。

虽然对标准的科迪勒拉模型,一些学者提出过挑战,但他们是少数派并且没有得到广泛认可。例如,Moores等(Moores,1970,1998;Moores and Day,1984;Moores et al.,2002)长期坚持认为北美大陆西部部分地向西俯冲至一个岛弧之下。Mattauer等(1983)认识到阿尔卑斯山脉和加拿大斯瓦普(Shuswap)地体在变形和变质方面存在相似性,进而提出了他们的北美大陆西缘的俯冲模型。Cham-

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berlain 等(1985)和 Lambert 等(1988)的理念非常有先见之明,因为他们构建了一个加拿大科迪勒拉模型,在这一模型中,大多数的外来地体发生离岸汇聚,然后向北迁移与北美西部整体碰撞。最近,标准的科迪勒拉模型重新受到一个动态碰撞模型的挑战,在这一模型中,北美大陆向西的俯冲导致了白垩纪——古近纪的碰撞,拼合形成一个含岛弧的带状大陆 (Johnston, 2008; Hildebrand, 2009)。 在碰撞模型中,两者之间的缝合带位于现今的塞维尔--落基山脉 (Sevier-Rocky Mt.) 逆冲带内,恰好是北美大陆架西部边缘。

在本书中,我将论证碰撞模型如何能提供一个统一且合理的论据,解释包括加利福尼亚州西部在 内的广泛分布且看起来毫无关系的一些观察到的现象。大多数的这些关系在早前的文献资料中完全是 未预见到的,但它们有力地证明了碰撞模型的有效性。为了实现这一目标,我将首先介绍相关地质概 况,然后利用碰撞模型中的不同方面来解释加利福尼亚州地区以前无法解释的一些特征,最后,我将 提供一个集合造山带各关键要素的模型。以下描述内容是向读者提供一个背景知识,以便让读者了解 其中基本争论的由来。这种综述并不是为了佐证碰撞模型的合理性,这一回顾可能有些过时,但都是 可查询的,并包含基本的争论(Johnston, 2008; Hildebrand, 2009)。在我早期的研究中,我更多地关 注北美克拉通及其与卢比亚超级地体的相互作用。在这里,更多地关注超级地体自身的合并汇聚过 程。即便是无法接受碰撞模型的读者,也会从这里展示的科迪勒拉山系聚合模型中找到可以验证的有 价值信息。有关 Hildebrand (2009)的基本思想将在文中适当的地方作必要的重复,以便更好地理解 本书内容。

卢比亚是一个长的、线性的巨型地体或带状大陆,由几乎所有的北美科迪勒拉山系的外来地体或 超级地体组成。它拥有一个新元古界和古生代的地体核,随着时间的推移,它将各种各样的地体镶嵌 添加,并得以逐渐变大。它最大的增长出现在中-晚侏罗世期间,那时候,各种各样的岛弧,包括大 洋岛弧和大陆岛弧,增生到两个不同的地块上,即为西雅里塔(Sierrita)和原卢比亚(Proto-Rubia) 地块;两个地块转而在160 Ma时发生碰撞并形成卢比亚带状大陆(Rubian ribbon continent)。

### 第2章 现有的弧后盆地模型问题

对于那些尚未阅读到之前综合性文献的读者来说,值得花时间回顾现已被接受的北美科迪勒拉模型或者弧后模型中存在的若干内在的主要缺陷。笔者并不打算逐一指出那个模型中存在的所有问题,因为之前的文献已有更详细的论述 (Johnston, 2008; Hildebrand, 2009),在我看来,更为重要的是任何成功的模型都应解决一些关键问题。

1)在当前的北美科迪勒拉山系造山运动解释中,被动大陆边缘或冒地槽被认为是从克拉通向西 延伸至内华达州中部和加利福尼亚州东部(Stewart, 1970, 1972, 1976; Stewart and Poole, 1974; Armin and Mayer, 1983)。这样一个解释的主要问题在于被动大陆边缘面向西的地台边缘出现在现今 塞维尔褶皱-逆冲带内,到达或靠近盐湖城东面的瓦萨奇区(Wasatch Front) (Armstrong and Oriel, 1965; Peterson, 1977; Rose, 1977; Doelling, 1980; Palmer and Hintze, 1992); 然而,那里有另一个 地台(图 2.1),它的地台边缘主要位于内华达中部和加州东部(Kepper, 1981; McCollum and McCol-



图 2.1 大盆地分区内两个古生代大陆架边缘的位置示意图 (据 Morrow and Sandberg (2008)修改)

示意图显示安特勒大陆架边缘位于北美边界以西约 200~400 km。骑在向东逆冲断层之上的外来岩石覆盖了北美大陆架边缘附近的所有地区,除了位于怀俄明州西南的怀俄明凸起和加拿大科迪勒拉山脉南部的主要山岭之外,该区大陆架边缘得以保存并被称为踢马圈(Kicking Horse Rim)(未标出)

lum, 1984; Heck and Speed, 1987; Montañez and Osleger, 1996; Morrow and Sandberg, 2008; Sheehan, 1986; Harris and Sheehan, 1998; Stevens et al., 1998; Stevens and Stone, 2007)。北美或落基山脉的边缘形成于早寒武世,主要由陆源碎屑岩组成,上覆一个中寒武世碳酸盐岩台地,而在地台西部,即为Hildebrand (2009)命名的安特勒 (Antler)地台,再往北称为卡斯尔 (Cassiar)地台,它们最初形成于新元 古界时期,主要岩性为陆源碎屑,上覆早寒武世含古杯动物门的生物礁 (Oriel and Armstrong, 1971; Stewart, 1972; Fritz, 1975; Read, 1980; Pope and Sears, 1997)。Johnston (2008)和 Hildebrand (2009)均认为两个地台被深水相沉积相岩石分隔,两者具有不同的构造、沉积和岩浆演化历史;因此,提出两者处于不同的板块之上,只是在白垩纪—古近纪—新近纪的科迪勒拉造山运动时期才拼合在一起。

2)在弧后模型中难以解释的另外一个特征是在北美边缘缺乏新元古界末至早寒武世的裂谷盆地沉积及相应的裂谷火山沉积。除了少数几个地方,大多数现代裂谷和裂谷边缘(图2.2)都以发育大量火山沉积为典型特征(Ebinger, 1989; Ebinger and Casey, 2001; Menzies et al., 2002; Sawyer et al., 2007)。当然,存在这种可能,即部分裂谷边缘延伸过于宽阔广泛、无火山且非对称(Lister et al., 1991),以至于裂谷地壳主要在边缘一侧分布,就像现在的北大西洋边缘一样(Keen and Dehler, 1997);但是由于另一侧裂谷边缘的宽度减小,通过古地理恢复,塞维尔褶皱-逆冲带西侧很可能并没有被北美地壳推倒。根据现今的边缘,如果从阿拉斯加到墨西哥的整个边缘完全是非岩浆成因岩石的话,这似乎是不太可能的。劳伦系(Laurentian)地壳延伸超过整个造山带的长度,而总体缺乏前寒武纪末—寒武纪喷发的火山裂谷沉积,这是目前通用的模型完全无法诠释的。



图 2.2 在现今大陆边缘分布的被动陆缘类型 (据 Menzies et al. (2002)修改) 注意:与火山边界相比,非火山成因裂谷边界的分布和范围都是极其有限的;这表明北美科迪勒拉裂 谷相火山岩的缺失是特有的,并且需要一个合理解释。Hildebrand 和 Bowring (1999)认为,大多数裂 谷相岩石俯冲与板片断离有关

一些学者(Stewart, 1972; Burchfi el and Davis, 1975; Lund, 2008)认为温德米尔(Windermere)超群的岩石和其等同物乃至更老的岩石(Dehler et al., 2010),可代表北美大陆西部边缘的裂谷型沉积,但是由于它们比被动大陆边缘的演化时间还要早大约85~100 Ma,大陆边缘不可能保持足够的热量与早古生代的沉降相匹配(Bond and Kominz, 1984; Devlin and Bond, 1988)。此外,温德米尔超群不含"本来应该"具有的火山岩。

3) 在大部分古生代的岩石中连续出现的镁铁质岩浆作用通常被认为代表了被动大陆边缘的漂移 相产物(图 2.3),但如塞尔温盆地和科奇卡(Kechika)海槽(Goodfellow et al., 1995; Cecile, 2010)却很难与被动大陆边缘背景相吻合,因为最近在爱达荷中部识别出一套 664~486 Ma 的碱性深 成岩体(图 2.4)侵入到了贝尔特(Belt)超群岩石及其古生代冒地槽型沉积盖层(Lund et al., 2010; Gillerman et al., 2008)。此外,至少有一个深成岩体很可能是晚寒武世期间发生了去顶作用 (Link and Thomas, 2009; Link and Janecke, 2009),这是冒地槽外侧部分出现的一个独特特征。



### 图 2.3 在加拿大科迪勒拉山脉盆地相岩石中从寒武纪到泥盆纪的岩浆活动位置复原图 (据 Goodfellow et al. (1995)修改)

在传统模式下,这些岩石一般被认为是冒地槽的一部分,但此盆地相岩浆活动比漂移阶段被动陆缘盆地边缘的更典型。 卢比亚和北美之间的缝合线被认为位于盆地相岩石和北美古生代碳酸盐岩台地(Johnston, 2008; Hildebrand, 2009)之间, 在此图中,很可能是位于碱性深成岩和爆发角砾岩筒群南部,踢马圈(Kicking Horse Rim)的西侧。



(据 Lund et al., 2010)

4) 一般来说,从寒武纪到白垩纪的北美克拉通大陆架上,在变形的样式或者外来物源沉积方面 都不存在碰撞的证据。例如,内华达西部的广大地区假设是在密西西比期之后作为北美大陆的一部分 被合并到一起的,但是在北美大陆架看不到与160 Ma的变形事件相关的任何变形或沉积证据,包括 在黑石戈壁 (Black Rock Desert) 和附近地体存在的7~14 km 的地壳加厚和主逆冲断层作用 (Wyld et al.,2001; Wyld,2002)。整个内陆带在侏罗纪时期也有强烈的变形和变质作用,在那里随着向西伏卧 逆冲推覆体的发展演化,地壳增厚到原来的两倍 (例如, Camilleri et al., 1997),然后,大陆架没有受 到变形作用影响,所处位置通过盆岭省正断层的恢复两者相距 80~100 km (Hildebrand, 2009)。

类似的关系也存在于加拿大落基山,在塞尔扣克山(Selkirk)扇形构造内,坐落于莫纳希山(Monashee)杂岩的东侧翼部(图 2.5),那里出露一个侵蚀窗,可能是埋藏在库特尼(Kootenay)地体之下的北美克拉通基底和盖层岩石形成的双重构造被剥露的结果,该区 187~173 Ma 的深成岩体在变形期之前和或变形期间侵位,在173~168 Ma 之前岩石快速抬升剥露,压力从7千巴●降低到3千巴(Colpron et al., 1996)。Scrip 推覆体,一个向西的等斜构造位于北侧,它拥有一个倒转的翼部,沿着走向方向延伸达50~60 km,并且很可能在同一时期形成(Raeside and Simony, 1983)。在南部远端也存在类似的构造(Höy, 1977),但是尚未能够得到很好地约束,形成的时间也很可能在178~164 Ma 之间(Read and Wheeler, 1975)。Colpron 等(1996, 1998)反驳到这些事件发生在地层外侧和邻近北美的冒地槽(Colpron and Price, 1995),但是这一时期在北美克拉通界面上没有任何地质记录,如变形、沉积、主要侵入作用、褶皱、加厚和抬升剥露等(图 2.6)。

事实上,这个时期在北美地台上沉积的是磷灰岩, P<sub>2</sub>O<sub>5</sub> 含量达到 30%,这种类型岩石的形成归因于冷的、富含营养的海水上升,并以沿海洋东部边缘分布为特色 (Poulton and Aitken, 1989; Parrish and Curtis, 1982)。

同时显示了爱达荷州中南部上寒武统圣查尔斯组(St. Charles) 蠕虫溪(Worm Creek)成员中碎屑锆石 U-Pb分析结果,基于碎屑锆石年龄以 498 Ma 为主,表明 497 Ma 的深溪(Deep Creek)深成岩体(D)在侵入后被迅速 去顶剥蚀(Link and Thomas, 2009; Link and Janecke, 2009)。其他缩写是指个别岩体,不在这里讨论

<sup>● 1</sup>千巴 (kbar) = 100 兆帕 (Mpa)。









图 2.6 Illecillewaet 复向斜压力-时间曲线图 (据 Colpron et al. (1996)修改)

加拿大落基山中沿着莫纳希(Monashee)杂岩东侧出现的 Illecillewaet 复向斜压力-时间曲线图。显示复向斜内的岩 石在 170~160 Ma 期间快速剥露和同时发生的在开阔水域沉积的北美(NOAM)磷灰石矿床,表明这两个区域在 那个时候并不靠近。在本书中,我认为 Illecillewaet 复向斜形成于卢比亚超级地体内发生的广泛碰撞事件期间, 即它与北美大陆碰撞之前。

美国莫里森组、弗尼 (Fernie) 组以及加拿大库特尼群内的侏罗纪沉积岩是被动大陆边缘序列中 唯一的侏罗纪岩石单元,目前已知它们包含西侧的物源,但是不含火成岩-变质岩碎屑,它们比西部 内陆盆地内的前渊初始沉积岩至少早 25~40 Ma。此外,莫里森组向西没被加厚,相反从位于犹他州 -科罗拉多州边界的地台边缘有一定距离的沉积中心向西逐渐变薄 (Heller et al., 1986; DeCelles, 2004)。

5)最近的野外地质和 U-Pb 年代学工作证实,加拿大的许多外来地体,包括斯莱德山 (Slide Mountain)大洋地带,在三叠纪发生汇聚拼合,而不是弧后模型所要求的侏罗纪 (Beranek and Mortensen, 2007; Beranek et al., 2010a)。此外,坚实的古地磁资料表明这一地块和位于克拉通西侧的 其他科迪勒拉地体,直到近 70~60 Ma 时才与克拉通完全拼贴在一起 (Enkin, 2006; Enkin et al., 2006a, Kent and Irving, 2010)。

6) 贝尔特--珀塞尔(Belt-Purcell) 超群的岩石通常被解释为沉积于劳伦系(Laurentian) 地壳之上,但在贝尔特--珀塞尔超群的变质沉积岩和侵入体岩石中发现了 1.2~1.0 Ga 的变质作用和变形事件(Anderson and Davis, 1995; Nesheim et al., 2009; Zirakparvar et al., 2010),而在北美克拉通北西部从未发现,因此该现象应被认为是经过长距离迁移所导致。

7) 弧后模型不能很好地解释强烈的(地壳) 缩短、高级变质作用(>9 kb,约 800℃)、汇聚的 温度-时间轨迹,以及塞维尔内陆带的伸展垮塌,该内陆带在晚白垩世期间至少发生了 70 km 的水平 缩短和 30 km 的垂向加厚(Camilleri et al., 1997)。这就意味着那时候的地壳厚度是正常克拉通地壳 厚度的两倍。考虑到这些变形通常是"薄皮"的,尚不清楚一个弧后环境下如何能产生所需的挤压 应力,以及一个弧后盆地的减薄地壳在变形过程中如何产生明显的压力。实验和地震数据显示在其他

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造山带中类似的带容易被解释为发育良好的侵蚀双重构造,由下盘岩石和外来岩石的飞来峰组成,分 布在前陆盆地一侧 (Malavieille, 2010; Schmid et al., 2004)。

8) 弧后模型的最新变化是大约 400 km 的北美克拉通地壳可能已经向西俯冲到内华达山脉之下, 才能平衡盖层中上部地壳的缩短 (DeCelles et al., 2009)。但是,长达 400 km 的克拉通地壳如何俯冲 到足够的深度发生熔融,而不被大洋岩石圈拖曳至地幔,这是需要继续研究探讨的问题。

9)目前已知北美地台的变形时间和其西侧岩石的变形时间明显不一致(图2.7)。例如,在拿大落基山脉,地台岩石一直到桑托期(86.3~83.6 Ma)一坎帕期(83.6~72.1 Ma)(Santonian-Campanian)才发生变形,但是其西侧温德米尔高地的岩石在108 Ma之前就已经卷入到主逆冲作用中(Larson et al., 2006)。类似地,从内陆带到塞维尔褶皱-逆冲带的岩石保留了侏罗纪和晚白垩世两个强烈变形时期的证据(Camilleri et al., 1997),而北美地台的岩石并未显示侏罗纪变形的证据,但有阿普特期(126~113 Ma)—诺曼期(100~93.9 Ma)为主的向东薄皮逆冲作用(DeCelles, 2004)。



图 2.7 沿着北美大陆边缘两个直接相邻区域不匹配的变形年龄差异图解 在加拿大分区,温德米尔湖和珀塞尔带超群的岩石在 108 Ma 之前发生变形和逆冲,而在北美地台东 部仅仅几千米处的岩石,直到桑托期—坎帕期(Santonian-Campanian)都没有发生变形和逆冲。在大 盆地分区,北美地台的岩石主要在 124~105 Ma 之间发生变形和向东逆冲,而内陆带的岩石仅仅西 侧的显示出侏罗纪和白垩纪末变形。

现有弧后模型的缺陷证明:我们需要重新审视科迪勒拉山系造山作用这个概念。相对于40年前 科迪勒拉模型刚刚建立之时,现在我们对于板块构造塑造地球上部圈层的运动机制已经知之甚多,因 而出现新的认识自然不足为奇。随着新地质事实的发掘和新概念的发展形成,可以看出弧后模型中一 些部分是不正确的,需要修正,有些甚至是需要被抛弃的。下面所讨论的不是最终的(或最后的) 结果,而是一个暂时的尝试,就是将现有知识整合成为一个符合实际的模型,这一模型可以随着新资 料的搜集和分析,接受进一步验证和再完善。

### 第3章 分段造山带

北美科迪勒拉山系造山带的一个最主要特征就是具有分段性(King, 1966)。根据地质对比和不同阶段的工作积累, Hildebrand (2009)将该科迪勒拉从北至南划分成三个分段:加拿大、大盆地(Great Basin)和索诺兰沙漠分区(Sonoran;图 2.5)。北部前两段大致由奥罗菲诺(Orofino)断裂分隔,而大盆地和索诺兰沙漠则以具有左旋走滑性质的菲尼克斯(Phoenix)断裂为界。基底前渊的碎屑锆石剖面(Leier and Gehrels, 2011)很清楚地显示在奥罗菲诺断裂附近出现明显的间断,这进一步指示了该断裂为一个有意义的分段边界。一个主要的右行断层,可由沿内华达山脉和爱达荷岩基的亚特兰大圆形突出体复原,推测该断层位于蛇河(Snake River)平原内非常年轻的熔岩之下(Hildebrand, 2009)。另外一个被识别出的是阿拉斯加段,它主要位于加拿大地体之间一个明显的间断以西,现在大部分被年轻的沉积物覆盖,如育空-塔纳纳(Yukon-Tanana)和塞尔温盆地,以及阿拉斯加地体北部和西部的大部分区域(图 2.5)。

理解分区内的地质和构造演化对于阐明造山带的演化是极为关键的,因为分区之间不仅在岩浆和构造演化上有大的差别(Armstrong, 1974),大量证据表明沿着造山带地体发生过显著的纬向迁移,这就存在这种可能,即一个地体在一段时间内位于某一个分区,而在另一段时间内迁移到了另外一个分区,而且因为沿着相邻分区的走向,变形样式或岩浆作用类型等可能发生显著的改变(Oldow et al., 1989)。海岸深成杂岩提供了一个很好的例子,因为它在 80 Ma 的时候位于索诺兰沙漠分区,但是在 58 Ma 的时候就完全处于加拿大分区了。同样,加拿大分区东部主要为落基山褶皱-逆冲带的薄皮逆冲构造样式(Price, 1981),它在时间上与大盆地分区的拉勒米厚皮变形是一致的。因此,重建和建立模型如果不考虑这些方面是不可能获得成功的。

在这里,我首先简要回顾一下位于加拿大和大盆地分区内的北美大陆或劳伦系和被动大陆边缘 等,然后描述单个分区的地质特征,最后基于它们的起因和最终的汇聚,提出一个明显合理的、可被 检验的模型。

### 第4章 劳伦被动陆缘

最古老的、通过热沉降沉积在北美克拉通(Bond and Kominz, 1984)上的被动大陆边缘岩石为 早寒武世石英硅质碎屑岩,该套岩石上覆沉积岩为大陆架-大陆斜坡相碳酸盐岩,上覆碳酸盐岩向西 发生突变,逐渐相变为稀薄的沉积盆地相页岩(Rigo, 1968; Stewart, 1970)。古生代碳酸盐岩从大 陆架向盆地相转变的过程可以在现今的怀俄明州扇形地背斜轴(怀俄明凸起)观察到,该区邻近犹 他州-怀俄明州边界(Peterson, 1977; Rose, 1977; Doelling, 1980; Palmer and Hintze, 1992)的塞维 尔褶皱-逆冲带东部(图 4.1),在加拿大落基山的主山脉亦可以观察到(Cook, 1970; Aitken, 1971)。底部的寒武纪砂岩向西粒度渐渐变细,并在大陆架边缘相变为页岩(Oriel and Armstrong, 1971; Middleton, 2001)。整体而言,从地台到隆升过渡带的岩石或多或少地在原地得以保留,但也 有例外,比如局部出现的向东海侵,以及在犹他州中部和东部出现的宾夕法尼亚纪隆起和凹陷,持续 时间至少从寒武纪延续到侏罗纪(Hansen, 1976; Koch, 1976; Rose, 1977; Blakey, 2008)。大陆架-大陆斜坡的转变带被一些学者称之为瓦萨奇(Wasatch)脊线(Hintze, 1988; Poole et al., 1992),但 这种叫法并不严谨,因为脊线是岩石圈拉伸过程中距离陆地最近的那个点,并且大陆架边缘很可能标 志着上地壳伸展过程(脆性断裂)中距离陆地最近的点。下地壳的伸展很可能在更远的西端。

在寒武纪期间,北美大陆是边缘有碳酸盐岩镶边的台地,并且从地形上没有证据显示较高的 地体将大量的碎屑带出碳酸盐岩台地。因此,北美大陆寒武纪的盆地相很可能是相当饥饿的("饥 饿盆地")。对于北美大陆西部而言,其整体的沉积相框架应该是一个突然出现的碳酸盐岩边缘, 同时带有一个狭窄的大陆斜坡相碎屑扇状体和一个薄的、少沉积物的、以等深流沉积为主的隆 起相。

位于西部的远离地台的岩石,出现另外一个面向西的地台,Hildebrand (2009)将其称之为安特 勒边界(图 2.1 和 4.1)。这个地台的岩石被罗伯茨山(Roberts)和宝山(Golconda)外来体逆掩推 覆。它的边缘要比北美大陆的边缘老,因为它形成于新元古代,并包含一个早寒武世的碳酸盐岩滩, 以出现大量的古杯动物门(Archeocyathids)化石和独特的鲕粒岩层及内碎屑灰岩岩层为典型特征 (Fritz, 1975; Pope and Sears, 1997)。面向西的古生代安特勒地台边缘相在内华达州和加利福尼亚州 均有完好的地质记录(图 2.1),可以肯定的是:安特勒地台和塞维尔地体是明显不同的(Kepper, 1981; McCollum and McCollum, 1984; Heck and Speed, 1987; Montañez and Osleger, 1996; Morrow and Sandberg, 2008; Sheehan, 1986; Harris and Sheehan, 1998; Stevens et al., 1998; Stevens and Stone, 2007)。这并不意味着安特勒地台边缘相最初并不位于北美西部,只是如果是这样的话,它是从边缘 分离出来,并且相对于现在的位置来说是外来的。

在加拿大的科迪勒拉山,那里的地质历史与美国类似。成熟的浅海相碎屑岩构成一个向西加厚的 楔形体不整合在克拉通基底之上(图 4.2)。碎屑岩的最下部被两个边缘平行的岩石相覆盖:一个内 侧的碳酸盐岩滩和一个外侧的大陆斜坡相到盆地相岩石,两者被一个称为踢马圈(Kicking Horse Rim)的海藻礁杂岩分隔(Aitken,1971)。大陆架到大陆斜坡相的过渡出现在现今的主山脉(图 4.3),该区靠近阿尔伯塔-不列颠哥伦比亚省(Alberta-British Columbia)边界,在那里,中寒武世的 碳酸盐岩台地尖灭成为泥质的斜坡相到盆地相沉积,这些沉积中最为著名的堆积就是博格斯 (Burgess)页岩(Cook,1970; Price and Mountjoy,1970; McIlreath,1977)。发生相变的地方也是断 层作用和重力垮塌陡坡的位置,相变标志着褶皱的透入性应变和样式的变化(Dahlstrom,1977; Stewart et al., 1993)。



图 4.1 墨西哥西北部、美国西部和加拿大最西南部地质简图

说明本文所讨论的各种构造要素。在逆冲断层带虚线标记的是外来体的大致东界,注意,那也是缝合线。ATL—亚特兰大裂片 (Atlanta lobe); bb—巴卡盆地(Baca Basin); bb—博尔德岩基(Boulder batholith); BC—比特鲁特杂岩(Bitterroot complex); bh— 黑山 (Black Hills); BIT—比特鲁特裂片 (Bitterroot lobe); bm—蓝山地体 (Blue Mountains terranes); br—黑石沙漠 (Black Rock Desert); c-瀑布核心 (Cascades core); cc--克利尔沃特杂岩 (Clearwater complex); cmb--疯狂山盆地 (Crazy Mountains Basin); cn—查尔斯顿-尼波山凸 (Charleston-Nebo salient); d—死亡谷 (Death Valley); db—丹佛盆地 (Denver Basin); ev—埃尔克霍恩 火山岩 (Elkhorn volcanics); F—炉溪断层 (Furnace Creek fault); fc—法国人岩帽 (Frenchman Cap); GA—宝山外来体 (Golconda allochthon); grb—绿河盆地(Green River Basin); hs—海伦娜岬(Helena salient); k—水壶杂岩或穹窿(Kettle complex or dome)。 lftb--卢宁-Fencemaker 逆冲断层带 (Luning-Fencemaker thrust belt); mb--麦考伊山组 (McCoy Mountains formation); mftb--大玛丽 亚褶皱-逆冲断层带 (Big Maria fold-thrust belt); mmt—骡子山逆冲断层系统 (Mule Mountains thrust system); msms—莫哈韦沙漠-索诺拉大型平移断层 (Mojave-Sonora megashear); ns—内华达山脉北部 (northern Sierra Nevada); oc—奥肯那根杂岩或穹窿 (Okanagan complex or dome); pn-松子地块 (Pine Nut block); pr-牧师河杂岩 (Priest River complex); prb-粉河盆地 (Powder River basin); R—兰德片岩 (Rand schist); rb—拉顿盆地 (Raton basin); rr—阿尔比恩-筏河-松鸡溪山脉 (Albion-Raft River-Grouse Creek Ranges); s-春山 (Spring Mountains); SFTB-塞维尔褶皱-逆冲断层带 (Sevier fold-thrust belt); sjb-圣胡安盆地 (San Juan Basin); SJIFTB-圣胡安岛褶皱-逆冲断层带 (San Juan Islands fold-thrust belt); SOB-舒斯瓦普-奥米尼卡南部带 (Shuswapsouthern Omineca belt); sw-Swakane 片麻岩 (Swakane gneiss); ub-因塔盆地 (Uinta Basin); UCSB-加州大学圣塔芭芭拉分校 (University of California at Santa Barbara); v-瓦尔哈拉杂岩或穹窿 (Valhalla complex or dome)。vp-比斯卡伊诺半岛 (Vizcaino Peninsula); wb—瓦沙基盆地 (Washakie Basin); wr—风河盆地 (Wind River Basin)。据 Nourse (2002), Orocopia 带复原到前圣安地 列亚斯(San Andreas)断层的形态,但现代海岸线留下来作参考。还需要注意的是更年轻的盆岭省内的伸展没有复原。在这里及 后面的图,我用科罗拉多高原地貌图,科迪勒拉核杂岩的伸展方向据 Wust (1986)







图 4.3 地质剖面图 (据 Collom 等 (2009)修改) 显示加拿大南部落基山脉主要山岭的踢马圈 (Kicking Horse Rim)地台、斜坡和盆地相岩石之间的关系。 大教堂 (Cathedral)组和艾尔登 (Eldon)组在西部陡坡均有约 200 m 的落差

在加拿大,与安特勒大陆架对等的地质体是卡斯尔地体,它位于远离大陆架盆地沉积岩的西侧 (Johnston, 2008)。就像它们在大盆地地区的对等地台一样,卡斯尔地台岩石由一个浅水相、碳酸盐 岩-硅质碎屑岩混合的岩石构成,年龄的跨度从新元古界一直持续到了中生代。通过大量的证据,包 括生物分区、基底岩石年龄和中生代构造演化的对比,Johnston (2008)提出卡斯尔地台岩石是属于 一个更大的带状大陆的一部分,该大陆相对于北美克拉通,在白垩纪以前是外来的。

### 第5章 大盆地分区

### 5.1 塞维尔褶皱-逆冲带

构成科迪勒拉造山带东部的一个重要要素是塞维尔褶皱-逆冲带 (Sevier Fold-Thrust Belt), 它是 一个薄皮变形带,该带分布范围从南部的内华达泉山(Spring Mt.)一直延伸至北部的奥罗菲诺断层 (图 2.5 和图 4.1)。在这条带内,北美被动大陆边缘的岩石从基底拆离,然后被向东逆冲到一个基底 滑脱带之上,并以位于边缘的寒武纪页岩剖面部分最具代表性 (Armstrong and Oriel, 1965; Armstrong, 1968; Allmendinger, 1992; Burchfi el et al., 1974a, 1974b, 1992, 1998; DeCelles and Coogan, 2006)。尽管大多数的学者认为该逆冲带是北美大陆下部大洋岩石圈向东俯冲的远端和弧后表现 形式,而 Hildebrand (2009)提出该逆冲带是一个典型的由碰撞引起的逆冲带,它的形成是由于北美 西缘被拉到先前拼合的卢比亚超级地体之下 (图 5.1)。在结构上,北美地台阶梯之上的逆冲岩片是 巨大的逆冲岩席,厚度达28km,宽度和长度可达数百千米(Sears, 1988; DeCelles and Coogan, 2006; Fermor and Moffat, 1992)。这些包括巨型逆冲岩席,在大盆地地区包含超过7~10 km 厚的新元 古界-寒武系沉积岩 (Christie-Blick, 1982, 1997; DeCelles and Coogan, 2006), 往北, 既位于加拿大 分区又位于大盆地分区的巨型外来体,包含了贝尔特-珀塞尔-温德米尔 (Belt-Purcell-Windermere) 超群的岩石,这些岩石被携带到了路易斯(Lewis)-埃尔多拉多(Eldorado)-Hoadley 逆冲杂岩中 (Cook and van der Velden, 1995; Mudge and Earhart, 1980; Sears, 2001)。巨型逆冲岩席和贝尔特--珀 塞尔-温德米尔外来体均没有在地层上有与北美克拉通对等的沉积序列;因此 Hildebrand (2009)认 为它们两个都是外来的产物。



图 5.1 大盆地分区塞维尔造山运动板块模型 简图说明 Hildebrand (2009 年)的大盆地分区塞维尔 (Sevier)造山运动板块模型。类似的,北 美大陆向西俯冲到卢比亚 (Rubian)带状大陆之下,这里还预想拉勒米事件也发生在加拿大分区 和索诺兰沙漠分区

碰撞时间的沉积记录在不同地点的起始时间不同,但是几乎所有的地方,碰撞时间均晚于砂砾岩和砾岩的沉积时间,这些砂砾岩和砾岩中包含了各种各样的沉积碎屑,如燧石岩、石英岩、灰岩和粉砂岩——这些沉积碎屑来自北美地台内古老岩石的剥蚀搬运,然后向东分散迁移,最终在一个区域不整合和一个钙结砾岩-硅结砾岩古土壤层之上形成另外一个薄层,其中的古土壤层是在莫里森(Morrison)组和随后的侧向同沉积岩中形成的(Schultheis and Mountjoy, 1978; Leckie and Smith, 1992; Heller and Paola, 1989; Yingling and Heller, 1992; Currie, 2002; Ross et al., 2005; Zaleha and Wiesemann, 2005; Zaleha, 2006; Roca and Nadon, 2007; Greenhalgh and Britt, 2007)。这些砂砾岩在不同地

区命名不同,如 Cadomin、Kootenai、Lakota、Cloverly、Ephraim、Buckhorn、Pryor等,并且它们分布 广泛(Heller et al., 2003),出现在边缘附近上下浮动(图 5.2)。在大盆地分区,砂砾岩被阿尔布期 和森诺曼期海相泥岩和粉砂岩覆盖,该海相沉积标志着塞维尔前渊盆地的首次沉积记录,区域上称之 为西部内陆(Western Interior)盆地(Kauffman, 1977; Hunt et al., 2011)。



图 5.2 白垩纪砾石在科迪勒拉前渊底部的分布示意图 (图改编自 Heller and paola, 1989) 通常认为这些砾石包含源自北美地台沉积单元的沉积岩碎屑。香草拱形(Sweetgrass arch)似乎最初分离于盆地的两 个不同的部分,在不同的时间发展演化。Hildebrand(2009)认为砾石源自北美地台的剥蚀,由于地台慢慢向上拱 并越过外缘隆起到向西倾的海沟

从加拿大分区卡多明(Cadomin)组最新搜集和分析的碎屑锆石年龄数据来看,卡多明组与大盆 地分区内的同等沉积岩相比,前者要比后者年轻 30 Ma 或者更多(Leier and Gehrels, 2011)。上述两 个地区被香草弧分隔,而白垩纪时期,香草弧在北美克拉通是非常活跃的(Lorentz, 1982; Podruski, 1988)。

Hildebrand (2009) 认为不整合和上覆砾岩标志着北美克拉通运移到海沟的外侧前缘隆起之上 (Currie, 1998)。当大陆通过构造凸起的时候 (McAdoo et al., 1978; Forsyth, 1980; Jacobi, 1981; Stockmal et al., 1986; Yu and Chou, 2001), 它会向上弯曲,同时它的被动边缘将会因侵蚀而出露,因此古生代和中生代沉积碎屑岩从隆起区被剥蚀到周围的盆地内。这种类型的砾岩与特提斯喜马拉雅的砾岩在碎屑岩性和地层背景上具有相似性,而特提斯喜马拉雅的砾岩也被认为是弯曲隆起和地台通过 海沟外侧隆起的标志 (Zhang et al., 2012)。因此,在 Hildebrand (2009) 的模型中,卢比亚与北美大

陆的碰撞是由北美克拉通前缘向西俯冲引起的。另一方面, Heller 等(2003)认为砾岩在大陆倾斜的 时候会由于动力地形学的原因而发生分散传播,这与前面所述并不矛盾。

根据砂砾岩上方的火山灰床锆石 U-Pb 数据(Greenhalgh and Britt, 2007)、碎屑锆石数据(Britt et al., 2007)和一个119.4±2.6 Ma的含铀碳酸盐岩 U-Pb 年龄(Ludvigson et al., 2010),说明大 盆地分区内的逆冲作用时间可以限制在近124~120 Ma(图5.3)。薄皮构造变形结束时间以前渊盆地 深海相沉积向局部的和沉积学上孤立的非海成盆地相的转变为标志(图2.5、图4.1和图5.4),后者 是拉勒米厚皮构造变形的典型特征,时间被约束在80 Ma和70 Ma之间(Dickinson et al., 1988; Raynolds and Johnson, 2003; Cather, 2004)。在犹他州, 鹿角(Buckhorn)砾岩底部有一个碎屑锆石 剖面,这个剖面与科罗拉多高原的古生代—中生代剖面几乎没有什么差别,通过前渊沉积剖面碎屑锆 石采集限定了一个倒转的时间相,记录了来自塞维尔褶皱-逆冲带西端的外来地层完整的去顶沉积过 程(Lawton et al., 2010; Hunt et al., 2011)。



图 5.3 科迪勒拉前渊的下部岩石与源自地台的砾石地层对比图 (据 Heller and Paola (1989)修改) 展示美国西部 125 Ma 的砾石和加拿大北部香草拱形的年轻砾石

在爱达荷州南东部, DeCelles et al. (1993) 指出:在相对细粒的贝希勒(Bechler)组内很可能存在阿普第期的粗砾岩,它代表了来自米德巴黎(Meade-Paris)逆冲断层系统前锋的物质,而该逆冲断层系统是将厚的新元古界—早古生代大型逆冲岩席带到北美地台边缘之上的主断层(图 5.5)。 在犹他州中部,帕旺(Pavant)逆冲断层是在构造上处于最低位的断层,它携带大型逆冲岩席的岩石将其带到北美地台之上(DeCelles and Coogan, 2006),有一种解释就是这些粗碎屑构成了阿普特期-阿尔布期的圣节组(San Pitch Formation)(DeCelles et al., 1995)。

一些学者(DeCelles and Currie, 1996; DeCelles, 2004; Fuentes et al., 2010)认为加拿大弗尼盆地和美国西部莫里森组内的侏罗纪岩石是构造凸起后沉积的,该沉积与 Royse(1993b)最早提出的 "Phantom 前渊凹陷"相关。Hildebrand(2009)认识到这种提法只是一种可能,但是无法得到验证,







显示科迪勒拉山系造山运动的拉勒米期间各个盆地的形成、演化和终结时代。在塞维尔阶段的薄皮(thin-skinned Sevier)变形时期形成的较老且更连续的前渊盆地被瓦解;在北美大陆和卢比亚带状大陆之间的拉勒米碰撞期间, 拉勒米阶段的厚皮(thick-skinned)变形形成较年轻的孤立盆地。注意:所有盆地都形成于麦斯里希特期(Maastrichtian)



#### 图 5.5 落基山脉和塞维尔褶皱-逆冲带剖面对比图

展示最西端逆冲岩席的相似性,它们都携带大量较厚的前寒武纪—古生代岩石,典型的逆冲岩席遍及整个科迪勒拉山系,它 们如果来自北美大陆的最外缘,将要求有庞大的地壳斜坡。本书认为相对于北美大陆而言,这些大型逆冲岩席的岩石是外来 的。罗布森山-碧玉地区(Mount Robson-Jasper area)据 Mountjoy(1979);卡尔加里(Calgary)地区据 Price and Fermor (1985);蒙大拿西北部据 Fuentes等(2012)简化;冰川国家公园(加拿大)据 Price(1962)修改;爱达荷-怀俄 明剖面据 DeCelles等(1993)修改 因而是无法令人满意的,因为前渊沉积的痕迹早已消失得无影无踪。进一步讲,在莫里森沉积与前渊 沉积之间存在一个 25 Ma 的间断,这一间断并没有被充分考虑进去,在这一间隔期,莫里森盆地是倒 转的和被分隔的,并伴有 50 m 的局部起伏和局部骨层堆积 (Eberth et al., 2006);在组中含少量的三 叠纪—侏罗纪碎屑锆石 (Fuentes et al., 2009),很可能起源于经过大气运移的凝灰岩单元的再循环, 这种类型凝灰岩在整个盆地分布很广 (Kowallis et al., 1998, 2007)。

在大盆地分区内部,一个发育于寒武纪页岩中的主逆冲断层带,将北美地台的岩石向东运移至克 拉通之上,同时也将所谓的巨型岩片内厚层的新元古界沉积序列带到地台岩石之上(DeCelles,2004; DeCelles and Coogan,2006),在蛇河(Snake River)平原的南东侧塞维尔带形成一个广阔的突出部 分,该部分向南延伸,穿过爱达荷、怀俄明州最西部和犹他州北东部(图 4.1)。该分区至少包含八 个主逆冲断层系统,这也是褶皱-逆冲带在美国唯一大面积出露的地区,在那里,北美地台边缘并没 有被逆冲岩片推翻倒置,该逆冲岩片携带安特勒地台的岩石,逆冲岩片本身的岩性为厚层新元古界陆 源碎屑岩(Peterson,1977; Rose,1977; Palmer and Hintze,1992)。西部的逆冲断层,如米德 (Meade)和巴黎(Paris)逆冲断层,携带厚层的元古宇沉积岩,而其他更多向东的逆冲断层则起源 于寒武纪页岩内的拆离断层(Armstrong and Cressman, 1963; Armstrong and Oriel, 1965; Royse et al., 1975; Lamerson, 1982; Royse, 1993a)。

其中一个叫做奥格登(Ogden)逆冲断层系统,是一个具有古元古界结晶基底的背斜型双重构造,该结晶基底构成现在的法明顿杂岩(Bryant, 1984; Yonkee, 1992; Yonkee et al., 1989, 2003; Andreasen et al., 2011)。结晶基底之下的逆冲断层中的一部分包含了怀俄明克拉通的太古宇岩石(Foster et al., 2006)。法明顿杂岩出露在犹他州北东部的大陆架边缘西侧(Rose, 1977),在构造上位于Paris 逆冲断层的下部。古元古界的岩石带很可能向北延伸至爱达荷,在那里,古元古界的结晶基底出露在舱药湖(Cabin-Medicine Lake)系统内,该系统就在爱达荷岩基的东侧(Skipp and Hait, 1977; Skipp, 1987)和蒙大拿州南西侧的滕多伊(Tendoy)山脉。根据 Sr 同位素(Armstrong et al., 1977; Fleck and Criss, 1985)、红宝石(Ruby)和附近山脉出露的基底构造窗以及捕虏体的证据(Evans et al., 2002),对逆冲带内的缩短量进行了古地理复原,结果显示:法明顿杂岩内的岩石复原到北美克拉通地壳的西侧,结合同位素和地质证据,太古宇的基底很可能只延伸到克拉通的边缘(Hanan et al., 2008),这只是相对于现今的位置,表明法明顿峡谷(Farmington Canyon)杂岩是外来的。

在犹他州普洛佛 (Provo) 南侧, 是另外一个向东的再进入部分, 称为查尔斯顿尼波 (Charleston – Nebo) 凸起 (图 4.1), 在那里, 查尔斯顿尼博系统的逆冲断层携带了一个大的、发生倒转的、几 乎伏卧的背斜, 该背斜由厚层的宾夕法尼亚纪—二叠纪沉积岩石构成, 这些岩石现在不在北美地台上, 同时背斜内还有少量的古元古界结晶基底岩石, 即称之为 Santequin 杂岩 (Tucker, 1983)。逆冲带更远的西部, 例如绵羊石 (Sheeprock), 携带了厚层的前寒武系碎屑岩层 (Christie – Blick, 1982, 1983, 1997; Rodgers, 1989)。

逆冲带东部的大片地区被同造山沉积岩埋藏,而瓦萨奇(Wasatch)山前以西的整个地区被新生 代正断层强烈地肢解。

在犹他州中南部存在四个主要的逆冲断层系统,犹他州最西部有一个峡谷山脉-哇哇-帕旺 (Canyon Range-Wah Wah-Pavant)系统,这些逆冲断层既携带了厚达 4~10 km 的以硅质碎屑岩为主 的新元古界岩石,又携带了厚达 12 km 的古生代地层,而在北美克拉通地台的东侧只有约 1.5 km 的 厚度(Hintze, 1988)。峡谷山脉逆冲断层即为 DeCelles (2004)和 DeCelles and Coogan (2006)提出 的那种"大逆冲型"(megathrust)。

像桑塔金(Santaquin)海湾以北一样,犹他州中南部的褶皱-逆冲断层前 50 km 也主要被造山沉 积物埋藏(DeCelles, 2004),后来被更年轻的正断层破坏。这一地区东部的断层同样是根植于寒武纪 页岩内的拆离断层,这一特点与北部类似(Lawton et al., 1997; DeCelles et al., 1995)。

在拉斯维加斯地区,存在若干主要的逆冲断层系统。在构造上最低的威尔逊峭壁 (Wilson Cliffs)

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逆冲断层将寒武纪碳酸盐岩和碎屑岩置于北美地台的风成阿兹特克(Aztec)砂岩之上(Burchfi el et al., 1974a, 1998)。再往西,新元古界的厚层沉积岩,即称之为帕朗(Pahrump)群,连同它们的古生代盖层在一系列逆冲断层内不整合覆盖在基底之上(Burchfi el et al., 1974a, 1974b; Brady et al., 2000; Snow, 1992; Wernicke et al., 1988)。

在大盆地分区,大型逆冲岩席和它们本身的新元古界厚层沉积岩石在北美地台上的就位主要发生 在阿普特期—森诺曼期(124~94 Ma),这一推理是基于对粗粒沉积堆积相的研究得出的,这些粗粒 沉积堆积相或与逆冲断层部分重叠,例如覆盖于峡谷山脉(Canyon Range)逆冲断层上的砾岩(De-Celles and Coogan, 2006; Lawton et al., 2007), 或与逆冲断层作用同时发生沉积, 例如贝希勒 (Bechler)砾岩,该砾岩被米德(Meade)逆冲断层最早的逆冲作用超越(DeCelles et al., 1993)。在 犹他州内,被认为是邻近前渊沉积的岩石被包含在印第安诺拉(Indianola)群和雪松(Cedar)山组 内 (DeCelles and Coogan, 2006; Hunt et al., 2011, and references therein)。在雪松山组最上段,即称 之为"Mussentuchit",经过放射性同位素测试定年,获得其为森诺曼期最早期产物,该段直接位于更 细的和宽阔盆地相的达科他(Dakota)组之下(Cifelli et al., 1997; Garrison et al., 2007; Biek et al., 2009)。峡谷山脉下部砾岩定年工作还非常薄弱,但是 Lawton 等(2007)认为一套碳酸盐岩碎屑特别 富集的独特砾岩与圣节(San Pitch)组的最下段建立了联系,圣节组位于峡谷山脉的东面,它包含阿 尔布期中期到晚期的孢粉化石 (Sprinkel et al., 1999)。因此,如果峡谷山脉和圣节组之间的相关性 是正确的话,大盆地分区内的大型逆冲岩席的主体就位时间就发生在阿尔布期-森诺曼期,即125~ 94 Ma,但是也可能不晚于中阿尔布期(近 105 Ma)。达科他上部沉积岩包含的碎屑锆石存在三个峰 期, 分别为 121、116 和 110 Ma (Ludvigson et al., 2010), 这与周围老的沉积物反映的源区显然不同, 早前提到过,周围老的沉积物包含一个倒转褶皱的碎屑锆石,该锆石来自于大型逆冲岩席内的新元古 界和下古生界岩石 (Lawton et al., 2010; Hunt et al., 2011)。

#### 5.2 内陆带

大盆地分区内位于大型逆冲岩席正西的新元古界沉积岩就是内陆带(图 4.1)。自从 Armstrong (1968) 识别出这条带之后,它的成因已被证实令人难以捉摸,因为它包含:多期变形和低变质沉积 岩、高级变质岩、若干不同年龄的结晶基底岩石,以及偏铝质和过铝质侵入体——上述所有的这些岩 石都被逆冲断层和正断层错断,并发生显著的位移,它的东部和西部边界多少有些模糊不清,因为至 少发生过两个主要时期的张裂和不同程度的剥露;但是一般而言,内陆带是向北延伸的条带,以含有 古新世—中新世的变质核杂岩、同时有侏罗纪和白垩纪的逆冲断层和变质作用、主要为向西的侏罗纪 褶皱、以及通常有稀少的侏罗纪—白垩纪的深成岩体等为特征。

内陆带的岩石出露在古新世—中新世变质核杂岩中,在内华达东北部、犹他州西北部和爱达荷州 南部的阿尔比恩(Albion,印第安纳州)、筏河(Raft River)和松鸡溪(Grouse Creek)等地被称为 是出露最好的典型区域(Armstrong, 1968; Snoke, 1980; Howard, 1980; Todd, 1980; Snoke and Miller, 1988; Wells, 1992)。因为在古新世—中新世被加厚的和热的内陆带伸展垮塌以及更为年轻的 盆岭伸展作用,现今内陆带内的构造地形是显而易见的。

阿尔比恩-筏河-松鸡溪山脉内的岩石可以划分为一个原地岩系,它由太古宙结晶基底构成,其上被石英岩和泥质片岩薄层不整合覆盖(Compton, 1972),然后又被古生代的绿片岩-角闪岩相变质沉积岩构造覆盖(Wells et al., 1997)。在这个原地岩系内,通过小尺度的构造研究,变形程度从逆冲断层下盘往下逐渐变弱,然而变质作用是滞后的,或者可能是与逆冲作用同时进行的(Compton, 1980; Miller, 1980; Snoke and Miller, 1988)。现在出现在上覆外来体中的早白垩世和侏罗纪花岗岩,在原地岩系内未被描述过。原地岩系的太古宙结晶基底被认为是以怀俄明为代表的北美克拉通(Miller, 1980; Snoke and Miller, 1988)。

这个地区的岩石通常包含两个变形幕次,第一次发生在晚侏罗世,第二次发生在晚白垩世(Camilleri et al., 1997; McGrew et al., 2000)。侏罗纪深成岩体产出在外来体内,侏罗纪的岩浆作用与小 褶皱形成、逆冲作用,以及侵入体周围的局部变质晕的形成,在时间上存在重叠 (Camilleri et al., 1997)。

第二次变形和变质事件更为强烈和普遍。这一时期逆冲断层作用的收缩脉冲至少导致了 70 km 的 地壳缩短,以及 30 km 的地壳加厚 (Camilleri et al., 1997)。在这一地区的变质组合指示了深部埋藏 和变质作用可能在晚侏罗世就已经开始,随后在大约 85 Ma 的晚白垩世达到变质峰期,变质峰期温度 为 800℃,压力大于 9 kb,随后出现一个快速隆升 (McGrew et al., 2000)。那个时候,当岩浆岩侵入 和开始跨越伸展拆离构造时 (Miller et al., 1987; Camilleri, 1992),大部分剥露作用已在中新世基本 完成。

内华达南西部的红宝石-东洪堡特(Ruby Range-East Humboldt)山脉,同样包含两期中生代造 山事件的岩石记录:①153 Ma 的深成岩侵位、多期褶皱作用和高角闪岩相的变质作用;②晚白垩世 混合岩化作用、变质作用和变形作用(Snoke and Miller, 1988; Hudec, 1992; McGrew et al., 2000)。 白垩纪混合岩化的高角闪岩相岩石发生构造叠加,并包含一个伏卧的等倾褶皱,核部为太古宙基底岩 石,上覆一套新元古界到密西西比纪沉积岩,该套沉积岩可能位于元古宙片麻岩之上(Howard et al., 1979; Lush et al., 1988; McGrew et al., 2000)。混合岩化作用和推覆体的就位是同步的,均在 84.8 ± 2.8 Ma,形成 750~800℃和9~10 kb 的变质峰期(Hodges et al., 1992; McGrew et al., 2000)。剥露作 用和隆升的初始时间尚未被精确限定,但是剥露作用持续范围在 63~50 Ma 之间(Snoke and Miller, 1988)。

红宝石山脉以南位于块状的蛇山脉 (图 4.1),它在构造上明显简单一些,但是仍然存在两期深成岩作用和变质作用的证据,一次在近 160 Ma,另外一次的变质作用峰期为 79 Ma,初始剥露阶段在 57~50 Ma 之间 (Snoke and Miller, 1988; McGrew et al., 2000)。在蛇山脉和东洪堡特-红宝石山脉中, 两者逆冲带内的大型逆冲岩席是相似的,均为外来的新元古代和古生代层序地层。

内陆带向南延伸进入死亡谷(Death Valley)地区,内陆带岩石在白垩纪(91.5±1.4 Ma)发生变质,变质峰期的温度和压力分别为近 620~680℃和 7~9 kb,随后发生了晚白垩世或三叠纪早期的伸展,这在内陆带的其他地区也比较典型(Hodges and Walker, 1990, 1992; Applegate and Hodges, 1995; Mattinson et al., 2007)。

#### 5.3 内华达中部

大型逆冲板片和内陆带的西部、通常认为的外来构造拼合体内,以及最东段的岩石组合等,是位 于安特勒大陆架上的地台相早寒武世到泥盆纪灰岩和硅质碎屑岩 (Poole et al., 1977)。地台的岩石被 其他地体逆掩覆盖,而这些所谓的其他地体相对于北美来说是外来体 (Silberling et al., 1992; Oldow et al., 1989)。罗伯茨山外来体的岩石 (图 2.5),由一个外来叠加构造杂岩组成,岩石组成为寒武纪 一泥盆纪粉砂岩、燧石、泥板岩、重晶石和镁铁质火山岩 (图 5.6),这些岩石是在安特勒造山过程 中在卢比亚边缘就位的,而安特勒造山作用的时间发生在晚泥盆世到早密西西比世 (Merriam and Anderson, 1942; Smith and Ketner, 1968; Poole and Sandberg, 1977; Nilsen and Stewart, 1980; Johnson and Pendergast, 1981; Johnson and Visconti, 1992; E. L. Miller et al., 1992b)。在这些外来体就位过程中和 之后的时间段里,粗碎屑向东散布,并在碰撞前的安特勒大陆架之上形成一个碎屑楔形体 (Poole, 1974, 1977; Harbaugh and Dickinson, 1981; Speed and Sleep, 1982)。

罗伯茨山外来体的岩石也在爱达荷州蛇河平原北部的先锋(Pioneer)山脉出露(Wilson et al., 1994; Link et al., 1996),并可能向北延伸至加拿大。一些学者(Turner et al., 1989; Smith and Gehrels, 1992a, 1992b; Smith et al., 1993; Root, 2001)认为古生代岩石和中晚泥盆世变形,以及伴随的一个造山前渊的形成演化均与安特勒造山作用有关,其中的造山带前渊从华盛顿州北部一直延伸到了加拿大北部的麦肯齐河(Mackenzie)三角洲。

位于罗伯茨山外来体顶部的是 Havallah 沉积序列岩石,属于另外一个外来的复杂变形岩石组合, 统称为宝山(Golconda)外来体(图 2.5),Havallah 沉积序列包含晚泥盆世到三叠纪最早期的燧石-

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展示罗伯茨山外来体的不同岩性。注意,碱性熔岩贯穿整个剖面,一个特征就是它与冷的、无岩浆活动的 被动边缘不相符。正如本文所讨论,岩石学特征比那些大陆边缘盆地更加典型

泥板岩沉积序列,并夹有枕状玄武岩的透镜体,这些岩石的就位发生在早三叠世的索诺玛 (Sonoman)造山运动期间 (Silberling and Roberts, 1962; Speed, 1977; Silberling, 1975; Dickinson et al., 1983)。不整合覆盖在罗伯茨山外来体之上,又位于宝山外来体之下的岩石,被 Dickinson 等 (1983)称之为安特勒上覆沉积序列,该套岩石在宝山外来体就位之前就发生了变形 (Cashman et al., 2011; Trexler et al., 2003, 2004)。

一个年龄为 250~248 Ma 的中酸性火山沉积序列,称为科帕托(Koipato)火山岩,不整合覆盖在 宝山外来体之上,他们的 Sr 和 Nd 同位素值显示其源区具有古元古界地壳的贡献,据此 Vetz (2011) 认为宝山外来体在火山岩喷发的时候就已经就位在罗伯茨山东侧岩石之上。短暂的岩浆大爆发可能是 俯冲板块断裂、随后发生宝山-罗伯茨山碰撞的表现形式。一个面向西的三叠纪碳酸盐岩台地覆盖于 火山岩之上,但是在东面,碎屑岩位于两者之间 (Oldow, 1984)。

### 5.4 白印优山脉

自印优山脉坐落在内华达山脉的东侧,形成了一个山脉系列,几乎同塞维尔山脉一样壮丽(图 2.5 和图 2.7),该山脉由一套 7 km 厚的古生代—中生代岩石构成(Stewart, 1970),该套岩石由于变质、变形而并没有明显向西延伸进入到内华达山脉,但已知向东延伸并在内华达州埃斯梅拉达县(Esmeralda County)出露,呈现一个弓状带(Albers and Stewart, 1972;图 4.1)。部分同时期的岩石 22

出露于死亡谷地区,两个地区的岩石都与白印优山脉内的岩石有关,但具体的岩石剖面却有很大差 异,要求每一个地层沉积间隔都有岩相变化:尽管出露良好,但两者之间的相变过渡却是缺失的。死 亡谷西部的岩石最初似乎位于现今的墨西哥索诺拉(Stewart et al., 1984, 2002; Stewart, 2005)。白印 优山板块至少在侏罗纪就明显地与邻近的内华达板块拼贴,因为那个时期的火山岩单元明显地沿走向 切割欧文斯山谷(Dunne and Walker, 1993)。类似地,死亡谷地区南东部的岩石在晚古生代发生拼 贴,因为二叠纪的逆冲断层(Snow, 1992)明显构成一个连续带,并和侏罗纪的深成岩体一样切割了 它们的接触带(Dunne et al., 1978)。

### 5.5 死亡谷

在内华达分区南部的死亡谷(图 4.1),帕朗(Pahrump)群的岩石和它们的结晶基底是外来的(Burchfi el et al., 1974a, 1974b; Brady et al., 2000; Snow, 1992)。木头峡谷(Wood Canyon)组的上覆埃迪卡拉纪—寒武纪沉积岩,含有的碎屑锆石年龄峰值为近 1.1 Ga,一个稍微年轻一些的扎布里斯基(Zabriskie)石英岩中含有大量 2.0~3.4 Ga 的锆石颗粒(Stewart et al., 2001)——而这些源区的年龄值在劳伦古大陆(Laurentia)西部是显著缺失的。在层序的偏上部分是中晚寒武世的博南扎王组(Bonanza King),该组的偏北侧是内华达州中部和犹他州最西部的安特勒大陆架的一部分(Kepper, 1981; McCollum and McCollum, 1984; Montañ ez and Osleger, 1996; Morrow and Sandberg, 2008)。在北美大陆架的西侧边缘,博南扎王组的岩石与科罗拉多高原内同时期的 Muav 组很难对应得上。博南扎王组的岩石和白印优山地区的岩石都发生褶皱,并被运移到逆冲断层上,而逆冲断层活动时间为近294~284 Ma 的二叠纪时期(Snow, 1992; Stevens et al., 1998; Stevens and Stone, 2007),但这一逆冲作用在北美地台内是未知事件。在靠近拉斯维加斯的泉山(Spring Mt.),博南扎王组的碳酸盐岩构造覆盖于阿芝特克(Aztec)砂岩之上,该砂岩是北美地台的一套侏罗纪风成砂岩(Burchfi el et al., 1998)。

该沉积层序再往上的地层,大致对应于博南扎王组的地台边缘相,是泥盆纪和晚奥陶世—志留纪的碳酸盐岩台地边缘相沉积岩(Sheehan, 1986; Harris and Sheehan, 1998; Morrow and Sandberg, 2008),这套面向西的大陆架相到边缘相转换沉积序列出露于内华达中东部和犹他州的北西缘,在那里它们作为安特勒大陆架的一部分出现在内陆带的西部(图 2.1)。类似地,宾夕法尼亚纪—早二叠世鸟泉(Bird Spring)碳酸盐岩陆棚边缘位于加州东部,并面向西(Stevens and Stone, 2007)。总的说来,大陆架边缘的位置、非北美克拉通的碎屑锆石和二叠纪的逆冲断层均表明死亡谷地区的结晶基底、帕朗(Pahrump)群和上覆古生代地层,相对于北美克拉通来说很可能是外来的。

我们现在跳过中生代大岩基和侏罗纪岛弧地体来描述内华达山脉西部的岩石,这样做的原因是为 了能够提供一个讨论白垩纪内华达岩基成因的框架,并首先描述包裹它们的围岩。

### 5.6 内华达山脉西部变质岩带

内华达山脉西部变质岩带是一条古生代—中生代混合岛弧和俯冲杂岩体的拼贴组合,其整体呈现为向西逐渐变年轻(Saleeby et al., 1989; Edelman, 1990),局部地区含有远程搬运的二叠纪麦克劳德(McCloud)和古地中海(Tethyan)的动物群(Miller, 1987),可划分为四个主要的构造带或地体(Day et al., 1985)。最古老和最东侧的为内华达北部地体(Coney et al., 1980),其下部由古生代嘘飞(Shoo Fly)杂岩组成,该杂岩包含向西的逆冲岩席,而这些岩席由奥陶纪到志留纪的沉积岩石和泥盆纪到奥陶纪的蛇绿混杂岩组成,并被385~364 Ma的鲍曼湖(Bowman)岩基穿切(Hanson et al., 1988; Harwood, 1992);在剖面上更高部分是向北东的逆冲岩席,包含古生代变质沉积岩,可能还有三个火山岩弧的火山岩和火山碎屑岩:①泥盆纪到宾夕法尼亚纪的泰勒斯维尔(Taylorsville,犹他州)火山沉积层序;②一个二叠纪的火山沉积层序;③中生代(变质岩带)东部的晚三叠世至中侏罗世的水壶石-汝拉山(Kettle Rock-Mount Jura)和塔特尔溪-水手峡谷(Tuttle Creek-Sailor Canyon)

火山沉积层序,变质岩带的西部边缘,被一个由较老混杂岩岩石构成的、向西倾的逆冲断层构造覆盖(Day et al., 1985; Christe and Hannah, 1990; Harwood, 1992; Christe, 2011)。中生代的岛弧序列是钙碱性火山岩带的延伸部分,该序列沿着内华达山展布,并通常被认为是由于大洋向东俯冲于北美大陆之下而形成(Burchfi el and Davis, 1972, 1975)。由于缺少合适的岛弧年龄、古地理因素和碎屑锆石分析, Wright 和 Wyld (2006)认为古生代的嘘飞杂岩曾经是冈瓦纳周缘的地体,然后向太平洋漂移,但是该杂岩何时与东部地体拼贴的问题尚未得到解决。

内华达地体北部的西侧是以断裂为界的羽毛河(Feather River)橄榄岩(图 5.7),即为一个古生 代一中生代缝合带(Edelman et al., 1989b),其岩石组成为发生变质和构造变形的橄榄岩、纯橄岩、 蛇纹岩和少量的变辉长岩、角闪岩和变质沉积岩,其中的一部分是由硬柱石蓝片岩(称为红蚂蚁 (Red Ant)片岩)组成的构造岩片(Mayfi eld and Day, 2000; Schweickert et al., 1980; Hietanen, 1981)。沿着橄榄岩带西缘产出的断层呈向东陡倾,并与卡拉维拉斯(Calaveras)混杂岩之上的橄榄 岩并列,但褶皱的向西倒转使得 Day et al. (1985)认识到边界断层亦发生过倒转,该边界断层向西 倾斜得如此之低以至于混杂岩向东逆冲到橄榄岩之上。沿着断层产出的片岩中角闪石的年龄变化范围 从 345~235 Ma,其中一些是高压变质年龄,因此变质底板可能形成于洋内的初始俯冲时期(Smart and Wakabayashi, 2009)。

羽毛河以西的橄榄岩位于中心带(图 5.7),该带为被肢解的等斜褶皱,以及各式各样的超基性岩、深成岩、火山岩和沉积岩等变质混合体(Day et al., 1985; Dilek, 1989)。除了沉积混杂岩和卡拉维拉斯岩石组合的破碎组,另有两个以断层为界的混杂岩,这两个混杂岩同时包含超基性岩和火山-侵入岩序列,被认为是岛弧产物,这两个混杂岩分别称之为:结合湖(Lake Combie)混杂岩和斯莱特溪(Slate Creek)混杂岩。

卡拉维拉斯(Calaveras)组合是一套构造岩石组合,其岩石组成为:玄武质到安山质枕状熔岩、 玻质碎屑岩、燧石岩、火山碎屑岩、千枚岩和二叠纪—石炭纪的大理岩,该组合位于羽毛河蛇绿岩和 前陆(Foothills)缝合带之下(Hietanen, 1981; Schweickert and Bogen, 1983; Hacker, 1993)。卡拉维 拉斯岩石组合内的灰岩包含麦克劳德动物群(Standlee and Nestell, 1985)。最近的碎屑锆石数据分析 结果表明混杂岩的年龄可能小于159~150 Ma,最年轻的碎屑锆石年龄峰值出现在岩石组合西部单元 (Van Guilder et al., 2010)。

侏罗纪的斯莱特溪 (Slate Creek) 杂岩,作为一个卡拉维拉斯杂岩以西以断层为界的单元出露 (图 5.7 和图 5.8),该杂岩主要位于小提琴溪 (Fiddle Creek)杂岩顶部,由三个假地层构成:①由 蛇纹岩基质混杂岩构成底部带,含有深成岩、火山岩和变沉积岩岩块;②一个由角闪石辉长岩、变质 辉绿岩和英云闪长岩岩石组合构成的深成岩中间间隔;③一个上部的火山岩单元,由隐晶质到辉石斑 晶的斑状绿岩、凝灰岩和局部火山碎屑岩组成 (Day et al., 1985; Edelman et al., 1989a, 1989b; Fagan et al., 2001),它在构造上位于小提琴溪杂岩之上 (图 5.7 和图 5.8),逆冲断层被 167 Ma 的斯 凯尔斯 (Scales) 深成岩体切割 (Day and Bickford, 2004)。变质深成岩和变质火山岩的年龄变化范围 从约 209→172 Ma,但也有一套较年轻的深成岩体,其年龄变化范围从 160→150 Ma (Edelman et al., 1989a, 1989b; Saleeby et al., 1989; Fagan et al., 2001; Day and Bickford, 2004)。

小提琴溪杂岩由两个不同岩石组合构成:①蛇绿混杂岩,含有大量的蛇绿岩岩块和被闪长质岩墙 切割的闪长岩;②燧石岩-泥岩单元,主要岩性为火山碎屑质砂岩和重力滑动沉积岩(滑塌岩),并 含有角闪岩、大理岩和极少量的枕状玄武岩等岩块(Dilek,1989; Edelman et al., 1989b)。该区的岩 石定年工作十分缺乏,但是燧石岩-泥岩单元的形成时代显然在中三叠世到近174 Ma 之间 (Hietanen, 1981; Hacker, 1993)。

结合湖(Lake Combie) 杂岩是另外一个以断层为界的侏罗纪杂岩带,与斯莱特溪杂岩类似且相关,因为它包含一个面理化和线理化的底部超基性岩石单元,构造上被辉长岩、石英闪长质侵入体和一个厚达5km的上部单元覆盖,由底部的基性熔岩流,向上渐变为以凝灰岩、角砾流和火山碎屑岩为主(Day et al., 1985)。



#### 图 5.7 内华达山脉和西部变质带地质简图

展示文中讨论的位置和地质特征。据 Irwin and Wooden (2001) 图修改,并补充 Bateman (1992)、Saleeby and Busby-Spera (1993)、Dunne 等 (1978) 和 Saleeby 等 (1978)的资料。

B--桦树溪 (Birch Creek) 岩体; BC-主教溪 (Bishop Creek) 顶垂体; Bi-Bendire 岩体; BL-鲍曼 (Bowman) 湖岩基; Bn-Bean 峡谷顶垂体; BP-边界峰岩体; BPC-大松树河 (Big Pine Creek) 顶垂体; By 一博伊登洞穴 (Boyden Cave) 顶垂 体; C-Concow 岩体; CG-粗粒金 (Coarse Gold) 顶垂体; Co-科利马岩体; DC-Dinkey Creek 顶垂体; dgo-魔鬼门 (Devil's Gate) 蛇绿岩; EMB-中生代东部带; GF-大森林; Guad-瓜达卢佩 (Guadalupe) 岩体; HC-Haypress Creek 岩 体; IZ-伊莎贝拉 (Isabella) 湖顶垂体; KC-克恩峡谷 (Kern Canyon) 顶垂体; KP-卡维亚峰 (Kaweah Peaks) 顶垂体; LKaR-卡维亚河 (Kaweah River) 下游顶垂体; LKR-国王河下游顶垂体; MB-巴克罗夫特山 (Mount Barcroft) 岩体; MC-大理石峡谷岩体; MG-戈达德山 (Mount Goddard) 顶垂体; MK-矿物国王 (Mineral King) 顶垂体; MM-莫里森山 (Mount Morrison) 顶垂体; Mmc-梅里马克 (Merrimac) 岩体; MT-塔利亚克山 (Mount Tallac) 顶垂体; OK-奥克里克溪 (Oak Creek) 顶垂体; PC-松河 (Pine Creek) 顶垂体; PF-帕普斯 (Papoose) 平地岩体; PG-皮诺格兰德 (Pino Grande) 岩体; Ph-松树山冈 (Pine Hill) 杂岩; PM-派尤特纪念碑 (Paiute Monument) 岩体; RR-里特山脉 (Ritter Range) 顶垂体; S-雪湖顶垂体; SF-Swede Flat 岩体; SHF-Sage Hen Flat 岩体; SL-鞍袋湖 (Saddlebag Lake) 顶垂体; SM-草莓矿山 (Strawberry Mine) 顶垂体; ST-斯坦达德 (Standard) 岩体; T-三分峰 (Triple Divide peak); TP-蒂哈查皮 (Tehachapi) 顶垂体; TR-图莱河 (Tule River) 顶垂体; YNP-约塞米蒂 (Yosemite) 国家公园; YV-Yokohl 峡谷顶垂体



图 5.8 变质带地质简图 (据 Day and Bickford (2004) 修改) 显示北西部和中心带以及西部变质带的地质特征

最西侧的地体称之为斯马特维尔(Smartville)杂岩,这是一个不完整的蛇绿岩,含有蛇纹岩化 的超基性岩、辉长岩、枕状玄武岩和席状岩墙,它们全部被一个1.5~2 km 的岛弧岩套覆盖,该岛弧 岩套的岩石组成为辉石安山质凝灰角砾岩、玄武质到安山质熔岩流和枕状熔岩,并含有少量的英安质 喷出岩、砂岩和砾岩(Xenophontos and Bond, 1978; Menzies et al., 1980),其中的火山岩通常被划分 为下部的拉斑质岩块和上部的中侏罗世钙碱性岩套,并被一个年龄为近 163 到 159 Ma 的伸展性席状 岩墙群侵入 (Beard and Day; 1987; Saleeby et al., 1989; Day and Bickford, 2004), 而席状岩墙被认为 是岛弧内发生伸展作用的产物 (Beard and Day, 1987; Dilek, 1989)。该杂岩形成了一个向东的逆冲 断层上盘,该上盘具有背斜的性质,并且该杂岩位于中心带混杂岩和破碎地层之上(Day et al., 1985; Moores, 2011,私人通信)。根据化石记录、地层叠覆律和一个获得年龄的凝灰岩,杂岩上部的火山岩属于牛津期-启莫里期的产物,比 157 Ma 的尤巴河(Yuba Rivers)深成岩体略老(Xenophontos, 1984; Saleeby et al., 1989; Day and Bickford, 2004),而尤巴河深成岩体沿着斯马特维尔与斯莱特溪-结合湖带(Slate creek-Lake Combie)和斯马特维尔杂岩的变质岩之间的逆冲断层产出(Bobbitt, 1982)。因此,斯马特维尔地块东侧发生碰撞的时间约为近 162 Ma (Day and Bickford, 2004),而岛弧岩浆作用一直持续到发生碰撞的时候。

再往南,沿着斯马特维尔杂岩的走向,很可能是属于同一个岛弧的一部分,是格外厚的侏罗纪岛 弧岩石——即为 Penon Blanco, Logtown Ridge 和 Jasper Points 组,它们位于蛇绿岩基底之上,岩性为 辉石斑岩,同时也属于一个上盘性质的背斜(Bogen, 1985)。这些岩石组合明显与一套稀少的相似年 龄为 170~160 Ma 橄榄岩-闪长岩侵入杂岩相关,沿着内华达山前陆西部出露,并被解释为是岛弧岩 浆作用的产物(Snkoke et al., 1992;译者注:应该是一套阿拉斯加型杂岩体)。这些岩石被 153~151 Ma 的瓜达卢佩(Guadalupe,加利福尼亚州)火成岩杂岩切割(Ernst et al., 2009b; Haeussler and Paterson, 1993; Saleeby et al., 1989)。现在出现一个问题,即高度变形的马里波萨(Mariposa)组中的细粒变沉积岩过去被认为是与火山岩相互穿插的(Bogen, 1985; Snow and Ernst, 2008),然而从该 变沉积岩中获得的碎屑锆石年龄数据显示其年龄分布范围为 155~152 Ma,这就意味着侵入杂岩与马 里波萨组是同时代形成的(Ernst et al., 2009b);事实上,最下部的砂岩单元,同时也是剖面互层火 山岩单元的上部,其碎屑锆石给出最年轻的年龄为 160 Ma,这就表明存在两个叠加盆地或者盆地内可能发生过一次重要的间断事件。

#### 5.7 克拉马斯山脉

克拉马斯山脉(Klamath Mt.)(图 2.5,图 4.1 和图 5.9)闻名遐迩,并在加州西北部和俄勒冈 西南部构成一个孤立地块。它由一个以逆冲断层为界的叠瓦状堆叠的地体构成,而这些逆冲断层通常 被认为是向东倾的(Irwin,1981)。地体的概念就是在这里被Irwin(1972)提出,该学者最近将克拉 马斯山脉划分为东、中、西三个地体,每一个地体自身又划分为若干次级地体,并通常将其与内华达 山脉西部的岩石相对应。

克拉马斯东部地体由怀里卡 (Yreka)、崔尼蒂 (Trinity) 和雷丁 (Redding) 三个次级地体构成 (图 5.9),包含克拉马斯地体中最古老的岩石。崔尼蒂次级地体由蛇纹石化橄榄岩组成,并被新元古 界和古生代斜长花岗岩和辉长岩切割,上覆岩石为镁铁质火山岩;怀里卡次级地体由一个早古生代的 变沉积岩推覆体、角闪岩、超基性岩、混杂岩和新元古代的羚羊山 (Antelope) 石英岩等叠瓦状堆叠 组成;雷丁次级地体包含一个早古生代的岛弧地体、密西西比纪变沉积岩、一个包含麦克劳德 (Mc-Clound) 动物群的二叠纪地体,以及一个三叠纪——侏罗纪的岛弧序列 (Potter et al., 1977; Hotz, 1977; Boudier et al., 1989; Peacock and Norris, 1989; Irwin, 2003)。这三个次级地体的岩石大致分别 与内华达地体北部的嘘飞 (Shoo Fly)、泰勒斯维尔和羽毛河带相对应 (Irwin and Wooden, 2001; Irwin, 2003)。

变质地体中部的岩石(图 5.9),琼斯堡(Fort Jones)次级地体的岩石部分是从东部地体分离出来的,包含一个 220 Ma 的绿片岩增生楔,但其主要岩性为角闪岩,并在构造上被片岩和大理岩覆盖(Hacker and Peacock, 1990)。中带与崔尼蒂(Trinity)次级地体的边界是一个向东倾的深大断裂,即"崔尼蒂逆冲断层",该逆冲断层被认为向东深入至崔尼蒂次级地体之下达 100km 或更多(Zucca et al., 1986)。古生代到早侏罗世的北福克(North Fork)次级地体(图 5.9)由变质蛇绿岩、镁铁质火山岩、放射虫燧石岩和灰岩组成, Irwin (2003)将该次级地体与内华达地体北部的水手峡谷,汝拉山和水壶石序列相对应。

在北福克(North Fork)次级地体的西侧坐落着海福克(Hayfork)地体(图 5.9),这一地体由 下列部分组成:①二叠纪—三叠纪破碎地层的一个东侧分区,以及火山岩和沉积岩的混杂岩,该混杂


图 5.9 克拉马斯山地区地质简图 (据 Irwin (2003); Snoke and Barnes (2006); Allen and Barnes (2006)修改) 显示克拉马斯山 (Klamath Mt.) 地区的地体、基岩和各种深成岩套 岩中包括燧石岩和一些角闪岩、灰岩、片岩和蛇纹石化超基性岩的岩块;②一个位于西侧的中侏罗世 玄武质到安山质的火山岛弧序列,该套岩石被认为增生到了海福克东部地体之上(Wright and Fahan, 1988; Irwin, 2010)。西侧岛弧地体的火山和沉积序列,其定年结果显示为177~168 Ma,并被诸多高 K<sub>2</sub>O 的侵入杂岩侵位,这些侵入杂岩的岩性为橄榄-单斜超基性岩、二辉石辉长岩、闪长岩、和二长 闪长岩,以及年龄在170 Ma 上下的中性到酸性含角闪石岩石,表明该岛弧在169 至164 Ma 之间增生 到了更东部的地体之上(Wright and Fahan, 1988; Barnes et al., 2006a)。在碰撞之前,海福克岛弧: ①位于响尾蛇溪(Rattlesnake Creek)地体之上,即一个三叠纪的俯冲杂岩;②从底部的大洋玄武岩 向上渐变为岛弧玄武岩;③包含年龄范围在207 至193 Ma 的辉长质到石英闪长质小侵入体(Wright and Wyld, 1994)。

再往西,另外一个岛弧杂岩,即为罗格-切托克(Rogue-Chetco)岛弧(Garcia,1979,1982), 是被约瑟芬(Josephine)蛇绿岩板片和它的盆地盖层沉积序列、加利斯(Galice)复理石,从海福 克-响尾蛇(Hayfork-Rattlesnake)杂岩中分离出来的。岛弧岩石被运移到史密斯河(Smith River)次 级地体中的加利斯(Galice)复理石内,其中史密斯河次级地体位于奥尔良(Orleans)逆冲断层之 上,而发生运移的时间介于150~153 Ma之间,因为加利斯岩石中碎屑锆石具有153 Ma的最年轻年 龄,同时一个150 Ma的侵入体切割了这一逆冲断层(Miller et al., 2003)。加利斯(Galice)复理石 沉积于164~162 Ma的约瑟芬(Josephine)蛇绿岩之上,而该蛇绿岩反过来又在罗格谷(Rogue Valley)次级地体之上,这个次级地体是一个同时代的157 Ma岛弧,部分位于地体的西侧(Harper et al., 1994)。

几个重要的侏罗纪—白垩纪深成岩套产出于克拉马斯中:①一个称为伍利溪(Wooley Creek)岩套,年龄为167~152 Ma,成分复杂,由超基性岩、辉长岩、石英闪长岩、英云闪长岩、奥长花岗岩和花岗闪长岩等组成;②年龄为151~147 Ma的西部岩套,由超基性岩到花岗闪长岩构成,该岩套侵入该地区的时间正好在155~150 Ma变形减弱阶段和在此阶段之后;③一个年龄为大约140 Ma的岩套,岩石组合为英云闪长岩、奥长花岗岩和花岗闪长岩,通常认为反映了岩浆混合的证据,并有古老地壳物质的注入(Allen and Barnes, 2006; Barnes et al., 2006b)。

## 5.8 侏罗纪岩浆岩岩石

一条以侏罗纪火山岩、沉积岩和深成岩岩石为主的变化多样的带,通常认为代表了一个相对低位的、面向西的弧,位于推测的、尚未出露的北美克拉通古西部边缘(Hamilton, 1969a; Burchfi el and Davis, 1972, 1975; Busby-Spera, 1988; Busby-Spera et al., 1990; Fisher, 1990),构成一条线性岩带,北起加州北部,一直向南延伸经过内华达州,直至加州的莫哈韦沙漠(Mojave)(图 2.5)。另一条类似走向的南东东带,穿越索诺兰沙漠(Sonoran),然后经过亚利桑那州南部-墨西哥北部,这条带一般认为与前面所述线性带是相互关联的(Tosdal et al., 1989; Haxel et al., 2005)。然而,另外一条走向为北东的带,穿越了内华达州西北部一直到蛇河(Snake River)平原(Crafford, 2007, 2008)。在内华达州,莫哈韦-索诺兰沙漠分区与内华达山分区之间走向和地质情况发生突变,表明其间存在断层(图 2.5 和图 2.7)。这些杂岩和其他各式各样的岩石将在下文按地区描述。

## 5.8.1 内华达山区

在内华达山区北部, 侏罗纪的岛弧岩石出露在靠近汝拉山(Mount Jura)的逆冲岩席内,这些岩石以三叠纪末到侏罗纪的塔特尔溪-汝拉山沉积序列以及水手峡谷组为典型代表,它们由厚的、大倾角倾斜甚至倒转的陆源沉积岩组成,其源区主要来自火山岩,包括一系列安山质-英安质-流纹质的熔岩流、近源角砾和凝灰岩,这些沉积岩的整体厚度达到了 8~11 km (Christe and Hannah, 1990; Harwood, 1992, 1993; Stewart et al., 1997; Lewis and Girty, 2001; Templeton and Hanson, 2003)。水壶石沉积序列中包含 180 Ma 的高钾火山岩,同时被 178 Ma 的斑岩型 Cu-Au 矿床切割,随后被 161~148 Ma 的汝拉山沉积序列不整合覆盖 (Christe, 1993, 2011; Dilles and Stephens, 2011)。Emigrant Gap 复

式岩体和 Haypress Creek 侵入体切割了早先发生变形的水手峡谷组和塔特尔溪组的沉积岩,同时还切割了嘘飞杂岩,其中嘘飞杂岩中各个期次的岩性年龄变化范围为 168 至 163 Ma (Girty et al., 1993a, 1995; John et al., 1994; Hanson et al., 1996, 2000)。但是,在这个剖面上的薄凝灰岩单元——在向东的泰勒斯维尔 (Taylorsville) 逆冲断层下面发生局部倒转——最近获得的年龄为 128 ± 3 Ma (Christe, 2010, 2011)。

变质的玄武质-安山质角砾岩、砾岩和浅成侵入体,很可能与塔特尔溪组相关,同样产出于太浩湖(Lake Tahoe)以北的威尔第(Verdi)山脉东南部(Pauly and Brooks, 2002)。太浩湖西南部,具有类似年龄的岩石出露于 Mount Tallac 顶垂体中(Saucedo, 2005),岩石组成包含火山碎屑砂岩、砾岩和单一岩性的火山角砾岩,被 164 ± 7 Ma 的金字塔峰(Pyramid Peak)花岗岩和体积较小的中性深成岩体侵位(Sabine, 1992; Fisher, 1990)。太浩湖的南东侧,即位于沃克河(Walker River)集水区,Schweickert(1976)描述了若干侏罗纪的浅成岩体和相关火山岩及外生碎屑岩,所有的这些岩石都发生了褶皱。在松子(Pine Nut)山脉,刚超过内华达和松子断层南西侧的边界,分布着约 2 km 的火山碎屑岩,同时含少量安山质熔岩,并存在三叠纪的碳酸盐岩夹层,这套岩石被厚大于 3 km、年龄范围从 170 至 162 Ma 的沉积序列覆盖,该序列的岩性主要为粉砂岩、火山砾岩、安山质熔岩和凝灰岩(凝灰流),虽然未定年,但很可能是被侏罗纪和白垩纪的深成岩体切割(Wyld and Wright, 1993)。

靠近内华达的耶灵顿 (Yerington), 168.5 Ma 的耶灵顿和 166 Ma 的三叶草 (Shamrock) 岩基侵入 到三叠纪安山质到流纹质熔岩的厚层沉积序列中, 该熔岩流被 233 Ma 的侵入体切割, 同时被 1800 m 厚的晚三叠世—早侏罗世沉积序列覆盖, 该沉积序列的主要岩性为非火山相的砂岩和灰岩, 同时这一 沉积序列又被一次短暂存在的爆发式的火山岩覆盖, 爆发时间为 169~165 Ma, 主要岩性为安山岩、 英安岩和玄武质熔岩流, 同时还存在少量相关的沉积岩和火山碎屑岩 (Dilles and Wright, 1988; John et al., 1994; Proffett and Dilles, 1984, 2008)。耶灵顿岩基赋存有丰富的斑岩型铜矿 (Dilles, 1987)。

在内华达山脉内部,许多顶垂体的岩石可以划分为多个序列:①在内华达山脉东端莫诺湖(Mono Lake)附近及南侧的莫里森山(Mount Morrison)地块,主体向西陡倾,为发生变质和多期变形的古生代岩石;②Koip沉积序列:三叠纪—侏罗纪的变火山岩和变沉积岩,该序列不整合于古生代岩石之上,并呈现向西逐渐变年轻的趋势;③白垩纪变火山岩和沉积岩,不整合于Koip和金斯(Kings)沉积序列之上或局部与上述两个沉积序列呈断层接触关系;④金斯沉积序列:主要为面向东的火山岩和沉积岩,该序列产出于部分金斯-卡维亚(Kaweah)岩基和Yokohl Valley顶垂体的西部;⑤Kem 高原顶垂体,包含了未知源区的变沉积岩,但是一些学者认为与罗伯茨山(Roberts)外来体有关(Dunne and Suczek, 1991; Chapman et al., 2012);⑥强变形的、以变质碎屑岩为主的伊莎贝拉(Isabella)顶垂体(Bateman, 1992; Saleeby et al., 1978; Saleeby et al., 1990; Saleeby and Busby-Spera, 1986, 1993; Stevens and Greene, 2000)。顶垂体内岩石的整体走向相对于内华达山脉稍微偏东。

有一个雪湖(Snow Lake)顶垂体(图 5.7)明显与其他的不同(Wahrhaftig, 2000),它包含了 148 Ma的辉长质岩墙和前寒武纪晚期到寒武纪的变沉积岩,其中 148 Ma的岩墙与 150~148 Ma的独立岩墙群相关,而对于后者,根据碎屑锆石剖面以及存在虫形石(Scolithus)判断与死亡谷--莫哈韦沙漠地区南部的岩石类似(Smith, 1962; Stewart, 1970),并认为前寒武纪晚期到寒武纪的变沉积岩是从死亡谷--莫哈韦地区南部迁移过去的(Lahren et al., 1990; Lahren and Schweickert, 1989, 1994; Schweickert and Lahren, 1990; Grasse et al., 2001a; Memeti et al., 2010a)。这些学者提出金斯沉积序列南部的岩石可能是相同层序的一部分,并认为通过岩基内的雪湖--莫哈韦断裂把这些岩石向东分离出去;但是,界定特征并未出现在这些顶垂体中(Saleeby et al., 1978),因此雪湖顶垂体可能代表了一个断层楔形体,它从莫哈韦地区向北迁移,也可能存在两个岩基内的断层或者一个断层系统的斜面(Kistler, 1993)。

在鞍袋湖(Saddlebag Lake)顶垂体(图 5.7)东侧附近几个更小的顶垂体东侧,是一个向西倾 30

的、厚达数千米的沉积岩堆积体,岩石组成为砾岩、流纹质熔岩流凝灰岩,安山质的熔岩、角砾岩, 以及相关的外生碎屑岩,这些岩石不整合覆于古生代和早三叠世岩石之上(Schweickert and Lahren, 1993a),至少有一个火山杂岩——具有流纹质喷出相岩石,获得了一个222±5 Ma的U-Pb 锆石年 龄,以及一个可能成因上相关的深成岩体获得了近232~219 Ma的锆石U-Pb 年龄——产出于鞍袋湖 顶垂体内,该杂岩发生褶皱,并沿着向东的逆冲断层发生迁移,发生运移的时间早于年龄为168Ma 的莫诺穹窿 (Mono Dome)内花岗闪长岩的侵入时间(Schweickert and Lahren, 1993a, 1993b, 1999; Barth et al., 2011)。

沿着鞍袋湖顶垂体的走向往南东方向是里特山脉(Ritter Range)顶垂体,它包含了一个向西倾斜的沉积序列,岩石组成为发生多期变形的古生代沉积岩和一个可能的逆冲堆积体,该逆冲堆积体由晚三叠世到早侏罗世(214~186 Ma)的变火山岩和变沉积岩组成,这套岩石又被一套164 Ma的沉积序列不整合覆盖,该沉积序列的岩石组合为变熔结凝灰岩、熔岩流和角砾岩,以小型断层为边界的年龄为140~130 Ma的火山岩片——由断层从其他岩石中分离出来的——一个向西倾斜的中白垩世(近100 Ma)火山填充沉积序列,上述沉积层序被一个98 Ma的、被认为是复活(resurgent)的岩体侵入(Rinehart and Ross, 1964; Huber and Rinehart, 1965; Russell and Nokleberg, 1977; Fiske and Tobisch, 1978, 1994; Greene et al., 1997; Tobisch et al., 1986, 2000)。一套年龄为226~218 Ma 被称为"白钨矿侵入体岩套"的深成岩石,侵入到更老的火山岩中(Bateman, 1965a, 1992; Barth et al., 2011)。

再往南东方向,狭窄的戈达德山顶垂体(图 5.7)包含了变质和变形的、向南西倾的火山碎屑 岩,并夹有交替稀疏出现的凝灰岩、凝灰角砾岩、熔岩流凝灰岩(143 ± 3 Ma 和 131 ± 6 Ma)、熔岩 流(156 ± 2 Ma 和 140 ± 1 Ma)、若干侵入体(159~156 Ma)(其中的一些比它们的围岩要老),以及 少量的碳酸盐岩(Bateman, 1965b; Bateman and Moore, 1965; Lockwood and Lydon, 1975; Moore, 1978; Tobisch et al., 1986)。上述沉积序列(和更老的侵入体)在剪切变形时期的区域变质作用过程中 可达角闪岩相,展示出老地层覆盖新地层和新地层覆盖老地层的构造接触关系,具体的变形时间被限制 在最年轻的凝灰岩年龄131 ± 6 Ma 和发生面理化的吉文斯山(Mount Givens)花岗闪长岩年龄90 ± 3.5 Ma 之间,而花岗闪长岩侵位到了戈达德顶垂体的南西部(Tobisch et al., 1986; Bateman, 1992)。

沿着走向往南东延伸,一系列侏罗纪岩床状岩体将戈达德山顶垂体和向南西倾斜的橡树溪(Oak Creek,科罗拉多州)顶垂体相连接(图5.7),其中橡树溪顶垂体的岩石组合包含一个165 Ma 的变 质变形熔岩流、流纹质和英安质熔岩流凝灰岩和时间上相关的侵入体,这些岩石被年龄为近109±2 Ma 的一套火山沉积序列覆盖,而火山沉积序列的岩性为中基性成分的熔岩、凝灰岩和火山碎屑岩(Moore, 1963; Saleeby et al., 1990)。靠近橡树溪顶垂体的一个糜棱化正片麻岩的年龄为164±4 Ma,变形的侏罗纪深成岩体的年龄为165~164 Ma (Mahan et al., 2003)。这些相对老的变形深成岩体被独立的岩墙群切割,这是首次被描述的典型地区(Moore and Hopson, 1961; Moore, 1963)。向西进入红杉(Sequoia)国家公园的 Giant Forest 地区(图5.7),一些顶垂体包含有侏罗纪的表壳岩系和深成岩体,其中一些岩石的年龄可达162 Ma,同时许多这样的顶垂体中包含一些岩墙,而这些岩墙可能是独立岩墙群的一部分(Sisson and Moore, 1994; Moore and Sisson, 1987)。在靠近三分峰(Triple Divide Peak)的地区(Moore, 1981; Moore and Sisson, 1987),岩墙可能沿着它们的围岩发生了褶皱,因为那些褶皱明显比典型的褶皱更为开阔,其中在翼部走向北西的地方,对应岩墙的走向也为水西;而在翼部走向南西的地方,对应岩墙的走向也为南西。这些几何学关系表明:至少在局部,这些岩墙可能原来就是岩床,只是褶皱作用让它们呈现为"岩墙"。

在橡树溪顶垂体以西约 30~40 km 处坐落着博伊登洞穴(Boyden Cave)顶垂体(图 5.7),有一个倾斜角度很大、厚度约为 5 km 的古生代变沉积岩,岩性包括石英岩、大理岩、泥质岩和砂岩,这 套岩石是从中白垩世沉积序列中分离出来的,而中白垩世沉积序列由岩性为大倾角的、年龄为 105~ 100 Ma 的流纹质熔岩流凝灰岩和硅质-中性火山岩以及火山碎屑岩的混合物共同组成,它们后来被年 轻的侵入体侵位(Saleeby et al., 1990)。位于 Giant Forest 地区的博伊登洞穴顶垂体的南西侧,存在 若干小的顶垂体,这些小顶垂体包含侏罗纪—白垩纪的岩石。定年结果为 162 Ma 的 Yuncca 山脉的花 岗闪长岩, 侵位到了未定年的沉积序列中, 该沉积序列包含变玄武岩和安山质熔岩, 其中的白垩纪熔 岩流凝灰岩中一小部分不整合覆于较老岩石之上 (Sisson and Moore, 1994)。

往南,顶垂体内的大多数岩石发生了强烈的变质和构造叠加(Wood and Saleeby, 1998),这就使 得恢复原岩变得异常困难,但是在伊莎贝拉湖(Isabella Lake)顶垂体中也出露一些已知的白垩纪岩 石露头(图 5.7),这些岩石不整合覆盖于先前变质的、属于金斯沉积序列的变沉积岩之上,这套不 整合面之上的岩石称之为"厄斯金湖"沉积序列(Saleeby and Busby-Spera, 1986)。厄斯金湖 (Erskin Lake)沉积序列内的岩性包含有陡倾角的熔结凝灰岩单元、火山角砾岩和一个可能的火山颈, 这些火山岩的锆石 U-Pb 年龄分布范围在近107 到102 Ma 之间(Saleeby et al., 2008)。

在孤松(Lone Pine)以西的亚拉巴马山(Alabama Hills)出露的侏罗纪岩石(Stone et al., 2000),两个已经发生褶皱(图 5.7),而主体向南西倾斜,沉积序列露头:一个是下部层序的厚度为约 2 km 的熔岩流凝灰岩和火山碎屑岩,它们被岩墙和岩床侵位;上部层序与下部呈不整合接触关系,上部层序岩性由变形和蚀变的沉积岩组成,同时还伴有厚度大于 450 m、年龄为 170±4 Ma 的流纹质熔岩流凝灰岩(Dunne and Walker, 1993)。

在白印优山脉的多个地区出露有侏罗纪的岩浆岩(图 2.5 和图 5.7)。在北部,最老的是含有大理岩夹层的硅质火山岩,以顶垂体的形式出露于 165 Ma 的巴克罗夫特(Barcroft)深成岩体内,在 3 km 厚的下部为变质的玄武安山质到流纹质熔岩、凝灰岩和浅成侵入体,被深成岩体切割,而上部又被另外一个变火山岩沉积序列覆盖,上覆的这一套火山岩中含有近 154 Ma 的流纹质凝灰岩(Hanson et al., 1987)。他们同时还报道了切割火山沉积序列的浅成侵入体的年龄为 137 Ma。再往北东内华达州的埃斯梅拉达(Esmeralda)县境内,出露若干很可能为侏罗纪的深成岩体,例如西尔韦尼亚(Sylvania)和帕尔梅托(Palmetto),但是其具体年龄以及相关火山岩的年龄尚不清楚(Albers and Stewart, 1972)。

在印优山脉(Inyo Mt.) 南部存在着三个相互分离以不整合为界的沉积序列,它们的岩石组成主要为向南西倾的变沉积岩和变火山岩,不整合覆于早三叠世岩石之上:①一个厚 500 m 的沉积序列, 岩性为砾岩、角砾岩、砂岩和少量的玄武质熔岩;②一个 169~168 Ma 的沉积序列,厚度可达 800 m, 主要岩性为熔结凝灰岩岩片、安山质和流纹质熔岩,夹有少量卵状火山砾岩;③2.5~3 km 厚的火山 成因砾岩、砂岩和细粒沉积岩,夹有少量的熔岩流凝灰岩、熔岩流和碎屑流,其中的一个英安质凝灰 岩的定年结果为 148 Ma (Dunne and Walker, 1993; Dunne et al., 1978, 1998)。

在南部的阿格斯(Argus)、斯莱特(Slate)及周边区域(图 5.7),出露缺少年龄数据的侏罗纪岩石,这些岩石包括从安山质过渡到玄武质的熔岩与火山碎屑岩互层,硅质火山灰流和空落凝灰岩(air-fall tuffs),一部分岩石被深成杂岩体切割,如 186 Ma 的 Bendire 深成岩体,174 Ma 的亨特山脉岩基(Dunne,1979)及161 Ma 的 Maturango 深成岩体(Dunne et al., 1978)。另一个沉积层序称为温泉(Warm Spring)组,岩石组合为砾岩、砂岩、尘雨凝灰岩和安山斑岩,出露在帕纳明特(Panamint)山脉的比尤特(Butte)峡谷地区,被K-Ar 年龄为145~137 Ma 的深成岩体切割(Abbott, 1972),

厄尔巴索(El Paso)山脉位于加洛克(Garlock)断裂的北侧,存在一个石英闪长质到石英二长质的深成岩体,其中的角闪石 K-Ar 年龄为152 Ma,黑云母 K-Ar 年龄为146 Ma,该深成岩侵入到先前发生过变质和变形的岩石当中(Carr et al., 1997)。Carr 等人通过填图识别出两个构造并置体,以及独特的早古生代层序:一个为西侧的细粒变黏土质和黏土质-变燧石单元,另一个为东侧的地台相,岩石组成为变泥岩、大理岩、沉积石英岩和笔石板岩,它们全部被密西西比纪的砾岩和变泥岩不整合覆盖,这一不整合事件被认为与安特勒造山作用导致的前渊盆地填充有关(Carr et al., 1980),而上覆的这些变沉积岩又先后被晚二叠世的片麻质深成岩体和一个无面理化的早三叠世深成岩岩套侵入,该岩套成分范围从辉长岩到花岗岩。

# 5.8.2 内华达西北部

在松子(Pine Nut)断层以北的内华达的北西部(图 4.1 和图 5.10),有一个北东延伸的杂岩带,由三叠纪到白垩纪的沉积岩、变质岩和岩浆岩构成(Crafford, 2007, 2008; Oldow, 1984; Oldow et al.,

1989)。在那里,运动学方向是沿着向南东汇聚的 Luning-Fencemaker 褶皱-逆冲带进行的——一些学者认为该褶皱-逆冲带的形成发展史是与逐渐向东迁移的弧后缩短过程相协调一致的,这个过程中产生了塞维尔远东褶皱-逆冲带(Burchfi el et al., 1992; Saleeby and Busby-Spera, 1992; DeCelles and Coogan, 2006) ——并置的黑石岛弧地体与洪堡特(南达科他州)岩浆杂岩位于拥有几千米厚且发生了强烈变形的三叠纪盆地相岩石之上(Burke and Silberling, 1973; Speed. 1978),基底岩石的性质目前尚未知,盆地相岩石逆冲到了发生微弱变形且面向西的三叠纪碳酸盐岩台地之上,该碳酸盐岩台地则是沉积在过去发生过变形的宝山和罗伯茨山脉外来体之上(Oldow, 1984; Elison and Speed, 1988, 1989; Quinn et al., 1997; Dilek et al., 1988; Dilek and Moores, 1995; Wyld, 2000, 2002)。盆地相岩石 的碎屑锆石年龄谱显示其变化范围在 1145 到 948 Ma 之间(Manuszak et al., 2000),这就表明盆地相



图 5.10 美国内华达西部岛弧、地台和盆地岩层之间关系简图 (转引自 Crafford (2007, 2008); Oldow (1984))

洪堡特(Humboldt)杂岩的位置,大致出露在出现几个逆冲岩片覆盖的宽阔区域(Dilek and Moores, 1995)

岩石并不是像 Speed (1988) 认为的那样起源于北美。岛弧地体在不晚于早侏罗世与盆地相岩石发生接触,但是盆地相岩石与东部大陆架的沉积关系尚不清楚。

黑石地体包含厚达 10 km 的晚古生代沉积和火山岩地层,其上又被中三叠世到早侏罗世的中基性火山成因岩石以及各式各样的、年龄范围在 170 至 160 Ma 之间的几个侏罗纪侵入岩体覆盖(Quinn et al., 1997; Wyld, 2000)。大约近 165 Ma 的洪堡特岩浆杂岩由一个同源的从超基性到花岗质的深成岩套组成,该杂岩被一套含火山岩沉积序列不整合覆盖,该火山岩沉积序列的岩石组成为枕状玄武岩、安山质-英安质熔岩流和一系列火山岩角砾岩和凝灰岩(Dilek and Moores, 1995)。黑石地体的变形开始时间最早可追溯到侏罗纪早期,但是主要的变形幕与其在崩塌盆地之上的就位时间相一致,即在163 Ma 稍后的时间,随后发生了 7~14 km 的剥露,这就使得变形时间和北美克拉通晚侏罗世的沉积时间之间发生联系的可能性极为微弱,因为那个时间既没有发生变形,也没有发生造山沉积(Wyld et al., 2003; Wyld and Wright, 2009)。早白垩世冲积碎屑沉积物,夹少量定年结果为 125~124 Ma 凝灰流,不整合覆盖在隆升剥蚀的侏罗纪岩石之上,并被一个 123±1 Ma 的深成岩体切割(Martin et al., 2010)。晚白垩世的深成岩体(93~88.5 Ma)侵位到了更老的岩石中,它们的年龄和成分很大程度上与沿着内华达山脉山脊产出的"内华达山脊岩浆事件岩浆岩"相类似(Smith et al., 1971; Van Buer and Miller, 2010)。

## 5.8.3 莫哈韦沙漠地区

大部分关于莫哈韦沙漠地区地质情况明显是建立在一个古生代基底之上,即那些位于前寒武结晶 基底之上、分布稀疏的露头,但是由于莫哈韦沙漠地区存在很多新生代走滑断层,且该区一半以上被 冲积层或新生代火山岩或者沙漠覆盖(未出露岩层),所以重建该区中生代甚至更老时代的古地理分 布将是一项极富挑战性的工作(Burchfi el and Davis, 1981)。

在莫哈韦沙漠地区内存在两个不同的地体,一个以深水沉积岩、晚二叠世变形及 260 Ma 的深成 岩浆活动/安山质火山活动为典型特征,而所有的变形都是通过向西的褶皱和逆冲断层实现的;另外 一个地体则由浅水相的沉积岩组成,并伴随有密西西比纪之后的变形,与 246~243 Ma 侵位的构造期 后深成岩体并置(Martin and Walker, 1995; Miller et al., 1995; Carr et al., 1997; Walker et al., 2002)。在维克多维尔(Victorville,加利福尼亚州)地区附近(图 5.11),锦绣谷(Fairview Valley) 组不整合于年龄为 243~241 Ma 的深成岩体之上(Miller et al., 1995; Barth et al., 1997)。而锦绣谷 组是一个等倾褶皱的沉积序列,定年工作非常薄弱,可能是侏罗纪的沉积岩石(Walker, 1987; Schermer et al., 2002)。

在蒂哈查皮(Tehachapi)山内的羚羊(Antelope)峡谷北部,加洛克断裂以南,有一个由大理 岩、片岩、变火山岩和超基性岩组成的断续带,统称为豆峡谷(Bean Canyon)组,该组大部分被推 测为白垩纪的深成岩体包围(Dibblee, 1967; Ross, 1989)。豆峡谷组的年龄尚未得到很好地约束,因 为没有从中发现化石,但是根据大理岩和玄武岩单元,Wood和 Saleeby(1998)推测或许与位于伊莎 贝拉湖顶垂体(Pendant)内的晚三叠世—早侏罗世金斯沉积序列相关联;但是一个英安质的变质凝 灰岩最近被 Chapman 等(2012)定年,明确地获得了 273 Ma 的中二叠世年龄。

哈韦沙漠中部地区大面积出露中侏罗世火山岩,被称为响尾蛇(Sidewinder)火山岩系列 (Bowen, 1954),该系列可能在时间上与锦绣谷组相关,可以划分为上下两部分,下部年龄范围为 179~164 Ma,岩石组成为流纹质-英安质火山洼地内(intracauldron)熔灰流,其上被一个流纹质-玄 武质熔岩层序的上部序列不整合覆盖,该上部序列又被年龄为152 ± 6 Ma的流纹质岩墙侵入 (Schermer and Busby, 1994; Schermer et al., 2002)。他们同时对一个厚层熔结凝灰岩进行了U-Pb定 年,该熔结凝灰岩是从下部沉积序列中经过剥露、沉积和可能的断裂过程之后分离出来的,定年结果 为151±1 Ma,该年龄指示了这一凝灰岩比下部层序年轻得多,它更为可能与上部沉积序列产生更为 密切的联系。最近,Fohey-Breting et al. (2010)在响尾蛇火山岩中,通过离子探针来厘定熔灰流单 元中的其中三个样品,进一步证实了早前的结论。他试图将凝灰岩与某个或某些具体侵入岩体建立联

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系:但是最老的 180±3 Ma 的凝灰岩没有找到对等的侵入岩体;一个 161 Ma 的凝灰岩大致与 167~161 Ma 的金条(Bullion)山脉侵入岩套及相关深成岩体存在一定的成因联系,而后者是基特峰-特里戈(Kitt Peak-Trigo)山脉超单元的一部分(Tosdal et al., 1989);一个 150±2 Ma 单元的凝灰岩或多或少与同时期的 155~151 Ma 的钙碱性深成岩体和 149 Ma 黑花岗岩体存在一定的关系。

年龄大致相同的岩石也出露在罗德曼(Rodman)、奥德(Ord)和弗莱山(Fry Mt.)(图 5.11), 其中厚层序列的变质变形灰流凝灰岩、安山岩流,以及各种表层碎屑岩,被年龄为171~166 Ma的中 侏罗世深成岩体切割,而后又被一个侏罗纪末的未变形独立岩墙切割,其中一些可能已经加入了熔岩 流(Karish et al., 1987; James, 1989)。

在几个小范围内,如位于贝克(Baker)南部的 Cowhole 山脉(图 5.11),在以断层为界的古凹 陷中,170 Ma 矿化-流纹岩灰岩流体凝灰岩、火山碎屑岩和安山岩夹有风成砂岩,并被 Cowhole 火山 岩覆盖,该火山岩由约 500 m 的熔结灰流凝灰岩、各种各样的火山岩角砾岩、英安岩,以及位于顶部 附近的流纹质角砾岩组成,又被年龄为 169±2 Ma 的半深成岩床侵入。上述所有的岩石都可能被独立 岩墙切割(Marzolf and Cole, 1987; Wadsworth et al., 1995; Busby et al., 2002)。



#### 图 5.11 彩色地貌图

显示了文中讨论的加利福尼亚州南部和莫哈韦沙漠(Mojave Desert)山脉的位置。底图根据约翰斯·霍普金斯大学应用物理 实验室(Johns Hopkins Applied Physics Laboratory)数据绘制,数据来源是使用1角秒航天飞机雷达地形测量的数字高程 (2000年2月航天飞机进行了为期11天的STS-99雷达地形测绘任务(SRTM),STS-99是航天飞机历史上第九十六次航天 飞机任务,也是奋进号航天飞机的第十四次太空飞行,译者注)

位于莫哈韦地块中东部的一系列山脉,包括 Clipper、Ship、派尤特(Piute)和老妇(Old Woman)山脉,还有 Kibeck山,它们位于靠近侏罗纪岩浆作用的北东末端,其中包含了各种各样的 深成岩体和变形特征(Howard et al., 1997; Howard, 2002)。Howard et al. (1995)提出 161±10 Ma 的以二长闪长岩为主的戈德哈默(Goldhammer)深成岩体是在一次逆冲事件过程中侵位的,即元古 宙的结晶基底逆冲到了古生代的地层之上。构造期后深成岩体包括 150~145 Ma 的 Ship 山脉深成岩体,一个混合杂岩体,由花岗岩、石英二长岩、辉长岩、闪长岩和二长闪长岩组成,以及一个近 145

Ma (Gerber et al., 1995) 的岩墙岩套。

在 Palen 山脉(图 5.11),174±8~162±3 Ma 的圆顶石(Dome Rock)沉积序列覆盖在 Palen 组的 砾岩和砂岩之上,4 km 厚的圆顶石沉积序列由英安质到流纹质凝灰岩、熔岩流、浅成侵入体,以及 一个穹窿杂岩和各式各样的外生碎屑岩组成,而 Palen 组又被以白垩纪为主的麦考伊山(McCoy)组 覆盖(Fackler-Adams et al.,1997; Busby et al.,2002)。再向东,在惠普尔山(Whipple)变质核杂 岩的下部出露了一套变形的花岗闪长质-石英闪长质深成岩体,即为 Whipple Wash 岩套,定年结果为 89±3 Ma (Anderson and Cullers, 1990)。

越过亚利桑那州南东边界的圆顶石山脉 (图 5.11),出露有古生代变沉积岩和中生代火山岩剖面,被中生代的深成岩体侵入。一个 165±3 Ma 流纹质凝灰岩和一个 164 Ma 的花岗闪长岩的喷发或结晶时间早于向南西发生侧卧褶皱作用的时间,而一个 161 Ma 的淡色花岗岩则晚于这一变形 (褶皱作用)时间 (Boettcher et al., 2002)。

在中新世中莫哈韦变质核杂岩的下盘出露了一个 100 km×20 km 的强烈变质变形带,发生变质变形的时间为侏罗纪和白垩纪(图 5.11)(Fletcher et al., 1995)。下盘的斑状变火山岩,即 Hodge 火山沉积序列,获得 U-Pb 锆石年龄为近 170~164 Ma,而一个构造期后花岗岩具有 151±11 Ma 的年龄,同时,一个切割白垩纪变形组构的白云母-石榴子石花岗岩给出了 83±1 Ma 的年龄(Boettcher and Walker, 1993)。在 Shadow山(图 5.11),Martin等(2002)报道了一个新元古界—中生代的沉积岩石序列,与死亡谷中的沉积序列类似,即在发生平卧褶皱作用期间的变形和变质时间早于 148±1.5 Ma 的辉长岩和闪长岩侵位时间,而晚于 144~143 Ma 花岗岩结晶时间。

在巴斯托(加利福尼亚州)的东南部产出大量 167 Ma 的侵入杂岩,称为金条山(Bullion)侵入 岩套,岩石组成主要为花岗岩、石英二长岩和石英二长闪长岩;它们出露于金条、平托(Pinto)和 老鹰(Eagle)山脉(图 5.11);这些岩石可能与同时期的戴尔湖(Dale Lake)火山岩相关,而戴尔 湖火山岩以中性熔岩、外生碎屑岩和凝灰岩为主,其中凝灰岩中包含独特的卵形、淡紫色碱性长石斑 晶,这些长石斑晶与侵入杂岩中石英二长岩中的长石非常类似(Mayo et al., 1998; Howard, 2002)。

在巴斯托的东北部的 Tiefort 山(图 5.11), Schermer 等(2001) 描述了一套未知年龄的沉积岩和 变火山岩,推测其年龄为中生代,这套岩石被 164~160 Ma 的片理化深成岩体侵入,但是走向朝北的 长英质和镁铁质岩墙(148 ± 14 Ma)切割了较老深成岩体中的片理;而在 Tiefort 山南部,伟晶岩年 龄为 82 Ma,正片麻岩 U-Pb 年龄为 105 Ma,表明该区存在侏罗纪和白垩纪两期变形事件。

在 Cronese 山的东侧(图 5.11),局部发生强剪切化和逆冲作用的绿片岩相变火山岩,原岩为凝灰岩和熔岩流、变质深成岩(166 ± 3 Ma)和变质沉积岩,这些岩石发生倒转并被一个年龄为155±1 Ma 的构造期后花岗岩切割(Walker et al., 1990b)。在 Alvord 山的西部,Miller and Walker (2002)描述了一个片理化的二长闪长岩-石英二长岩深成岩体,该岩体的定年结果为179 Ma (Miller et al. 1995),随后又被未发生片理化的年龄为149±3 Ma 的辉长质和角闪石辉绿岩墙以及中性组分年龄为83 Ma 的斑状岩脉切割。

#### 5.9 白垩纪岩基岩石

几个科迪勒拉型岩基通常作为长期的、构成北美西部岩浆带的组成部分(图 2.2 和图 5.12)。 Hildebrand (2009)依据几个火山-侵入岩带中的宁静期划分出明显的岩浆作用活动期(图 5.13): ①晚侏罗世—早白垩世伸展弧阶段,保存完好的火山岩,含有河流相和浅海相陆源碎屑岩夹层;② 120~80 Ma的岩浆活动阶段(图 5.14),科迪勒拉岩基的主体形成;③75~60 Ma 俯冲板块破裂岩浆 作用阶段,在某些地方岩浆岩带只有几十千米宽,却有上千千米长,由中性到酸性的侵入岩和火山岩 岩石组成;④岩浆弧阶段,包含典型的大约在 53 Ma 开始喷发的弧岩石,由板块向东俯冲引起,并延 续至今构成太平洋北西部。但非常重要的是,两个岩浆阶段之间通常有 5~10 Ma 的岩浆宁静间隔期。 所有的岩基都有着许多相同点,但是最为重要的一点是:在科迪勒拉基岩阶段,大部分岩基处于同一 时代,即约 120~80 Ma (图 5.14)。









海岸深成杂岩体据 Gehrels et al. (2009)修改,内达华山脉岩基据 Paterson et al. (2012)修改 显示不列颠哥伦比亚省海岸深成杂岩体的平均岩浆流(基于深成岩)-年龄与加利福尼亚州内华达山脉岩深成杂岩体的平均深 成岩增加速率的比较。海岸深成杂岩体图阐明了 Hildebrand (2009)解释的四个岩浆活动阶段。注意在这两个区域之间的流量 总的来说相似。主要的科迪勒拉岩基相(120~80 Ma)因塞维尔和拉勒米碰撞是有短暂边界的。大多数人认为科迪勒拉型岩浆 作用是由单一俯冲带引发的岛弧岩浆作用,但在这里岩浆作用很可能可以分为不同成因的两部分。海岸带和内达华地区之间一 个最重要的区别就是内达华岩基没有后拉勒米岩浆作用

尽管半岛山脉岩基和海岸深成杂岩体非常出名,但是内华达山脉岩基可能是最佳的中生代岩基研究对象。内华达山脉岩基的大部分为白垩系,但是其中常包含更老的中生代岩石(Bateman and Wahrhaftig, 1966; Bateman et al., 1963; Bateman, 1992; Saleeby et al., 2008)。本文作者把白垩纪岩浆作用 从较老的和较年轻的岩套中区分出来,不只是因为较老的岩石能指示弧的极性和形成较年轻岩石的初始俯冲时间,而是因为其间存在明显的岩浆间隔期,有时岩浆间隔期与变形期是相关的,是区分不同 构造背景的重要标志。因为白垩纪基岩的走向相对于较老的岩石稍偏北,按这里的定义严格意义上来 说这些深成岩体不属于岩基部分,前陆变质带中总体较老的岩石主要分布在岩基主体的西北部,在白 印优山脉分布在岩基主体的东南部:在前面部分只讨论了出现的主要岩基中的几个残余岩块。

我也对 120~80 Ma 的科迪勒拉岩基进行了细分,并详细考察它们的起源,这个时期的岩浆作用 发生在海岸深成杂岩体、半岛山岭岩基和内华达山脉岩基,明显由 2 个阶段或 2 个部分组成 (Gromet



图 5.14 科迪勒拉岩基距离-年龄图

(资料来源:内华达: Bateman (1992);半岛山岭: Ortega-Rivera (2003), Silver and Chappell (1988);盐碱地块:
Mattinson (1990);海岸深成杂岩: Gehrels et al. (2009))

四个集中爆发的科迪勒拉型岩基的距离-年龄图,表明主要年龄为120~80 Ma。密切注意年龄小于 80 Ma 的海岸深成杂岩 被认为是俯冲板片断离岩浆作用产物 (Hildebrand, 2009),同时很好地展示了 2 个前 105 Ma 的海岸深成杂岩套,它们是 在约 100 Ma 大部分科迪勒拉型岩基特征的转换挤压变形期间加入的

and Silver, 1987; Silver and Chappell, 1988; Kistler, 1990, 1993; Bateman et al., 1991; Todd et al., 2003; Lee et al., 2007; Lackey et al., 2008), 即在105~100 Ma 之间的变形和碰撞(Kimbrough et al., 2001; Tulloch and Kimbrough, 2003; Saleeby et al., 2008; Gehrels et al., 2009), 广而言之, 较老的变形前的岩浆岩岩石出现在西部, 而较年轻的碰撞后岩浆作用岩石在东部占优势。

# 5.9.1 内华达山脉岩基

内华达岩基内的深成岩石成分范围由辉长岩到淡色花岗岩,但是最为常见的岩石类型为石英闪长 岩、花岗闪长岩和花岗岩 (Bateman and Wahrhaftig, 1966; Bateman et al., 1963; Bateman, 1992; Ross, 1989)。一般而言,岩基内数以百计的深成岩体之间的关系彼此是截然的,或者被较老的变质 岩小岩片分隔开 (Bateman, 1992)。

岩基底下的结晶基底属性不详,但是根据深成岩体的成分和规模,下部基底一定是大陆性质的 (Hildebrand and Bowring, 1984)。自从 Moore (1959)识别出偏基性的深成岩体位于"偏中性成分" 的西侧后,根据地球化学成分,磁化率、年龄、放射性和稳定同位素的数据,诸多学者确信内华达岩 基可以进一步划分为西部较老的部分和东部较年轻的部分(Chen and Tilton, 1991; Bateman et al., 1991; Kistler, 1990, 1993; Saleeby et al., 2008; Lackey et al., 2008, 2012a, 2012b; Chapman et al., 2012)。目前所知的内华达岩基的围岩均大于 98 Ma,它们都发生了变形(Peck, 1980; Nokleberg and Kistler, 1980; Bateman et al., 1983a; Saleeby and Busby-Spera, 1986; Memeti et al., 2010a; Wood, 1997; Saleeby et al., 1990)。关于整体变形的时间,如果从岩浆岩角度考虑,显然发生在岩体侵位 (98~85 Ma)之前,而发生在 98~85 Ma 的大规模岩浆侵位事件被称为"内华达山脊岩浆事件" (Coleman and Glazner, 1998; Davis et al., 2012);如果从变质岩角度考虑,在内华达岩基最南侧发生 的 95 Ma之前(Saleeby et al., 2007, 2008)。

由于角闪石角岩具有狭窄到中等的变质晕,可以判断大多数岩基深成岩石的侵位层次属中-浅层 次水平(Bateman, 1992)。精细的地质压力计指示了肉眼可见的主要岩基结晶块体的压力为 3~4 kb,

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除了岩基最南端的压力达到6kb以上(P>6kb)以及大部分岩石达到高角闪岩相或麻粒岩相外,沿着内华达岩基东边的岩石压力以1~2.5kb为主(Ague and Brimhall, 1988; Ross, 1989; Wood and Saleeby, 1998; Saleeby et al., 2007; Nadin and Saleeby, 2008)。

尽管在内华达山区小范围有三叠纪和侏罗纪的岩体,但深成岩体的主体年龄仍然为125~82 Ma (Stern et al., 1981; Chen and Moore, 1982; Bateman, 1992; Irwin, 2003)。在岩基的中东部, Bateman (1992)详细研究了几个具同源捕虏体的侵入岩套,但不一定是同源岩浆,因为这些深成岩体在岩相学、成分、结构特征以及空间距离上(接近性)都具有显著不同。目前了解最为深入的是<100 Ma 的 图奥勒米 (Tuolumne)复式环状杂岩体,该杂岩体具有明显的分期性,具有不同的嵌套单元,这些单元由外向内变年轻,且岩石色率也逐渐降低(Calkins, 1930; Bateman and Chappell, 1979; Bateman et al., 1983b; Huber et al., 1989)。具有镶嵌构造的侵入杂岩体沿着内华达山脊东部侵位,侵入年龄介于约 98~86 Ma 之间,其主要特征为:外侧为年龄相对老的英云闪长岩和花岗闪长岩,逐渐向中间截然地变为较年轻的角闪石花岗闪长玢岩,杂岩体的核部位置甚至出现更年轻的含有钾长石巨晶的花岗岩和花岗闪长岩(Bateman, 1992; Coleman and Glazner, 1998; Hirt, 2007)。另外一个具有类似镶嵌构造的 Sahwave 杂岩体产于内华达山区西北部,其区域走向为北北东向(Van Buer and Miller, 2010)。根据 U-Pb 年龄的集中性分组特征,Coleman 和 Glazner (1998)认为图奥勒米(Tuolumne)侵入岩套和内华达山脊南侧和北侧的其他杂岩体类似,例如 Whitney 和 Mono Pass 侵入岩套(Gaschnig et al., 2006; Hirt, 2007),均形成于 10 Ma 的岩浆大爆发期间,并命名为"内华达山脊岩浆事件"。

往南, Nadin 和 Saleeby (2008)将内华达山脉划分为三个纵向带:①西部带,富含镁铁质和石英 闪长质岩石;②中轴带,深成岩体伴生有中白垩世硅质变火山岩和相关浅成侵入体的顶垂体;③东部 带,主要由大体积的复式深成岩体组成,其岩石组成从外向内依次为英云闪长岩、石英闪长岩、花岗 闪长岩和含钾长石斑晶的似斑状花岗岩,这些复式岩体整体上与"内华达山脊岩浆事件"产出的岩 体类似。正如 Kistler (1990)采用的方法一样, Nadin 和 Saleeby (2008)将初始<sup>87</sup>Sr/<sup>86</sup>Sr=0.706 的等 值线作为划分西部带和中轴带边界的标准 (译者注:参考了 Nadin 和 Saleeby (2008)的原文,认为 这样翻译更为忠实原意),而中轴带和东部带的界限为大型复式杂岩体 (译者注:应该为东内华达岩 基晚期复式岩体)的西界。

内华达山南部和特哈查比(Tehachapis,加利福尼亚州)的岩石从约100 Ma的9kb埋藏压力, 抬升到约95 Ma的4kb(Saleeby et al.,2007)。同样在南侧,Salleby等(2008)识别出一个2~5 km 宽的主剪切带(Busby-Spera and Saleeby,1990),即Proto-Kem 峡谷断裂-东特哈查比剪切带,分隔 了各种类型的深成岩体岩套,出露于该剪切带断层的西侧,包括105~98 Ma的熊谷(Bear Valley)岩 套,110~95 Ma的Needles岩套,以及105~102 Ma的科恩河(Kem River)岩套,而规模大的95~84 Ma的Domelands岩套和100~94 Ma的南福克岩套出露于断层的东侧。Nadin and Saleeby(2008)利用 地质压力计计算获得该剪切带中心部分发生了约10±5 km向东且向上的位移,但根据岩基环带的破 裂,可以推测南部发生的缩短量可达25 km。他们同时指出该期变形至少在95 Ma时已经进行,并且 在不晚于90 Ma的时候发生了右行走滑为主的剪切构造叠加,Wong(2005)亦推导出类似的认识。 该区存在77±5 Ma的冷却事件,显然是由于内华达山脉南部拆离系统边缘快速垮塌的结果(Chapman et al., 2012)。这也可能与Maheo等(2004)在更北部的惠特尼山(Whitney,位于美国加利福尼亚 州东部)侵入岩岩套中观察到的83~79 Ma快速冷却是同一事件。

一些古地磁研究结果表明,内华达山脉位于现在位置以南 700 ± 500 km 的地方 (Housen and Dorsey, 2005);而其他研究者认为,至少在约 83 Ma 以来,其相对于北美大陆的位置几乎没有变化 (Hillhouse and Grommé, 2011)。Kent 和 Irving (2010) 计算的新北美极点增加了约 500 km 的误差, 无论如何,由于较长滞留时间、较高的亚固线,温度可能会改变结构、年龄和磁铁矿颗粒大小,使岩 浆岩的矿物和磁性构造面会发生退磁效应 (Pullaiah et al., 1975)。

5.9.2 盐碱地块

盐碱(Salinian)地块的白垩纪深成岩岩石同时受到了 NE 和 SW 断层的隔离(图 2.5、图 4.1 和

图 5.15) (Ross, 1978),在结构、组分和时空上都与内华达山脉 (Sierra Nevada) 类似,包括整体向 东变年轻 (Ross, 1972; Mattinson, 1978b, 1990; Mattinson and James, 1985)。基于最近收集的 U-Pb 数据, Barth 等 (2003) 认为盐碱地块内的许多深成岩体岩石都与内华达山区内的大教堂峰 (Ca-thedral Peak) 侵入系列有着密切的联系;萨利纳斯山脉 (Sierra de Salinas) 片岩,一般认为与更南部 的兰德-Pelona-Orocopia 片岩对应,像其在南部一样,都逆冲到白垩纪花岗质岩石之下 (Ross, 1976; Jacobson et al., 2011)。



图 5.15 主要岩带和岩层地质简图 (据 Jennings (1977)和 Dumitru et al. (2010)修改) 显示弗朗西斯科 (Franciscan)杂岩、蛇纹岩带和大谷地群序列主要岩带和岩层

中新世之后圣安德烈亚斯断层右旋位移距离约 320 km (Crowell, 1962, 1975, 1981),根据是 23.5 Ma 的 Pinnacles 和 Neenach 火山分离 (Matthews, 1976),圣克鲁斯山脉 (Santa Cruz)中的 La Honda 盆地和 Temblor 山脉中圣华金河 (San Joaquin)盆地的渐新世—中新世海底陆源扇形沉积 (Graham et al., 1989; Critelli and Nilsen, 2000),以及海岸山脉的洛根 (Logan,犹他州) 辉长岩与圣埃米格迪奥 (San Emigdio)山中的 Eagle Rest 山峰的岩石相似等 (Ross, 1970; James et al., 1993),复原了内华达山脉以南大部分的盐碱 (Salinian)地块,但是,即使复原了如圣格雷戈里奥 (San Gregorio) -Hosgri 那样的右旋位移 130 km 的断层系统,仍有许多位于大谷地的西侧 (Ross, 1984)。因此,现在的复原工作是有问题的,并且可以肯定还有其他断层的错断。

## 5.10 大谷地沉积作用和变形作用

据 Ingersoll (2008) 所述,大谷地弧前盆地(图 4.1 和图 5.15)"是地球上弧前盆地最佳的研究 对象与其他弧前盆地相比,它是最典型的"。虽然 Ojakangas (1968) 是第一位研究该盆地沉积作用 的,但该研究是在板块学说出现之前,所以并没有得到很好的解释,后来 Dickinson (1970, 1971, 1976) 建立了普遍接受的弧前模型。盆地的填充物以晚侏罗世—早白垩世到麦斯里希特世(Maastrichtian) 的大谷地群为代表 (DeGraaff-Surpless et al., 2002; Surpless et al., 2006),该群包含厚度 大于 15 km 的连续沉积,不整合地覆盖在板片和角砾岩之上,被覆盖的包括海岸山脉蛇绿岩、蒂黑马 -科卢萨 (Tehama-Colusa) 蛇纹岩带、燧石岩以西和内华达山区基底以东 (Ingersoll, 1982; Hopson et al., 2008)。

根据碎屑矿物学研究, Ingersoll (1983)将大谷地群北部的岩石划分为六个岩相。最老的三个岩相含有大量的沉积岩和变质岩碎片,可能来自克拉马斯(Klamath)和内华达山区北部的地体;而较年轻的中晚白垩世岩相则含有更高比例的深成岩碎片(Ingersoll, 1983)。DeGraaff-Surpless等(2002)对大谷地地层中几个剖面的碎屑锆石进行了U-Pb分析,总体上说,最底部的岩相石头溪(Stony Creek)岩石中保留了来自蒂托阶(Tithonian)的生物地层带,其含有的最年轻锆石年龄从144→135 Ma,而其他单元内最年轻的锆石年龄范围为97~72 Ma,已有证据表明它们在晚白垩世沉积(Surpless et al., 2006)。大谷地群的晚白垩世部分主要由浊积岩组成,沉积环境为南北走向的不对称盆地内的盆地平原、冲积扇、斜坡和大陆架(Ingersoll, 1979),现在主要位于内华达山脉前陆以东和海岸山脉(Coast Range)蛇绿岩-弗朗西斯科杂岩以西之间。

Constenius 等(2000)发现在最下面的两个岩相之间有一个主要不连续(图 5.16),这两个岩相的结构、成分和整体的沉积特征都不相同,并与125±1 Ma的巴列姆阶—阿普第阶(Barremian – Aptian)的界限相吻合(Gradstein et al., 2004)。在大谷地的北部,不连续之下的岩石大都发生了错断、弯曲和局部侵蚀。对该区变形的原因知之甚少,至于它的成因尚无共识(Constenius et al., 2000; Wright and Wyld, 2007; Dumitru et al., 2010)。

在坎帕阶—麦斯里希特阶 (Campanian-Maastrichtian)期间,盆地显著的变化是它的西部突然隆 升,其沉积中心逐渐向东迁移 (图 5.17),沉积体系从深水到浅海至冲积相,古流向从向西转为向南 (Moxon and Graham, 1987; Moxon, 1988; Mitchell et al., 2010)。根据对沉积物成分的分析,表明沉 积物来源几乎没有改变,并且仅有的变化都与盆地西侧的隆升有关 (Almgren, 1984; McGuire, 1988)。来自弗朗西斯科杂岩的深变质碎屑岩位于西部,在麦斯里希特阶期间出现在大谷地群的沉积 物中 (Berkland, 1973)。一些学者 (Wentworth et al., 1984; Unruh et al., 1991; Wakabayashi and Unruh, 1995)认为存在一个向东的构造楔,尽管其活动时间主要集中在古近纪,但明显持续至今 (Unruh and Moores, 1992)。

虽然普遍认为大谷地群的岩石学剖面代表了内华达山脉剥蚀去顶,但其物源并不局限于内华达岩基,并且与北美西部所有的科迪勒拉岩基的侵蚀面相吻合。实际上,Wright和Wyld (2007)研究了从侏罗系顶部到白垩系底部的碎屑锆石,发现大部分具有近980 Ma和1.4~1.6 Ga的年龄,这表明大谷地群西部的岩石有被现在位于墨西哥 Oaxaquia (图2.5)附近的原始沉

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图 5.16 大谷地群的岩相划分 (据 Constenius et al., 2000)

展示 125Ma 不整合和上白垩统分裂和通道,见本文讨论。

CRO—海岸山脉 (Coast Range) 蛇纹岩; ss—砂岩; sh—页岩

积物取代的迹象。

因此,大谷地群的沉积岩反映了盆地于近 125 Ma 发生的剧烈变化,其中晚侏罗世—早白垩世的 沉积物遭受了变形和局部剥蚀,并且盆地西侧在近 80~75 Ma 发生了隆升,沉积相从深水相转为了浅 水相。隆升和剥蚀发生得非常迅速,以至于在约 67 Ma 的麦斯里希特阶期间,弗朗西斯科杂岩中的蓝 片岩都出露到地表,被剥蚀并搬运至盆地 (Berkland, 1973)。

沿克拉马斯山脉东北侧有一个称为霍恩布鲁克(Hornbrook)的残余盆地(图 4.1),一般认为其



(据 Dickinson and Seeley, 1979) 说明盆地轴心向东迁移以及上白垩统沉积相从深海相到浅海相突变。隆升在这里被解释 为由俯冲板片断离后的剥露而产生。

代表了与大谷地群一样的弧前盆地的一个片段(Nilsen, 1986; Kleinhans et al., 1984; Miller et al., 1992)。霍恩布鲁克向北东东倾, 125~85 Ma的陆源碎屑序列残余不整合地覆盖在克拉马斯山脉东部地体之上,并且占据了沿加利福尼亚州-俄勒冈州边界的区域呈弧形分布,自下而上具有从砂岩至粉砂岩的冲积扇沉积基本特征(Nilsen, 1993; Beverly, 2008)。虽然一般认为这些岩石都主要来自克拉马斯,但最近一个采自盆地中的碎屑锆石分析表明其中含有大量的年轻锆石,其年龄比附近任何已知的岩浆岩露头都要年轻,因此除了最下部以外,可能大部分都来自科迪勒拉岩基之一,如内华达岩基或海岸深成杂岩体(Beverly, 2008)。

# 5.11 海岸山脉蛇绿岩

海岸山脉蛇绿岩(图 5.15)是一个年龄在168~161 Ma的被肢解序列,由超镁铁质、辉长质和玄 武质岩石组成,认为其代表了形成于古海底的扩张脊,并且被远端的牛津阶(Oxfordian)凝灰质放 射虫硅质岩和泥岩不整合覆盖,向上变为更接近蒂托阶(Tithonian)的火山碎屑岩相(Hull et al., 1993; Hopson et al., 2008)。古地磁和生物地层学证据表明所有的蛇绿岩残余都形成于古赤道附近, 并且向北移动通过一个非沉积地带,随后发生了火山成因的沉积作用,晚侏罗世(152~144 Ma)的 火山喷发形成了蛇绿岩角砾,并最终被大谷地群中一个厚的侏罗系顶部的硅质碎屑浊积岩围裙所淹没 和埋藏(Pessagno et al., 2000; Hopson et al., 2008)。虽然这些学者认为蛇绿岩形成于开放的大洋, 但有些学者则倾向于认为是弧前(Shervais, 2001; Shervais et al., 2004, 2005)或弧后(Godfrey and Dilek, 2000)背景下的超俯冲作用成因。对于支持或反对各个模型的地球化学理由进行了总结,见 前述参考文献,限于篇幅不一一列出。

值得一提的是剖面上还有一种特殊且还没有得到很好解释的特征,那就是在蛇绿岩带和大谷地群 之间发现的由蛇绿岩碎片组成的蛇绿岩角砾单元,其厚度为0~500 m (Hopson et al., 1981; Robert-

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son, 1990; Hopson et al., 2008)。就像 Hopson 等(2008)的总结中所说的那样,角砾之下的蛇绿岩被一个复杂的断层系统所破坏,由于断层活动时间非常长,导致早期形成并部分胶结的角砾从较高的单元中脱落。在断层活动过程中还伴有基性岩浆活动和热液蚀变作用,并且在老人河(Elder Creek)角砾之下还残留了一个岩墙,其年龄为154±5 Ma(Blake et al., 1987)。Robertson(1990)认为简单的正断层和较高岩块的崩塌是形成角砾岩的原因,但 Hopson 等(2008)发现该模型并不能解释何以形成如此大量的角砾岩碎片,因此他认为角砾岩的形成是受洋壳中转换断层的影响。另一种可能也会形成类似的角砾岩、蚀变作用和岩浆作用,即正断层系统,数千米长且有100~500 m的纵向分割,并与4~9 Ma的碱性玄武质岩浆、红褐色火山凝灰岩和活化玻质碎屑岩共生,如果看看现在的日本海沟,显然其形成是因为在通过800 m高的外部海沟高地之前产生了大洋板块弯曲(Hirano et al., 2001, 2006)。弯曲和转换模型都可以解释远洋火山沉积岩覆盖在角砾之上的现象(Robertson, 1990; Hopson et al., 2008)。弗朗西斯科增生杂岩同样也包含一套碱性辉长岩,并在其并入增生楔之前侵入到沉积岩中(Mattinson and Echeverria, 1980; Mertz et al., 2001),这也可能反映通过了外部海沟高地。

一个以蛇纹岩为基质的被称为蒂黑马-科卢萨(Tehama-Colusa)的蛇纹岩质混杂岩(Hopson and Pessagno, 2005),出露在弗朗西斯科杂岩和沿萨克拉门托(Sacramento)峡谷以西的海岸山脉东侧与大谷地群岩石之间(图 5.15)。通常在地质图上将其作为海岸山脉蛇绿岩的一部分(Jennings, 1977),但是地球化学研究表明该混杂岩代表了被肢解的弗朗西斯科洋壳,并且有地幔物质加入到其深海薄层沉积物中(Shervais and Kimbrough, 1985)。该混杂岩显然只受到了低温热液蚀变的影响,并且通常出现在有边界的断层细小裂缝中,在构造上位于海岸山脉蛇绿岩-大谷地群组合之下,被认为与海岸山脉蛇绿岩相似,代表了起源于古赤道地区的侏罗纪洋壳物质(Hopson and Pessagno, 2005)。

#### 5.12 弗朗西斯科杂岩

弗朗西斯科杂岩出露于加利福尼亚州海岸山脉(图 5.15),并且由于其岩性混杂、高压低温变质 作用以及系统性地向东变质程度逐渐变深,通常被认为是俯冲杂岩的典型实例(Hamilton, 1969a; Ernst, 1970; Blake et al., 1988)。根据 Berkland 等(1972)的划分,杂岩内岩石通常分为三个带: 东部带、中部带和海岸带。

东部带包含两个不同的地体:皮克特峰(Pickett Peak)和约拉波利(Yolla Bolly),两地体中岩石的变质程度都达到了蓝片岩相(Blake and Jones, 1981),并被一个向东倾的逆冲断层所分割。皮克特峰含有两个亚单元,分别为南福克山片岩(SFMS—South Fork Mountain schist)和情人泉组(VSF—Valentine Spring formation),同样也被向东倾的逆冲断层隔开;两个亚单元的变形程度都很高,并位于约拉波利的东侧(Worrall, 1981)。SFMS局部含有一些洋中脊玄武岩(MORB)的逆冲岩片——类似于变质玄武岩,顶部为燧石(Wakabayashi et al., 2010),其中的碎屑锆石年龄为137 Ma,但总体上以细粒为主,其原岩可能为泥岩(Dumitru et al., 2010)。对其变质年龄最准确的估计来自白云母所给出的121Ma的<sup>40</sup>Ar/<sup>39</sup>Ar 坪年龄(Wakabayashi and Dumitru, 2007; Dumitru et al., 2010)。在构造上位于 SFMS 之下的情人泉组含有更加丰富的变质硬砂岩,且其变质程度明显比上述岩石要低,其碎屑锆石年龄也比较年轻(为123 Ma),<sup>40</sup>Ar/<sup>39</sup>Ar 给出的年龄为约117 Ma(Dumitru et al., 2010)。与皮克特峰相比,约拉波利地体含有更多的燧石并且在约119 Ma时(Mertz et al., 2001)被碱性玄武岩岩床所侵入(Blake and Jones, 1981; Isozaki and Blake, 1994)。虽然这通常被认为是海沟岩浆作用的产物,但这种侵入体的成分与5.9 Ma 日本外部扩张外侧的喷出岩浆非常类似,现在被正断层错断且位于日本海沟中(Hirano et al., 2001, 2006)。

中部带主要由构造混杂岩组成,含有从东部带分裂来的蓝片岩碎片和岩块,以及各种高级蓝片岩、榴辉岩和角闪岩等外来岩块,枕状熔岩(约88 Ma)被燧石或灰岩包裹,所有被剪切的黏土质杂基中并含有硬砂岩互层(Blake and Jone, 1981)。一些学者对该带内不同地区的碎屑锆石进行了研

究,所得到的最年轻年龄为110到78 Ma (Snow et al., 2010; Morisani et al., 2005; Tripathy et al., 2005; Joesten et al., 2004)。

海岸带包括弗朗西斯科杂岩西部大部分单元,出露在加利福尼亚州北部的海岸山脉地区(图 25),其中断续分布有含陆源碎屑物的枕状玄武岩、远洋灰岩和少量蓝片岩岩块(Blake et al., 1998)。该带东部的砂岩以长石砂岩为主,含有极少量火山和燧石碎片,并含有浊沸石;而从整体上看,该带并没有新生成的蓝片岩矿物,其变质程度比东部带更低(Blake and Jones, 1981)。根据碎屑锆石和微化石研究,可能形成于始新世到中新世(Evitt and Pierce, 1975; Blake et al., 1988; Tagami and Dumitru, 1996; Snow et al., 2010)。

在一篇具有划时代意义的文章中, Dumitru 等(2010)通过对碎屑锆石的 U-Pb 分析,表明弗朗 西斯科混杂岩的主要沉积期开始于约123 Ma,混杂岩中深变质外来岩块和小的岩块明显更加古老, 其变质年龄在169 Ma(?)到132 Ma之间(Mattinson, 1986; Ancziweicz et al., 2004)。许多外来岩 块为蓝片岩、角闪岩和榴辉岩,并具有阳起石-绿泥石环带,表明其之前曾被蛇纹岩所包围 (Coleman and Lanphere, 1971; Cloos, 1986),因此没有必要将其与混杂岩主体所处的俯冲带环境联 系起来。蒂黑马-科卢萨(Tehama-Colusa)蛇纹岩、海岸山脉蛇绿岩东侧(Hopson et al., 2008)和 内华达山区西部的岩石都向北倾,例如金斯-卡维亚蛇纹石混杂岩(Saleeby, 1977; Saleeby and Sharp, 1980),斯马特维尔-前陆弧地块(Menzies et al., 1980; Dilek, 1989; Day and Bickford, 2004)以及斯莱特溪-结合湖弧带(Edelman et al., 1989a, 1989b; Fagan et al., 2001),介于它们之 间的缝合线可能是外来岩块造成的,但也有可能是中部大谷地之下隐伏断层沿走向运动的结果 (Wright and Wyld, 2007),因此其起因可能比较复杂。

根据对弗朗西斯科内大量深变质岩块年龄的研究,通常认为向东的俯冲开始于约 169~165 Ma (Wakabayashi and Unruh, 1995; Anczkiewicz et al., 2004; Wakbayashi and Dumitru, 2007; Dumitru et al., 2010),但正如上文所讨论的那样,位于上部板块的斯马特维尔杂岩在 159 Ma 之前并没有增生,由此看来在碰撞完成之前,俯冲并不会向西发展到弧的其他边缘,虽然存在这种可能。除了一个石榴子石的 Lu-Hf 年龄之外,之前发表的所有弗朗西斯科岩块的年龄都在 159 Ma 的分析误差之内(Wakabayashi and Dumitru, 2007), Anczkiewicz 等(2004)发表的唯一一个 168 Ma 的较老年龄还是来自<sup>176</sup> Lu 衰变常数不确定的情况下。因此,尽管有来自较老地体的岩块,但弗朗西斯科地体本身的年龄不会早于 159 Ma。但是,一个大致呈南北(N-S)向的带状深成岩体群(图 5.7~图 5.9)于约 140 Ma 侵入到克拉马斯和内华达山前陆地带,表明俯冲有可能开始于 140 Ma 之前(Saleeby et al., 1989; Irwin and Wooden, 2001; Day and Bickford, 2004)。这可能是俯冲所带来的第一次岩浆产物,并最终形成了弗朗西斯科杂岩,但是它们代表的只是一些孤立的和活跃期仅为 2~3 Ma 的岩浆作用脉冲,另一种原因如俯冲板片断裂的可能性则更大。

显然,弗朗西斯科杂岩中最年轻的蓝片岩年龄为科尼亚斯阶到桑托阶 (Coniacian to Santonian), 并且出露在 Burnt 山地体和 Diablo Rang 的 Pacheco 山口处 (图 5.15) (Blake et al., 1985, 1988; Wakabayashi and Unruh, 1995; Wakabayashi and Dumitru, 2007; Ernst et al., 2009a; A. Jayko, 2010, 私人通讯),这表明当时俯冲体制发生了显著的变化。

## 5.13 晚白垩世变形与变质作用

科迪勒拉山系最著名的变形是晚白垩世到始新世有基底卷入的拉勒米隆升和出现在大盆地区段的 落基山地区盆地的形成(图4.1)(Dana, 1896)。隆升形成了以前寒武基底为核部的不对称背斜, 两翼被逆冲断层所错断,或者向深部陡倾构成单斜构造,多数情况下发生倒转,深盆地在邻区隆升的 过程中接受沉积。这种构造特征多数延伸长达数十到数百千米,盆地-隆升之间的构造起伏达 5~12 km,并且涉及到了整个地壳(Grose, 1974; Smithson et al., 1979; Brewer et al., 1982; Rodgers, 1987; Hamilton, 1988b)。

拉勒米特色是在科迪勒拉山系大盆地区段反映为构造样式和沉积作用的根本改变,包括从塞维尔

薄皮变形转变为厚皮变形,以及从浅海陆源盆地沉积为主改变为局部的、孤立的陆相盆地 (Dickinson et al., 1988; Beck et al., 1988)。虽然在两种变形类型之间有过一些空间和时间上的叠 加(Kulik and Schmidt, 1988),但后续的整体特征是连续陆源盆地沉积,之后随着局部沉积中 心形成演化并伴生了厚皮变形,因此在两个变形时期中只有轻微的时间上的重叠(Armstrong, 1968)。

塞维尔褶皱-逆冲带内逆冲作用继续,但相对于早白垩世阶段活动强度大大降低。在怀俄明的扇 形地背斜轴地区(怀俄明凸),主要的逆冲断层为克劳福德(Crawford,科罗拉多州)和阿布萨罗卡 岭(Absaroka,美国蒙大拿州南部怀俄明州西北山岭)逆冲断层,更老的阿普第期-森诺曼期 (Aptian-Cenomanian)逆冲断层被褶皱到更大的背斜中(Yonkee and Weil, 2011)。更南部,较老的 峡谷逆冲断层以及阿普第期-森诺曼期逆冲断层被褶皱到大型复背斜中,并有较小型的逆冲断层活动 和双重构造发育(DeCelles and Coogan, 2006)。

在大盆地的内陆带具有两个阶段的变形,一个发生在侏罗纪,另一个发生在晚白垩世(见举例, Snoke and Miller, 1988)。这两个变形或者至少是两期强烈的正断层活动使得许多细节的复原变得非 常困难,但总体上晚白垩世的变形中含有逆冲断层、后折式褶皱和推覆体等形式 (Camilleri et al., 1997; Snoke et al., 1997; McGrew and Peters, 1997)。

晚白垩世的变形还发生在特哈查比山(Tehachapi,加利福尼亚州)地区。那里 100 Ma 的深成岩体形成了斜卧褶皱并向西逆冲,在 95 Ma 之前被剥露(Wood, 1997; Saleeby et al., 2007, 2008)。 之后,不整合地覆盖在 92 Ma 花岗质岩石之上,上白垩统的 Witnet 组岩石(Chapman et al., 2012)可能也发生了褶皱,并被 92 Ma 的花岗质岩石从南部逆掩(Wood, 1997)。

在内华达山区, 戈达德顶垂体地区岩石在 131 Ma 之后到 90 Ma 之前发生了变形 (Tobisch et al., 1995; Bateman, 1992)。在内华达山区的所有地方, 顶垂体内的大多数白垩纪岩石都发生过至少两次 变形。

内华达山脉西部,大谷地群岩石的褶皱和逆冲变形发生在晚白垩世到古近纪(Unruh et al., 1991),在那时大型海底峡谷切入到较老的岩石中(图 5.16)(Williams et al., 1998)。

## 5.14 兰德-Pelona-Orocopia-Swakane 俯冲杂岩

一般认为俯冲杂岩的代表是一个呈 NW-SE 走向的带状延伸露头,出露范围从蒙特雷(Monterrey,墨西哥一座城市)湾南部到亚利桑那州西南部(Jacobson et al., 2011; Haxel, 2002),该带是由特有的片岩不连续露头组成,分别被命名为萨利纳斯山脉(Sierra de Salinas)、圣埃米格迪奥(San Emigdio)、兰德(Rand)、Pelona 和 Orocopia 片岩(图 5.18)。人们认为该片岩最初位于主逆冲断层之下,但现在一般认为其位于低角度正断层下盘,它们占据剥露区中心位置,岩石主要是长英质的并具有倒转层序,变质矿物组合主要属于钠长石-绿帘石角闪岩相岩石,人们认为这些岩石原岩是硬砂岩、砂岩、角岩、泥岩和玄武岩经过变形和变质作用所形成的(Haxel and Dillon, 1978; Ehlig, 1981; Frost et al., 1982; Malin et al., 1995; Oyarzabal et al., 1997; Wood and Saleeby, 1998; Haxel et al., 2002; Jacobson et al., 1986, 2002, 2007, 2011)。在现有的年龄分布上,一般原岩年龄与剥蚀年龄在约 5 Ma 范围内,片岩的原岩年龄和就位年龄均从西南到东北逐渐变老,从西南不到 60 Ma 到东北的 90 Ma(Grove et al., 2003b; Jacobson et al., 2011)。当加利福尼亚州南部断层恢复活动时(Powell, 1993; Nourse, 2002),东部的露头就形成了一个从加利福尼亚州南部到亚利桑那州西部的东西向延伸条带(图 4.1)。

虽然最近学者们将所有的片岩综合在一起构成了一个连续的北东—南西向递减的年龄(Jacobson et al., 2011),但有可能存在两个甚至三个不同的变形时期。在大多数前圣安德烈亚斯构造重建中,能够将兰德(Rand)、圣埃米格迪奥(San Emigdio)、萨利纳斯山脉(Sierra de Salinas)和门户脊



图 5.18 俯冲片岩分布简图 (据 Jennings (1977)和 Jacobson et al. (2011)修改)

(Portal Ridge) 露头以及内华达山脉南部的一组片岩和更加年轻的 Orocopia-Pelona 露头区别开来,后者的碎屑锆石年龄主要为早→中白垩世;然而许多东南部的 Pelona-Orocopia 露头的碎屑锆石年龄主要为元古宙和晚白垩世,当然并非全部如此 (Grove et al., 2003b; Jacobson et al., 2011)。此外,圣 埃米格迪奥山和内华达山脉南部的露头年龄要老 10~15 Ma,并且与加洛克断层南部的兰德片岩相比 具有截然不同的碎屑锆石年龄数据,但与卡塔利娜 (Catalina) 片岩在年龄范围和起源方面非常相似 (Jacobson et al., 2011; Chapman, 2011,私人通信)。因此,最好的区分方法可能是根据年龄来划分:前拉勒米片岩:圣卡塔利娜 (Santa Catalina)、圣埃米格迪奥和门户脊;和后拉勒米片岩 (图 5.19)。

出乎意料的是,非常相似的岩石出露于华盛顿州喀斯喀特(Cascades)北部(图4.1)海岸深成 杂岩南端并归属于 Swakane 片麻岩,以往通常却认为这种岩石不属于兰德(Rand)-Pelona-Orocopia 露头(Matzel et al., 2004)。该片麻岩为长英质角闪岩相变质岩(9~12 kbar 和 640~740 ℃),其沉积 原岩所含的碎屑锆石年龄变化范围为161→73 Ma,一个68 Ma 的淡色花岗岩被推测为该片岩部分熔融 的产物,其角闪石<sup>40</sup> Ar/<sup>39</sup> Ar 年龄为 57.9 ± 0.5 Ma (Matzel et al., 2004)。最近,Gatewood 和 Stowell (2012)认为沉积作用发生的时间应该提前到大约 75 Ma,年轻的锆石都是变质锆石,但锆石形态学 研究表明其为岩浆成因,而非变质成因(Bowring, 2012,私人通信)。总之,片麻岩在时间上是与 Pelona-Orocopia 片岩组合最相似的岩石(图 5.19)。



图 5.19 兰德-Swakane-Orocopia-Pelona 片岩及类似岩石的氩同位素和碎屑锆石 U-Pb 年龄 注意,他们已经被分为前拉勒米 (Laramide) 和后拉勒米 (Laramide) 两组,据 Jacobson et al. (2011); Swakane 数据据 Matzel et al. (2004)

# 第6章 索诺兰沙漠分区

## 6.1 横向山脉

横向山脉(Transverse Ranges)是从加州海岸到加州最东南部地区大致呈东西走向的一系列山脉(见图 5.11)。以前,该山脉大多呈南北走向延伸,但在过去的 20 Ma,西部山脉受太平洋板块俯冲的影响顺时针旋转了约 80°~110°(Kamerling and Luyendyk, 1985; Nicholson et al., 1994)。

在圣盖博(San Gabriel)山北部(图 5.11),文森特(Vincent)逆冲断层将 1.7 Ga 的门登霍尔(Mendenhall)麻粒片麻岩和 1.2 Ga 的斜长-正长-辉长杂岩逆冲到了 Pelona 片岩之上(Barth et al., 1995)。May and Walker(1989)将几个白垩纪地体划分到山脉南部,包括由早白垩世粒岩相片麻岩组成的库卡蒙加(Cucamonga)地块以及圣安东尼奥(San Antonio)地体,该地块由晚白垩世深成岩体、糜棱岩和片麻岩构成,且片麻岩伴有角闪岩相变质沉积岩石的顶垂体。

圣贝纳迪诺山(San Bernardino)的大部分岩石都出露在圣安德烈亚斯断层的东侧(图 5.11), 主要由古元古界花岗质岩石和片麻岩以及不整合覆盖在其之上的新元古界和古生代变质沉积岩组成 (Cameron, 1982)。

Powell (1993) 将横向山脉各种中生代岩浆岩 (图 5.11) 划分为三个北西走向的带:东部带、中部带和西部带,东部带以出露年龄为约 165→150 Ma 的贫石英、碱性长石斑岩体为特征,中部带以出露元古宙—古生代基底为特征,西部带主要以 120~85 Ma 的片理化深成岩为特征。最近 Barth 等 (2008a) 和 Needy 等 (2009) 的锆石 U-Pb 年代学测定研究揭示,在圣贝纳迪诺山、小圣贝纳迪诺山和圣盖博山西部有许多年龄在 156 和 149 Ma 之间的晚侏罗世碱性和高钾钙碱性闪长-辉长质至正长-石英二长岩深成岩体 (图 5.11),几个片理化的钙碱性岩体,包括在 Hexie、平托 (Pinto)和小圣贝纳迪诺山脉 (图 5.11) 出露、年龄在 80→74 Ma 之间的片状岩体,以及少量年龄为 181→167 Ma 零星出露的深成岩体。

年龄为159±7 Ma 的玉米泉(Corn Springs)花岗闪长岩出露在秋克华拉山(Chuckwalla)和小秋 克华拉山中,属于横向山脉东部的组成部分(图 5.11)并且被糜棱岩所切断,糜棱岩反过来又被 150 Ma 的斑状花岗岩和闪长岩,以及一个年轻得多的74±6 Ma 花岗闪长岩侵入(Davis et al., 1994)。有些学者认为159 Ma 之后的变形作用广泛存在于整个地区,并且地质压力计测量表明在变形 之后存在突然的挤压后伸展阶段,其伸展的证据主要是150 Ma 岩体的侵入和148 Ma 独立岩脉的侵 位。在秋克华拉山正北的老鹰(Eagle)山中的老鹰山侵入体是一个多相复式侵入体,岩性范围从闪 长岩到英云闪长岩再到二长花岗岩,在165 Ma 侵入到前寒武片麻岩和早古生代变质沉积岩中就位, 并被后期的独立岩脉所切割(Mayo et al., 1998; James, 1989)。

在加利福尼亚州的东南端(图 5.11),在巧克力山(Chocolate)中拆离体的上盘出露厚达 80~100 m 的英安质熔岩,并含有少量硬砂岩夹层,上覆石英砂屑岩并含有少量砾岩层的硅质黏土岩(Haxel et al., 1985; Jacobson et al., 2002)。Haxel等(1985)认为火山岩与亚利桑那州南部圆顶石山中的 Slumgullion 侏罗系变质火山岩和上覆的侏罗系——白垩系麦考伊(McCoy)组沉积岩有关联。

## 6.2 亚利桑那州南部

保罗纪火山岩在索诺兰沙漠的亚利桑那-索诺拉地区出露很普遍,像哈韦沙漠地区一样,该地区 在古近纪—新近纪发生过强烈的扩张,所以露头的分布很广,在广阔的冲积峡谷中被断层分割彼此分 开(图 6.1)。该区是在北西西走向的菲尼克斯(phoenix)断层北部,该断层将延伸区从科罗拉多高 50 原的转换带中分离出来(Hildebrand, 2009)。

在亚利桑那州中西部,中生代火山岩都经历了白垩纪逆冲和古近纪—新近纪伸展变形,主要出露 在几个地区,或者在拆离断层的上盘,或者在逆冲推覆体内(Reynolds et al., 1987, 1989; Richard et al., 1987)。以下描述基于这些研究,在罗海德山(Rawhide)(图 6.1),155 Ma 的 Planet 火山岩 位于 Buckskin-罗海德拆离断层的上盘,由 600 m 厚的各种变形的变质火山岩组成,如流纹岩和流纹 质凝灰岩,其中含火山碎屑岩和安山质熔岩夹层。在花岗质砂岩(Granite Wash)山(图 6.1),由变 质与复杂变形的浅成斑岩和含火山碎屑岩夹层的流纹岩组成,所有都被推测与麦考伊(McCoy)组相 关的沉积岩所不整合覆盖,在东北部被元古代和侏罗纪结晶岩石组成的逆冲岩席所逆掩。在 Harquahala山(图 6.1),由流纹质-流纹英安质凝灰岩、少量安山质-英安质熔岩和156±10 Ma 黑石 火山岩的表层碎屑岩组成,被浅成斑岩侵入并被与麦考伊(McCoy)组岩石相关的碎屑岩所覆盖,出 露在逆冲断层下盘的底部。



图 6.1 加利福尼亚州东南部、亚利桑那州南部和墨西哥北部地质简图 (据 Mauel et al. (2011)修改) 显示文中讨论的不同山脉以及各种中生代岩石单元的分布。图中墨西哥北部地区还表示了 Anderson and Silver (2005)所 描述的区域地质单元

在加利福尼亚州东南端、亚利桑那州西南部和索诺拉西北部,Tosdal等(1989) 描述了两个含变 质火山岩的序列:早侏罗世(205~170 Ma) 菲涅尔(Fresnal)峡谷地层,由超过7~8 km 的流纹质-英安质凝灰岩、熔岩流以及相关的碎屑岩组成,局部安山质熔岩和浅成斑岩集中;较年轻的阿蒂西亚(Artesia)序列,是一个变化性很大的混合体,由部分属牛津阶的变质沉积岩和变质火山岩组成,如 流纹质-玄武质熔岩和凝灰岩,并被柯维亚(Ko Vaya) 双峰式、碱性深成岩套中159~145 Ma 的深成 岩体所切割。根据成分和年龄,他们还将侏罗纪深成岩(图 6.2) 划分为三个不同的群:①广泛分布

的基特峰-小麦高峰(Kitt Peak-Trigo Peak)超单元,岩性从角闪闪长岩到花岗岩,成分连续,年龄逐渐变新(170~160 Ma);②卡戈穆沙舒(Cargo Muchacho)超单元,由于其在构造上是出露在麦考伊山(McCoy)组之上的逆冲席,故将其作为一个单独的岩套,但其在成分和年龄上(173~159 Ma)与基特峰-小麦高峰超单元相当;③有点双峰式特点的柯维亚超单元,由大量的花岗岩和少量闪长岩组成,所有都在159~145 Ma的年龄范围。



(据 Tosdal et al., 1989,修改) 展示后碰撞柯维亚岩套的双峰性质,认为其代表了曾发生过板片断离岩浆作用;更老的、组分更加连续的 基特峰-小麦高峰、卡戈穆沙舒(Cargo Muchacho)深成岩一般认为与弧岩浆作用有关

厚度大于 8 km 的托帕瓦 (Topawa) 群岩石出露在 Baboquivari 山 (图 6.1), 由约 170 Ma 的流纹 质-英安质火山岩、表层碎屑岩、少量碱性玄武岩和碱性流纹岩组成,全部被基特峰 (Kitt Peak) 岩 套中的 165~159 Ma 的浅成岩体、闪长质-花岗质深成岩体和柯维亚岩套中的一个 146 Ma 的条纹长石 花岗岩体所切割 (Haxel et al., 1980, 1982, 2005)。周围相邻地区也出露几处年龄相似的岩石:向 东到圣路易斯 (San Luis)、Las Guijas 和帕哈利托山 (Pajarito) (图 6.1)为 170±5 Ma 的铜岭 (Cobre Ridge)凝灰岩 (Riggs et al., 1993);在圣里塔山 (Santa Rita) 广泛出露约 183~170 Ma 的莱特森山 (Mount Wrightson)组,由安山质-英安质熔岩和角砾岩组成,其上主要被硅质熔岩和凝灰岩 覆盖;往北东到西雅里塔山 (Sierrita)(图 6.1)是厚达 1.3 km 的牛骨架 (Ox Frame)火山的安山质 熔岩、硅质熔岩和凝灰质岩石,并被一个 175 Ma 的深成岩体所切割 (Cooper, 1971; Spencer et al., 1978),晚侏罗世砾岩覆盖在其之上 (Bilodeau et al., 1987)。沿南南东向国境线到图森 (Tucson)(图 6.1),火山岩 K-Ar 定年大约 177 Ma,黑云母为 169 Ma,可能包括厚层的火山洼地内的流纹质灰 -流凝灰岩,并被格兰斯 (Glance)砾岩所覆盖,该区比斯比 (Bisbee,亚利桑那州)盆地内的基本单元 (Bilodeau, 1979; Bilodeau et al., 1987)。

# 6.3 墨西哥索诺拉省

在索诺拉北部, Anderson 等(2005)根据其工作将侏罗纪岩石划分为四个主要地区:从诺加莱斯(Nogales)(图 6.1)往东为有侏罗纪岩石零散分布的诺加莱斯-卡纳内-纳科札里(Nogales-Cananea-Nacozari)区,其覆盖于或者被推覆到前寒武结晶基底之上,由174 Ma 流纹质-英安质熔岩

和凝灰岩组成,含各种硅质碎屑沉积岩夹层,有些地方被浅成斑岩(174 Ma)和177~173 Ma的深成 岩体所切割。位于诺加莱斯-卡纳内-纳科札里区正西的是巴巴哥(Papago)南部区,该地区是 Topawa-铜岭(Cobre Ridge)-莱特森(Wrightson)和基特峰-小麦高峰岩石向北的延续,Haxel等 (1984)首次发现该区有大量的侏罗纪岩浆作用,并且没有出露前寒武基底岩石,其含有厚层状的流 纹质凝灰岩、熔岩、浅成斑岩和年龄测定为176~166 Ma的深成岩。莫哈韦沙漠-索诺拉区,是一个 已遭受过强烈变形的北西-南东向侏罗纪岩带,包括侧卧褶皱和位于陡倾糜棱岩北侧的逆冲断层,人 们认为其代表了莫哈韦沙漠-索诺拉巨横推断层。在约160 Ma之后与早白垩世之前的巨横推断层的移 动使得更多东部地区的岩石发生变形。卡沃尔卡(Caborea)区由覆盖在1.8~1.7 Ga结晶基底上的沉 积层序组成,含有少量可能为侏罗纪(未测年)的火山岩岩石露头,但在加利福尼亚州海湾附近有 三处发生变形和变质的岩石,其年龄分别为164、153和141 Ma,在该区东南侧是晚侏罗世库库尔佩 (Cucurpe)组,不整合覆盖在约700 m厚的中侏罗弧岩石组合之上,该岩石组合中英安质灰流凝灰岩 获得的年龄为168 Ma,由海相沉积岩和底部152~150 Ma 薄层硅质凝灰岩组成(Mauel et al., 2011)。

从杜兰戈(Durango,美国科罗拉多州西南部)附近向东南一直延伸到墨西哥海湾的是一个侏罗 纪变质火山岩区(图 2.5),通常被称为纳萨斯(Nazas)火山岩区(Barboza-Gudiño et al., 1998, 2008; Bartolini, 1998; Bartolini et al., 2003; Godínez-Urban et al., 2011)。火山岩层序推测可能是 被莫哈韦沙漠-索诺拉巨横推断层所分割的索诺兰沙漠北部岩石的延伸(James et al., 1993; Blickwede, 2001; Barboza-Gudiño et al., 2008),牛津阶(Oxfordian)沉积岩石不整合覆盖在其之上, 火山岩以厚达 3 km 的火山碎屑岩为主,底部含少量安山质-英安质熔岩、凝灰岩和火山碎屑岩夹层, 火山岩被一个年龄为 158 ± 4 Ma 且不同程度变形的流纹斑岩所侵入切割。与此不同的是,基于大量 的古地磁研究(Molina-Garza and Geissman, 1999)或与泛非期(Pan-African)岩石相似的碎屑锆石 年龄(Godínez-Urban et al., 2011)并不支持这一可能性。

#### 6.4 下加利福尼亚

下加利福尼亚(Baja California)位于墨西哥西部的一个多山半岛,延至太平洋和美国边界南部的加利福尼亚湾之间东南偏南。加利福尼亚州西南部和墨西哥下加利福尼亚半岛的侏罗纪岩石不是很著名且相对稀少,但是大量的残余露头存在表明在白垩纪岩浆作用之前该地区是一个重要的组成部分。在圣迭戈(San Diego)东部和东北部,侏罗纪深成岩体均为过铝质和铝质岩石,岩体普遍发育片麻理,年龄在170~160 Ma之间,大多侵入到晚三叠世—侏罗纪变质沉积岩中,现在呈陡倾产出,呈现一个45 km 宽、至少150 km 长的带 (Girty et al., 1993b; Shaw et al., 2003)。在下加利福尼亚半岛北部的圣佩德罗马蒂尔山脉 (Sierra San Pedro Martir)南端, Schmidt and Paterson (2002)填图和测定2个不同地点的黑云母正片麻岩年龄为164 Ma,分别在半岛山脉 (Peninsular Range)岩基的中部和正片麻岩分布广泛的东部区域。

在半岛中部的埃尔阿尔科州(El Arco)界线附近,出露有约6km的绿片岩相侏罗纪岩石,包括 安山质熔岩流、角砾岩、火山碎屑岩和连续的等倾斜褶皱沉积岩,被一个年龄为165±7Ma的矿化 花岗闪长斑岩所侵入(Valencia et al., 2006; Weber and Martínez, 2006)。另外,整个半岛的侏罗纪 岩石露头分布零散(Kimbrough, 2010,私人通信)。

#### 6.5 圣盖博-卡沃尔卡地块

位于 Pelona-Orocopia 片岩重建带南侧的岩石 (图 4.1) 可能代表了一个单独的地体或者地体群, 因为正如前文所说,片岩通常代表了俯冲杂岩的一部分,因此将该地块与北部的岩石区分出来。该区 许多已被新生代断层所破坏,但几次重建帮助确认南部地区在断裂破坏前是同一个地块 (Powell, 1993; Nourse, 2002)。因此其囊括了圣盖博、Orocopia、巧克力山 (Chocolate)及其南部地区 (图 4.1 和图 5.10)。

在圣盖博山, 覆于 Pelona 片岩之上(Ehig, 1982)的是发生了多期变形和变质、年龄 1.7~1.6

Ga 的前寒武片麻岩,并被年龄为1.2 Ga 的斜长-正长-辉长杂岩、早三叠世花岗闪长岩和各种各样的 白垩纪深成岩体侵入 (Dibblee, 1968, 1982; Ehlig, 1975, 1981; May and Walker, 1989; Powell, 1993; Nourse, 2002)。含紫苏辉石英云闪长片麻岩给出的锆石 U-Pb 年龄为 88~84 Ma,而未变形的 黑云母花岗岩获得的 U-Pb 年龄为 78±8 Ma (May and Walker, 1989)。从大多数白垩纪侵入单元获得 的黑云母 K-Ar 年龄都落在了 78~57 Ma 之间 (Miller and Morton, 1980; Mahaffie and Dokka, 1986)。这些年龄组合表明在约 80 Ma 的主变形事件后,紧接着有一个快速折返时期。

往东,在索诺拉-亚利桑那前缘的 Sierra los Alacranes 和 Sierra El Choclo Duro (图 6.1) 是一个 1.7~1.6 Ga 的前寒武片麻岩、高角闪岩相变质沉积岩地块,被 1.45 Ga 的斑状花岗岩和几个晚白垩 世石英闪长岩、花岗闪长岩和花岗岩侵入体所切割 (Nourse et al., 2005)。这些学者认为这些岩石向 南继续延伸并与卡沃尔卡 (Caborca) 地块内 1.78~1.69 Ga 的基底岩石有关 (Iriondo, 2001; Premo et al., 2003; Anderson and Silver, 2005),但在更北边的马扎察尔省 (Mazatzal),岩石的变形年龄不 同。已知卡沃尔卡地块含有一些年龄为约 1100 Ma 的斜长质杂岩 (Espinoza et al., 2003)。在早三叠 世时期,卡沃尔卡地块显然位于北纬~21°±4°,相对于早三叠世参考极点该古北极为顺时针旋转,表明相对于北美大陆早期的左旋位移,紧随其后的是年轻的右旋运动 (Steiner et al., 2005)。根据 Kent 和 Irving (2010)的早三叠世北美古坐标,确定的古纬度相当于现在的太平洋西北部位置。

## 6.6 格雷罗和其他墨西哥地块

墨西哥由多个地体组成,就北美大陆而言大多数都属于外来体(Campa and Coney, 1983)。格雷 罗(Guerrero)地体是一个大型复合地体,占据了墨西哥大陆的中部和西部的大部分地区,该地体由 五个分散的地体组成: Teloloapan、瓜纳华托(Guanajuato)、阿西利亚岛(Arcelia)、Tahue 和芝华塔 尼欧(Zihuatanejo) (图 2.5),所有的地体都具有侏罗纪末到白垩纪的火山岩连续层序,位于 Oaxaquia 和密斯特克(Mixteca)地体的西部(Centeno-García et al., 2008)。

Oaxaquia (Ortega-Gutiérrez et al., 1995)包含一个由变质斜长岩、正片麻岩和紫苏花岗岩组成的 前寒武纪格林威尔 (Grenville)基底 (Ruiz et al., 1988; Keppie et al., 2001, 2003; Solari et al., 2003; Ortega-Obregón et al., 2003),被以外来体为主的古生代沉积地层不整合覆盖,包括非劳伦系 生物相和二叠纪火山岩及其相关的沉积岩岩石,沿其西部边缘被连续的厚层状浊积岩所覆盖,所有岩 石在基默里奇阶 (Kimmeridgian)火山岩喷发和沉积之前都发生了强烈的变形 (Centeno-García and Silva-Romo, 1997; Jones et al., 1995; Barboza-Gudiño et al., 2004)。尽管瓦哈卡 (Oaxaca)地体有可 能来自北美大陆东部,根据其北西方向近垂直于北美格林威尔、外来动物群以及强烈的古生代变形, Oaxaquia 可能在古生代的大部分时间内都与北美大陆具有一定的距离 (Ortega-Gutiérrez et al., 1995)。该地体西部边的界一个明显标志是地壳厚度从 40 km 减小到以西的 20 km (Delgado-Argote et al., 1992)。一个称为 Juchatengo 的地体是具有类洋中脊玄武岩 (MORB-like)特征的二叠纪岩浆岩 小地体,位于 Oaxaquia 的西南侧,可能代表了在 290~219 Ma 深成岩侵位之前存在一个短期的 (short -lived) 古生代弧后盆地发展阶段 (Grajales-Nishimura et al., 1999)。

墨西哥地体(图2.5)含有前密西西比纪多期变形的变质岩(Ruiz et al., 1988; Yañez et al., 1991; Ortega-Gutiérrez et al., 1999),被二叠纪沉积岩以及中侏罗世火山岩和沉积岩所不整合覆盖(García-Díaz et al., 2004),沿其西部与格雷罗复合地体接触带是一套变形和变质的火山岩及沉积岩, 锆石 U-Pb 年龄为约 130 Ma (Campa Uranga and Iriondo, 2003, 2004),所含的碎屑锆石与变质基底大体相似(Talavera-Mendoza et al., 2007),不整合覆盖在变质岩之上的是阿尔必期-土仑期地台相碳酸盐岩和向东不断增厚的土仑期-古新世前渊沉积,并且发育同时代的向东的逆冲断层(Cerca et al., 2010)。在晚白垩世—古近纪主要是向东的褶皱-逆冲断层带中,Oaxaquia和密斯特克(Mixteca)地体都发生了变形(Suter, 1984, 1987; Hennings, 1994)。在逆冲带的东侧前缘形成了一个深海槽(图 2.5),被称为坦皮科-米桑特拉(Tampico-Misantla)前渊(Busch and Gavela, 1978)。

阿西利亚岛、瓜纳华托和 Teloloapan 地体 (图 2.5) 均为构造岩片,含有厚的白垩纪大洋岩石和

类岛弧型岩石(arc-like rock)岩片,向东逆冲推覆到 Oaxaquia 和密斯特克地体之上(Centeno-García et al., 2008)。Teloloapan 地体西部被阿西利亚岛地体所逆掩,含有约 3000 m 的玄武质-安山质 熔岩和角砾岩,下部含有早白垩世硅质碎屑岩,上部含阿普第期灰岩夹层,所有都被阿尔必期-土仑 期海相沉积岩覆盖(Monod et al., 2000; Talavera-Mendoza et al., 2007; Cerca et al., 2010)。阿西利 亚岛(Arcelia)地体由 2 km 厚的阿尔布期-森诺曼期拉班玄武质枕状熔岩和角砾岩组成,上覆含放射 虫硅质岩和页岩互层以及蛇纹岩小块体,一致认为它代表了大洋岛弧地体的一部分(Delgado-Argote et al., 1992; Elfas-Herrera et al., 2000; Mendoza and Suastegui, 2000)。在地体中向东逆冲推覆明显 与主要为科尼亚斯期—坎帕期(Coniacian-Campanian,分别为晚白垩世的第三个时期和倒数第二个时 期)的厚层红层沉积特征相符,其间夹有 84 Ma 的熔岩和砾岩,砾岩中所含有的一个安山质碎屑岩年 龄为 74 Ma(Martini et al., 2009; Martini and Ferrari, 2011)。

瓜纳华托地体(图 2.5),由辉长岩、英云闪长岩和超镁铁质岩石的逆冲碎片组成,向北北东逆 冲越过一个等斜褶皱、双峰式火山岩套和早白垩世复理石包,发生在阿普第期—阿尔必期碳酸盐岩不 整合覆盖在其上之前(Lapierre et al., 1992; Ortiz-Hernandez et al., 2003)。在圣米格尔德阿连德 (San Miguel de Allende),上阿普第期含远洋沉积岩石夹层的钙碱性玄武岩和玄武质安山岩在构造上 位于 Oaxaquia 西部边缘(Ortiz-Hernández et al., 2002)。

格雷罗复合地体最西部的两个地体——Tahue 和芝华塔尼欧(Zihuatanejo)地体(图 2.5)显然 具有相同的早白垩世历史。Tahue 地体是两个最西北部的地体之一,由一个变形和变质的奥陶纪弧地 体组成,被变形的宾夕法尼亚纪—二叠纪浊积岩所不整合覆盖(Centeno-García, 2005)。在某些地方 可能由于构造的原因,这些岩石被白垩纪弧火山岩岩石所覆盖,并被相关的基性-中性深成岩体所侵 入(Ortega-Gutiérrez et al., 1979; Henry and Fredrikson, 1987; Centeno-García et al., 2008)。芝华塔 尼欧地体由晚侏罗世硅质熔岩和 163~155 Ma 深成岩组成,分别覆盖和侵入到一个三叠纪增生杂岩 中,该增生杂岩由含枕状熔岩、角岩、蛇纹岩和灰岩岩块的复理石混杂岩组成,全部被变形的早白垩 世基性硅质熔岩、沉积岩所不整合覆盖,并被 105 Ma 及更年轻的深成岩体侵入(Centeno-García et al., 2008, 2011)。

有些学者认为格雷罗复合地体的中生代火山沉积岩岩石是沉积在减薄的北美克拉通地壳上(Cerca et al., 2007; Centeno-García et al., 2008; Martini et al., 2009),但据我所知无论在露头或者同位素方面都没有支持这种古老基底存在的证据。在卢比亚超级地体内有比北美西部更加典型的晚侏罗世—早白垩世火山岩广泛分布区域,很可能代表了一个复合岛弧 (Tardy et al., 1994)。此外,最近从芝华塔尼欧地体获得的碎屑锆石年龄主要在 110~105 Ma,并在约 1000 Ma 和 560~590 Ma 形成较小的年龄峰值 (Centeno-García et al., 2011),表明相对于劳伦古大陆,其与卢比亚关系更密切。

在芝华塔尼欧地体南部,沿海岸带的南马德雷山脉(Sierra Madre del Sur)有一个 Xolapa 杂岩体(图 2.5),是一个约 50 km 宽、650 km 长的高度混合岩化的正片麻岩地体,至少有一部分来自于格伦维尔期的地壳,被年龄为 160~136 Ma 的变形深成岩体和大量年龄为 66~46 Ma 的岩体所切割(Campa and Coney, 1983; Herrmann et al., 1994; Ducea et al., 2004)。杂岩体的北部边界明显是以大量糜棱岩为标志,指示顶部向北西方向运动,更加年轻的证据是始新世向西南逆冲和走滑断层运动(Nieto-Samaniego et al., 2006; Solari et al., 2007)。

在 Xolapa 杂岩和密斯特克地体东侧是玛雅(Mayan)地体(图 2.5),包含一个向西南倾的白垩 纪地台相碳酸盐岩沉积(科尔多瓦地台),位于侏罗纪陆相地层之上,在科尼亚斯期-马斯特里赫特 期 Zongolica 逆冲褶皱带中,沿东北部边缘的逆冲断层被启莫里奇期-始新世盆地相沉积岩所覆盖,该 盆地内所含的岩石组合,下部为海底玄武岩、砂岩、页岩和砾岩,向上变为灰岩、砂岩和泥岩,顶部 为复理石(Nieto-Samaniego et al., 2006)。玛雅地块南部包含的恰帕斯(Chiapas)地块(图 2.5)是 一个长条形地体,由二叠纪期间变质的古生代深成岩体组成,被晚白垩世—古近纪的深成岩体所切割 (de Cserna, 1989; Burkart, 1994)。在伯利兹(Belize)玛雅山脉,包含变质变形的宾夕法尼亚纪— 二叠纪沉积岩,不整合覆盖在晚志留世深成岩基底之上(Steiner and Walker, 1996)。沿着地体的西部边界是狭长的 Cuicateco 地体(图 2.5),主要由麦斯里希特阶片岩、绿岩、辉长岩和具有蛇绿岩特征的蛇纹岩组成,在晚白垩世—古新世期间,该地体向东逆冲到玛雅地体的红层沉积之上(Pérez-Gutiérrez et al., 2009)。

在北美大陆最南部, 危地马拉(Guatamalan) 缝合带杂岩以北是一个向西倾的白垩纪被动地台陆 缘, 位于玛雅地块的中元古界—三叠纪基底之上, 在坎帕阶的最上面沉降接受沉积, 在麦斯里希特阶 一达宁阶时期被造山带复理石埋藏(Fourcade et al., 1994), 并被超基性岩推覆体逆掩。在 Chuacús 杂岩中, 板块下部的结晶基底岩石在 76 Ma 时变质达到榴辉岩相, 这暗示北美大陆边缘部分大概在那 个时间俯冲到了至少 60 km 的深度, 并在 1 Ma 之后折返回到角闪岩相(Martens et al., 2012), 推测 可能是俯冲板片断离所致。这种变形变质作用通常归因于北美大陆板块试图向西俯冲到一个弧地体之 下, 该弧地体一般被称为加勒比大圆弧 (Pindell and Dewey, 1982; Burke, 1988; Pindell et al., 1988; Donnelly et al., 1990; Rosenfeld, 1993; Burkart, 1994)。

#### 6.7 加勒比大圆弧

自从 Wilson (1966)提出安替列群岛(Antillean)和斯科舍(Scotian)弧来自沿转换断层的太平 洋以来,不同学者认为它们是中生代时期太平洋盆地内分布更为广泛的岛弧的一部分(Moores, 1970; Malfait and Dinkelman, 1972; Burke, 1988; Pindell, 1990; Pindell and Kennan, 2009; Wright and Wyld, 2011)。Burke (1988)提出"加勒比大圆弧"这一概念,认为安替列群岛弧只是更大的 中生代扩张弧的一部分,并在晚白垩世时期与中美洲、墨西哥和南美北部的岩石碰撞。虽然某些学者 (Pindell et al., 2005; García-Casco et al., 2008)最近指出向西倾的俯冲带在 120~115 Ma 时就已经 存在,Burke (1988)观点的核心是该弧最初俯冲于向东倾的俯冲带之上,但在 85~80 Ma 与大洋高 原碰撞之后快速改为向西倾俯冲,该高原在约 90 Ma 形成于加拉帕戈斯群岛(Galápagos)热点之上 (Vallejo et al., 2006)。岛弧和大洋高原都进入到了大西洋当中,现在它们出现在加勒比海的内部及 周边。Sinton et al. (1997, 1998)和 Kerr et al. (2003)对加勒比海中残余的大洋高原进行了研究, 他们还研究了白垩纪岛弧碎片。在 Honduras-Nicaragua 的边境地区,一个 350 km 长的岛弧段,在当 地被称为休纳(Siuna)地体,在约 80~75 Ma 时侵位在 Chortis 地块向北发散的科隆(Colon)逆冲褶 皱带之上(Venable, 1994; Rogers et al., 2007)。

大圆弧及其附带的大洋高原还与南美的西北部发生了碰撞,显然在约75 Ma的晚坎帕阶达到了厄瓜多尔(Ecuador)(Jaillard et al., 2004; Luzieux et al., 2006)。在委内瑞拉和哥伦比亚,许多走滑断层在碰撞及碰撞后将外来体切成岩片并将它们向北部和东部运移(Altamira-Areyán, 2009)。

#### 6.8 半岛山脉岩基

白垩纪半岛山脉岩基(图 2.5和图 4.1),从加利福尼亚州南部一直延伸到墨西哥下加利福尼 亚半岛,延伸达 800 km (Gastil et al., 1975),由两个在岩相学、时间和空间明显不同的岩浆岩套 组成,每个具有不同的基底岩石 (Gastil, 1975;Gromet and Silver, 1987;Silver and Chappell, 1988;Walawender et al., 1990;Gastil et al., 1990)。较老的变形深成岩体年龄范围在约 140→105 Ma 之间,由许多环状杂岩体组成,主要出露在西部;在北部大致侵入到同时代的圣地亚哥峰 (Santiago Peak) 127~116 Ma火山岩和火山碎屑岩石中,在下加利福尼亚半岛侵入到 120~110 Ma 的浅-中深海沉积的阿尔图斯(Alisitos)群;岩石成分范围从辉长岩到二长花岗岩(Gastil, 1975; Johnson et al., 2002;Wetmore et al., 2003;Busby et al., 2006;Kimbrough et al., 2001;Gray et al., 2002)。被称为La Posta 岩套的东部岩群,年龄范围从 99~92 Ma,岩石组成由外带向内带分别由含角 闪石英云闪长岩、二云花岗闪长岩和二长花岗岩构成(Clinkenbeard and Walawender, 1989;Walawender et al., 1990),侵位深度为 5~20 km,侵入到低绿片岩相到角闪岩相变质围岩中,在许多地区 都发生了混合岩化(Gastil et al., 1975;Todd et al., 1988, 2003;Grove, 1993;Rothstein, 1997, 2003)。在 La Posta 岩套侵位之前、西部深成岩系列侵位之后,侵入岩及其各自的基底岩沿着一群逆冲断层平行排列,并在圣佩德罗马蒂尔山脉形成一个双向扇形构造,导致许多学者根据阿尔图斯外来 弧推导出了各种碰撞模型 (Johnson et al., 1999a; Schmidt and Paterson, 2002; Schmidt et al., 2002; Wetmore et al., 2002; Alsleben et al., 2008)。其他学者 (Gastil, 1993; Busby et al., 1998)则认为 是岛弧内的弧后扩张和形成"边缘弧" (fringing arc),或者基于美国境内扇形的连续深成岩作用,完 全没有必要中断岩浆活动 (Todd et al., 2003)。但是,古地磁数据表明相对于北美古极点,西部带经 历了长距离的漂移 (11°±4°,向北),而 La Posta 岩套则大致产出在极点附近 (Symons et al., 2003)。

紧随着这次变形事件, La Posta 深成岩体从 99→92 Ma 侵入就位,并主要侵位在缝合带的东侧, 在局部切断了缝合带并侵入到阿尔图斯弧西侧的岩石中(Silver and Chappell, 1988; Kimbrough et al., 2001)。La Posta 深成岩及其围岩的剥蚀发生在两个不连续的阶段:①森诺曼期—土伦期阶段;②晚 白垩世坎帕期—麦斯里希特期阶段(Grove, 1993; Lovera et al., 1999; Kimbrough et al., 2001; Grove et al., 2003 a)。第一阶段,时间上可能与深成岩的侵位重叠,由于受拆离断层的影响,深度为 10 km 的岩石被带到了陆上并且崩塌,与盆地中早森诺曼期-土伦期粗粒碎屑沉积脉动相符,西部盆 地包含 100~90 Ma 碎屑锆石年龄(George and Dokka, 1994; Lovera et al., 1999; Kimbrough et al., 2001)。根据它们的年龄, Grove 等(2003)将较年轻的冷却年龄与拉勒米事件联系起来。

# 6.9 莫哈韦沙漠和索诺兰沙漠白垩纪岩浆作用

在莫哈韦沙漠地区(图 5.11),125~80 Ma的岩浆活动通常情况下比较少见,但还是有许多残存的深成岩体测定年龄在此范围内。在拉斯维加斯南部,加利福尼亚州-亚利桑那州边界附近的克拉克-龙舌兰-伊万帕山(Clark-Mescal Range-Ivanpah),147~142 Ma 深成岩体发生变形且迁移到逆冲断层之上,部分覆盖在100 Ma的 Delfonte火山岩之上,很可能是发生在条顿(Teutonia)岩基93 Ma 侵位阶段之前(Fleck et al., 1994; Walker et al., 1995; Beckerman et al., 1982)。

晚白垩世—古近纪的深成岩与索诺拉地块 80~75 Ma 变形特征形成期间的区域逆冲事件相符。在加利福尼亚州东部的一个很好的例子就是老妇-派尤特山脉(Old Woman-Piute Range)岩基(图 5.11),该岩基的侵位发生在峰期变质和变形作用之后;岩基由铝质和过铝质花岗岩组成,测定的锆石 U-Pb 年龄为 71 ± 1 Ma;在侵位期间及侵位后不久从中地壳水平快速隆升剥蚀,作为证据的是其角闪石<sup>40</sup> Ar/<sup>39</sup> Ar 年龄为 73 ± 2 Ma 和黑云母为 70 ± 2 Ma(Foster et al., 1989; Miller et al., 1990)。在邻近地区,其他相似侵入杂岩包括约 70 Ma 的 Chemehuevi 山深成岩套(图 5.11),是由黑云母花岗岩和石榴子石-二云花岗岩组成的带状杂岩体(John and Wooden, 1990),以及 66.5 ± 2.5 Ma 的 Ireteba 深成岩体,该岩体为石榴子石-二云母过铝质花岗岩,具有高 Sr、Eu 和重稀土元素富集的埃达克型岩石的特征(Kapp et al., 2002)。在惠普尔山(Whipple)(图 5.11)的变质核杂岩,73 Ma 的 Axtel 石英闪长岩同样具有埃达克岩特征(Anderson and Cullers, 1990)。

约70 Ma的科克斯空山(Coxcomb)侵入岩套(Howard, 2002),由大量的石英二长闪长岩、花 岗闪长岩和花岗质深成岩体组成,出露在 Kilbeck 山和科克斯空山(Coxcomb)、Sheephole、卡柳梅特 (Calumet)和金条山(Bullion)等(图 5.11)。它们往往比 80 Ma的较早变形的深成岩更加富含石英 (John, 1981),并显示出某些埃达克质岩石的趋势(Economos, 2011,私人通信)。

出露在横向山脉东部的(图 5.11)是一个代表了古埋深达 22 km 的斜切剖面(Barth et al., 2008b),那里有超过 20 个 82~73 Ma 的深成岩体,包括一个北西走向的杂岩体,少量向北西倾的片状岩体,沿走向延伸了约 150 km (Powell, 1993),大量的由英云闪长岩-花岗闪长岩和黑云母+白云母±石榴子石花岗岩组成的岩体,并有轻微的变形 (Needy et al., 2009)。

晚白垩世—古近纪的岩浆作用同样发生在德克萨斯州西南部,位于里奥格兰德(Rio Grande)的 大转弯(Big Bend)西北部,那里有一个64 Ma 的红山(Red Hills)深成岩体,是一个含铜钼的斑岩 系列(Gilmer et al., 2003)。因此,在美国西南部,从加利福尼亚州到德克萨斯州西南部,该期岩浆 作用的整体走向为近东西向。 在科罗拉多高原南部的索诺拉段,包含另一个代表俯冲板片断离岩浆作用的岩浆岩带(Hildebrand, 2009)。这些包括在亚利桑那州南部和新墨西哥州区域广泛分布的拉勒米侵入岩,以及一条 76~55 Ma 的深成岩带,该带向南一直延伸到墨西哥以西大部分地区(Anderson et al., 1980; Damon et al., 1983; Zimmermann et al., 1988; Titley and Anthony, 1989; Barton et al., 1995; McDowell et al., 2001; Henry et al., 2003; Valencia-Moreno et al., 2006, 2007; Ramos-Velázquez et al., 2008; González-León et al., 2010)。最近在索诺拉北部的锆石 U-Pb 测年结果揭示,塔拉乌马拉(Tarahumara)组,是一个厚度超过2km、混杂着火山碎屑的单元,不整合覆盖在元古界到晚白垩世的逆冲和褶皱岩石之上,其沉积和喷发发生在76~70 Ma 之间,并被测定为70~50 Ma 之间的深成岩体所侵入(McDowell et al., 2001; González-León et al., 2010, 2011)。

## 6.10 俯冲和弧前杂岩

在索诺兰沙漠分区,相当于弗朗西斯科杂岩的岩石很好地出露在卡特琳娜(Catalina)岛(图 2.5 和图 5.18),那里有一套叠瓦状平缓逆冲岩席,从上到下岩石逐渐年轻且变质程度从角闪岩相到 蓝片岩相逐渐变低(Platt, 1975, 1976)。最上部的构造单元主要是含水和被交代的角闪岩,其上部 由包含蛇纹岩和超基性正片麻岩中的变基性岩岩块的混杂岩组成,下部主要为角闪片麻岩和泥质片岩 (Sorensen, 1988)。往下,Platt (1975) 识别出了一套绿片岩,由泥质片岩夹可能为变质火山岩的基 性岩片和少量蛇纹岩体组成。随后,Grove 和 Bebout (1995)将该单元进一步划分为绿帘石-角闪岩、绿帘石-蓝片岩和含硬柱石蓝片岩。最下部的构造单元由各种蓝片岩相变质硬砂岩、块状变质玄武 岩、基性凝灰岩或砂岩、变质燧石岩、含碳泥质片岩和含变质燧石岩、变质火山岩、超基性岩卵石、现在变质为透闪石-铬云母的岩石以及变质闪长岩-变质辉长岩碎屑的砾岩组成(Platt, 1975)。Grove 等 (2008a)报道的从叠瓦状逆冲推覆体岩石中所获得的碎屑锆石 U-Pb 年龄,表明上部角闪岩单元 不会老于 113 ± 3 Ma;绿帘石-蓝片岩为 100 ± 3 Ma,硬柱 石-蓝片岩为 97 ± 3 Ma,混杂岩中的石榴蓝片岩岩块为 135 Ma。

即使在更南端,沿着墨西哥下加利福尼亚半岛西部,弗朗西斯科型岩石出现在两个地区:比斯凯 诺半岛(Vizcaino)-塞德罗斯岛(Cedros)地区和沿着巴伊亚马格达莱纳(Bahia Magdalena)西岸的 岛上(图 2.5)。那里的一系列火山沉积岩通常被认为代表了一个岩浆弧,岩浆活动至少从166±3 Ma 持续到160 Ma,并在白垩纪期间再次活跃,位于晚三叠世—中侏罗世超俯冲蛇绿岩基底之上,换句 话说,在构造上位于一个含蓝片岩的增生杂岩之上,并且可以从一个含深变质岩块的蛇纹石化混杂岩 中将其区分开来(Kimbrough, 1985; Moore, 1985, 1986; Kimbrough and Moore, 2003; Sedlock, 2003)。混杂岩中的岩块主要由170~160 Ma的榴辉岩和角闪岩以及115~95 Ma的蓝片岩岩块组成 (Baldwin and Harrison, 1989, 1992)。上部与蓝片岩带相接触表明其代表了一个正断层并反映了其可 能属于晚白垩世—古近纪的剥露背景(Sedlock, 1996, 1999)。

可能与弧前岩石有关的两个小岩体出现在沿下加利福尼亚半岛西缘、比斯凯诺半岛-塞德罗斯岛 地区和沿马格达莱纳(Magdalena)以及 Almejas 湾西缘更南端的岛上(图 2.5)。在北部地区,可能 有 10 km 厚的上阿尔布期—森诺曼期到科尼亚斯期—麦斯里希特期的硅质碎屑浊积岩,不整合覆盖在 较老的弧-蛇绿岩-增生杂岩之上(Minch et al., 1976; Boles, 1986; Busby-Spera and Boles, 1986; Sedlock, 1993)。往南,有少量的相似岩石出现在圣玛格丽特(Santa Margarite)和马格达莱纳 (Magdalena)岛地区(Rangin, 1978; Blake et al., 1984b; Sedlock, 1993)。从塞德罗斯岛和比斯凯 诺半岛获得的古地磁数据表明,该区岩石沉积时所在的纬度分别为 20°±6°和 16°±7°(Smith and Busby-Spera, 1993)。

# 6.11 晚白垩世变形与变质作用

虽然被新生代走滑断层和正断层破坏和错断,但从索诺拉西部通过莫哈韦沙漠直到横向山和盐碱(Salinian)地块还是能够识别出大量的白垩纪变形变质作用。研究程度最高的地区位于亚利桑那州中

西部和加利福尼亚州东部,是一个高度发育构造和变质结晶基底的弧形区域,并且上覆被称为玛丽亚(Maria)褶皱和逆冲断层(图4.1)的薄层变质沉积岩——由具有与北美地台相似的岩性和地层中的岩石组成(Stone et al., 1983),但与其不一定有关系,或者说其在白垩纪期间才位于现在的位置——经变质到角闪岩相,局部减薄到其原始厚度的1%,平卧褶皱在 84 Ma 到 73 Ma 期间形成(Brown, 1980; Hamilton, 1982; Hoisch et al., 1988; Fletcher and Karlstrom, 1990; Spencer and Reynolds, 1990; Tosdal, 1990; Knapp and Heizler, 1990; Howard et al., 1997; Boettcher et al., 2002; Barth et al., 2004; Salem, 2009)。该带(图 4.1)北面被菲尼克斯(phoenix)断层截断(Hildebrand, 2009),往南,在南北两侧都有逆冲推覆构造,是麦考伊(McCoy)山组的线性逆冲边界带(Tosdal, 1990; Tosdal and Stone, 1994; Barth et al., 2004)。

麦考伊 (McCoy) 山组 (图 4.1), 出露在加利福尼亚州和亚利桑那州南部断层区,具有与大谷 地群极其相似的年龄,但通常被认为是弧后 (Barth et al., 2004; Jacobson et al., 2011)。或断裂环境 (Spencer et al., 2011)。该组具有一个7 km 厚的沉积层序,为冲积扇、河流和浅湖沉积环境,露头呈 东西向横贯加利福尼亚州-亚利桑那州边界,位于 Orocopia 片岩的北侧 (Pelka, 1973; Harding and Coney, 1985; Tosdal and Stone, 1994; Jacobson et al., 2011)。该组岩石不整合覆盖在年龄从 174~155 Ma 的中晚侏罗世火山岩和火山碎屑岩之上 (Fackler-Adams et al., 1997; Barth et al., 2004)。由于南端和北端的岩石都被逆冲断层所掩盖,其厚度不详 (Tosdal, 1990)。详细的碎屑锆石 U-Pb 测 年表明,基性砂岩由于不包含晚于 165 Ma 的锆石,可能属于卡洛夫期;而从最下部砂岩之中获得的 锆石年龄为 109 Ma,在露头顶部附近还发现一个锆石年龄为 84 Ma,并被一个年龄为 73.5±1.3 Ma 的 花岗质深成岩体所切断 (Barth et al., 2004)。

图森(Tucson)南部,向西延伸到 Ajo 和 Organ Pipe 国家纪念碑(图 6.1),是一组晚白垩世— 古近纪的逆冲断层,把前寒武结晶基底推到侏罗纪和白垩纪变质沉积岩、火山岩和深成岩岩石之上, 深成岩岩石具有向西倾的片理和向西南陡倾的线理(Haxel et al., 1984)。这些逆冲断层被年龄范围 从 74→64 Ma 的偏铝质深成岩体和 58~53 Ma 的过铝质岩体所切割(Miller et al., 1992)。Hildebrand (2009)认为年龄范围从 74→64 Ma 的深成岩可能是俯冲板片断离岩套的一部分,并且一直向南延伸 贯穿墨西哥西部,而较年轻的岩体群为后碰撞弧的一部分,同样也向南延伸到墨西哥西部。

往东南部, 白垩纪比斯比(Bisbee) 盆地的岩石形成轴向北西的褶皱, 并被逆断层所破坏(Davis, 1979)。该期的褶皱, 连同逆断层和逆冲断层一道沿走向一直延伸到新墨西哥州的西南部, 并贯穿奇瓦瓦(Chihuahua) 海槽(Corbitt and Woodward, 1970, 1973; Gries and Haenggi, 1970; Haenggi and Gries, 1970; Brown and Clemons, 1983; Lehman, 1991; Hennings, 1994; Clemens, 1998; Hildebrand et al., 2008)。

一个逆冲褶皱带从墨西哥北部一直延伸到瓜地马拉(Guatemala),形成了Oaxaquia(图 2.5)的 东部边缘,并且正如前文所述,其也在墨西哥恰帕斯、危地马拉、洪都拉斯和哥伦比亚等地出露 良好。

向西到加利福尼亚州的莫哈韦沙漠中北部地区,在莫哈韦沙漠的中央出露变质核杂岩的下盘板块(图 5.11),Fletcher 等(2002)的研究表明 105~85 Ma 的白垩纪混合岩和深成岩体侵入使变形复杂化。他们认为一个具独居石和谐年龄(86~84 Ma)的淡色花岗岩体可能同步变形,但是鉴于它含有两个主要的变形组构并且明显褶皱(译者注:前者可能是指 S-C 组构,后者指褶劈理),也可能早于变形期。

向西到达圣贝纳迪诺(San Bernardino,美国加利福尼亚州南部一座城市)山脉的西北端(图 5.11),卡洪山口(Cajon Pass)的钻孔穿透了一个稍微呈叶片状的81~75 Ma 深成岩体,并被浅浅的 倾向于水平的断层所分割(Silver and James, 1988; Silver et al., 1988)。横向山也包含大量的地体, 或明显属于白垩纪混合岩和糜棱岩逆冲带的构造岩片,包括圣盖博山中的库卡蒙加(Cucamonga)和 圣安东尼奥(San Antonio)岩片(图 5.11),其中一个84 Ma的英云闪长岩体发生了变形,而侵入在 其中的一个78±8 Ma 黑云母花岗岩体却没有变形(May, 1989; May and Walker, 1989; Powell,

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1993)。在横向山脉东部和莫哈韦沙漠南部区域的 K-Ar 测年研究结果显示,包括前寒武片麻岩在内的大多数岩石的黑云母年龄都在 70→57 Ma 范围 (Miller and Morton, 1980),说明这一时期存在冷却事件。半岛山脉往南的岩石显然也受到了这次变形事件的影响,并显示出在 80~68 Ma 之间快速剥露的证据 (Grove et al., 2003a)。

# 第7章 加拿大分区

## 7.1 被动大陆边缘

就像本文前面所讨论的那样,在加拿大和北美大陆西面自晚元古界就形成了被动大陆边缘(Bond and Kominz, 1984),由一个浅水地台相沉积层序、厚约1km的晚古生代碳酸盐岩组成,覆盖在分布广泛的Gog 群碎屑岩之上(图4.2),在局部地区覆盖在最上部的新元古界梅雅特(Miette)组之上(Wolberg, 1986; Aitken, 1989)。下古生界地台边缘在主要山岭出露良好,被称为踢马圈(Kicking Horse Rim)(Cook, 1970; Aitken, 1971)。来源不明的细粒深水沉积单元出现在环形礁的西侧,包括著名的波基斯页岩(Burgess shale)(图4.3)。在泥盆纪时期,大陆内部大体上呈拱形,以致于东部地区上泥盆世地台相碳酸盐岩不整合覆盖在志留纪、奥陶纪和晚寒武世的岩石之上(Price, 1981)。石炭系由地台相碳酸盐岩和以硅质碎屑相为主的砂岩组成(Beauchamp et al., 1986; Mamet et al., 1986)。二叠纪地层在大部分的边缘缺失,显然这与普遍存在的较年轻的剥蚀作用有关(Henderson, 1989)。与其他岩石向西北延伸不同,地台的三叠纪岩石主要出露于香草拱形的东侧,在国界附近呈北东走向的宽阔隆起(图5.2),在落基山脉被称为斯普雷(Spray)河组(Porter et al., 1982)。总之,三叠纪岩石以海相-边缘海相碳酸盐岩和硅质碎屑岩为主,并有少量的蒸发岩(Gibson and Barclay, 1989; Poulton, 1989)。

## 7.2 前渊

西加拿大盆地含有科迪勒拉山系前渊的岩石,该前渊通常被认为形成于中侏罗世的某个时期,大 约在 170~160 Ma,随着弗尼组沉积,被 Leckie and Smith (1992)称为第一旋回,但这种解释还存在 问题。该组的最底部——不整合覆盖在向东和向北东不断变老的地台相沉积岩之上——含有 P,O,含 量达30%的磷灰石单元,通常沿着开放海洋的东部边缘形成,因为冷的富营养水受向南自身环流的 影响从冷处上升,这种水流可能受季风或者水流偏转的影响(Poulton and Aitken,1989; Parrish and Curtis, 1982)。具有放射虫的黑色页岩、绿色砂岩和细粒硅质碎屑岩假整合覆盖在磷酸盐单元之上, 可能预示着开放海洋通道的终结,其下部完全缺失膨润土,并被库特尼 (Kootenay) 群的粗粒硅质碎 屑岩、泥煤和煤层所覆盖(Jansa, 1972; Stronach, 1984)。这些岩石沉积在一个古水流向北的、狭 窄的北西走向盆地中,其西部出露最好和最厚部分出现在冒地槽以东约 70 km 的地台边缘(Hamblin and Walker, 1979; Price and Fermor, 1985)。Gibson (1985) 对库特尼群的岩石进行了研究,发现绝 大多数碎屑都来自冒地槽的风化岩石。同样, Ross 等 (2005) 对弗尼和库特尼地区的 Nd 同位素和碎 屑锆石进行了研究,得出结论是两者都与泥盆纪到三叠纪的冒地槽特征吻合。根据深成岩浆活动、平 卧褶皱以及在库特尼地体内岩石剥露 4 km 等,可以推测出当时的地台边缘发生过碰撞 (Monger et al., 1982),并在弗尼-库特尼(Fernie-Kootenay)内发生了下陷和沉积作用,学者们还总结出了一个 分水岭来解释弗尼-库特尼盆地缺失变质作用和深成岩残余的现象(Ross et al., 2005; Evenchick et al., 2007)。直到桑托期—坎帕期(Santonian-Campanian)之前(Larson et al., 2006),或者在盆地 开始沉积后约 80~90 Ma 期间,逆冲断层并没有明显切割弗尼-库特尼,从盆地开始到前渊逆冲似乎 经历了一段特别长的时间。

不整合覆盖在弗尼-库特尼之上——也许年轻多达 25 Ma——的是 Cadomin 组的粗粒沉积岩和砂岩,在 114 Ma 后开始沉积,也可能比 95 Ma 更年轻(Leier and Gehrels, 2011),由砾石和砾岩组成,后者含来自冒地槽的燧石和石英碎屑(McLean, 1977; Leckie and Smith, 1992; Leckie and Cheel,

1997)。如前文所述,Hildebrand (2009)认为产出在加拿大和大盆地的这些内部地台相砾岩,在很大程度上形成于被动大陆边缘抬升和剥蚀的背景下,因为它们越过外缘隆起被搬运至海沟,并且科迪勒拉山系内的前渊沉积作用,随着大陆边缘被拉到海沟,开始沉积覆盖在砾岩之上。砾岩和砂岩被一个厚层的、由硅质碎屑沉积岩组成的碎屑楔形层所覆盖,包括布莱尔莫尔 (Blairmore)和曼维 (Manville)群在内,均由火山岩和深成岩碎屑、中白垩世碎屑锆石组成,并具有比 Cadomin 组岩石高得多的 ε<sub>м</sub>值 (Ross et al., 2005),与西部年轻火山-深成岩物质的隆升和剥蚀一致。

作为区域不整合的一部分,其内部形成的古土壤和峡谷系统切入到布莱尔莫尔和曼维群的中阿尔 必阶(Albian?)岩石中(Leckie et al., 1989)。除了局部的砾岩和砂岩带,森诺曼期—坎帕期页岩作 为间歇性缺氧环境和深水环境证据,构成了该区的主要部分(Leckie and Smith, 1992)。

在坎帕期—古新世,盆地内还有另一个古水流沉积的呈北西向的厚层碎屑楔形层,由砂岩以及安 山质-英安质火山碎屑、低级变质碎屑和泥质岩、碳酸盐岩和砂质岩组成(Mack and Jerzykiewicz, 1989)。这一阶段的沉积显然与冒地槽相地台内的主逆冲阶段有关(Larson et al., 2006),隆升和剥蚀 在约 58 Ma 时结束(Price and Mountjoy, 1970; Sears, 2001; Ross et al., 2005)。在北美地台,约 95 Ma 的克罗斯内斯特(Crowsnest)火山岩(Peterson et al., 1997)被逆冲断层所切断。同样,最西部 布尔若(Bourgeau)主逆冲断层也切割了地台相岩石,并且逆冲推覆到桑托阶--坎帕阶 Wapiabi 组之 上(Price,待刊; Larson et al., 2006)。紧接着一个主侵蚀时期,年龄在始新世到上新世之间的古近 纪—新近纪砂砾岩和砾岩覆盖在亚伯达(Alberta)和萨斯喀彻温(Saskatchewan)之上,砾石最大粒 径达 0.5 m(Leckie and Smith, 1992)。

## 7.3 加拿大地体和超级地体

三十年前, Monger 等(1982)发表了一篇极具挑战性的文章,其中列举了许多令人信服的证据 来证明加拿大科迪勒拉山系内部的许多地体可以归并为两个超级地体:地体 I 和地体 II,两个超级地 体在与北美碰撞之前互为海岸。他们最重要的观点之一是更远的外缘地体逐渐拼合到北美大陆。随 后,这些地体就成为广为人知的山间和岛屿超级地体。就今天的理解来讲,山间地体包括库特尼、斯 莱德(Slide)山、Quesnellia、卡什溪(Cache Creek)和 Stikinia,而岛屿地体包括亚历山大(Alexander)和兰格利亚(Wrangellia)(图 2.5)。他们认为山间地体在三叠纪时就缝合在一起,而岛屿地体 则在白垩纪才完成拼合。他们还认为北美的西缘当时位于奥米尼卡(Omineca)带的西缘(图 2.5)。 即现在所认识的兰格利亚和亚历山大地体在约 309 Ma 时的宾夕法尼亚纪(Pennsylvanian)拼合在一 起(Gardner et al., 1988)。在某些术语中,阿拉斯加的亚历山大、兰格利亚和半岛(Peninsular)地 体统称为兰格利亚的超级地体(见 Nokleberg et al., 2000)。

# 7.4 贝尔特--珀塞尔--温德米尔超群

从爱达荷州往北一直到不列颠哥伦比亚省,新元古界温德米尔群不整合覆盖在近 30 km 厚的中元 古界贝尔特--珀塞尔超群的变质沉积岩和基性岩基之上(Gabrielse, 1972)并且全部含有大型逆冲岩 片(图 2.5 和图 4.1)。结晶基底不知是否来自逆冲岩片,在蒙大拿和加拿大最南端,沿路易斯--埃尔 多拉多--霍德利--斯坦巴克(Lewis-Eldorado-Hoadley-Steinbach)逆冲推覆系统,一个宽 70~110 km、 长约 450 km、厚 14~16 km 的巨形贝尔特(Belt)超群岩片被推覆在北美地台之上——甚至在东部覆 盖在白垩纪前渊盆地的页岩之上(图 7.1 和图 7.2) (Mudge and Earhart, 1980; Mudge, 1982; Sears, 1988, 2001; Cook and van der Velden, 1995; Fuentes et al., 2012)。虽然传统上认为贝尔特-珀塞尔超群的岩石沉积在北美大陆的西缘,位于现在的位置附近(Price and Sears, 2000), Hildebrand(2009)认为相对于北美大陆来说其属于外来体。后续研究中支持外来体模型的数据包括 ①在贝尔特超群的岩石中识别出了一系列 664~486 Ma 的碱性侵入岩,并且属于爱达荷中部的古生代 冒地槽盖层(Lund et al., 2010; Gillerman et al., 2008); ②认识到至少一个深成岩体在晚寒武世很可 能曾发生过蚀顶作用(Link and Thomas, 2009; Link and Janecke, 2009),可作为冒地槽外部的一个 奇怪现象;③在贝尔特-珀塞尔变质沉积岩中发现了北美克拉通的西北部前所未知的1.2~1.0 Ga的 变质和变形作用(Nesheim et al., 2009; Zirakparvar et al., 2010)。

温德米尔超群(图7.1)包括多种多样的粗粒碎屑岩,部分为冰川成因,少量火山岩、页岩、长石砂砾岩以及深水碳酸盐岩(Ross, 1991)。一些学者(Stewart, 1972; Burchfiel and Davis, 1975; Lund, 2008)认为温德米尔超群和同等的、或更老的岩石(Dehler et al., 2010)代表了北美大陆西部边缘的峡谷沉积,但其并不含有大量的火山岩,且有些岩石比发育在被动大陆边缘的要老 84~100 Ma(Lund et al., 2003; Fanning and Link, 2004),它们可能不会保留足够的热量来匹配早古生代沉降的速率(Bond and Kominz, 1984; Devlin and Bond, 1988)。

# 7.5 塞尔温盆地

在塞尔温(Selwyn)盆地内有一个巨大的复合外来体穿过,为一个长约 700 km、宽 200 km 的 Earn、Road River 和 Hyland 群岩石,从新元古代到古生代(图 2.5),沿着道森(Dawson)和破脑壳(Broken Skull)断层逆冲推覆到北美砾岩之上,并且与下盘岩石截然不同(Gordey and Anderson, 1993)。外来体由细粒沉积岩、燧石岩、石灰质浊积岩和笔石页岩组成,并含碱性玄武岩、重晶石分层以及喷流沉积 Ag-Pb-Zn 矿床(Goodfellow et al., 1995; Mair et al., 2006)。由于这些特征,加上整个剖面存在断断续续出露的玄武岩(Goodfellow et al., 1995; Cecile, 2010),与Turner等(1989)首次提出的罗伯茨山外来体的特征非常相似,与被动大陆边缘相比,具有比有限边缘盆地内(如中国南海-台湾海峡)更加典型的沉积、岩浆和蚀变作用(Teng and Lin, 2004; Koski and Hein, 2004)。此外,最近的研究表明(McLeish et al., 2010; McLeish and Johnston, 2011)塞尔温盆地最南部的Kechika凹陷在泥盆纪期间形成了伏卧褶皱,这使得很难将其与北美东部具有相似年龄而未变形的岩石联系起来。

一个早寒武世含古杯动物门的碳酸盐岩地台,被称为卡斯尔地台(图 2.5),出现在塞尔温盆地的西部。年龄为 110~90 Ma 的深成岩套将卡斯尔地台的岩石与塞尔温盆地东部联系在一起(Johnston, 2008)。Johnston (2008)和 Hildebrand (2009)都认为卡斯尔地台与北美地台相比是一个不同的、更老的地台,主要由早寒武世硅质碎屑楔以及覆盖在其之上的中寒武世碳酸盐岩滩组成,在他们的模型中与前文所提到的安特勒大陆架相似。Pope 和 Sears (1997)注意到往南在爱达荷-蒙大拿地区冒地槽相岩石缺失,因此认为加拿大卡斯尔地台沿 Tintina—落基山脉北部的海沟断层向北逃逸。

# 7.6 奥米尼卡带岩浆活动

在加拿大科迪勒拉山系,从育空-塔纳纳(Yukon-Tanana)地体和塞尔温盆地向南一直到路易斯-克拉克线性构造(Lewis & Clark lineament),出露一大片白垩纪深成岩体(图 2.5 和图 5.12)(Monger et al., 1982)。就像南侧的美国西部一样,在早白垩世的约 140~135 Ma 至 120~115 Ma 期间,科 迪勒拉北部存在一个岩浆间歇期(Armstrong, 1988)。在北部,主要岩浆作用从约 120 Ma 爆发一直 持续到 96 Ma,并向东逐渐变年轻,含有铝质和过铝质两种岩体(Hart et al., 2004; Johnston, 2008)。 Anvil-Hyland-卡斯尔岩浆亚带侵入到塞尔温盆地和卡斯尔地台的主沉积岩层,在走向上延伸约 900 km,年龄在 110~96 Ma,有过铝质和铝质两种岩石(Driver et al., 2000; Hart et al., 2004)。在 Anvil -Hyland-卡斯尔岩浆亚带东侧,另一个侵入到塞尔温盆地沉积岩中的线状深成岩带称为 Tombstone-Tungsten带,通常由年龄为 96~90 Ma,具有 Au、Cu、Bi、W、Zn、Sn、Mo 和 Sb 不同矿化的次碱性 到碱性岩体组成(Hart et al., 2005)。较年轻的岩套沿走向延伸超过 1000 km 直到阿拉斯加,分别被 称为 Livengood 和 Fairbanks-Salcha 岩套(Reifenstuhl et al., 1997a, 1997b; Newberry et al., 1990, 1996)。因此,该区的深成岩与其他分区的完全不同。首先,科迪勒拉型岩浆活动终止于约 96 Ma, 而其他地区的岩浆活动终止于 82 Ma 左右, 96~90 Ma 的线性带状含金属的深成岩体与其他地区没有 明显的等同性。


#### 图 7.1 加拿大南部落基山褶皱-逆冲带地质简图

图中表示的温德米尔-贝尔特-珀塞尔超群这里被认为由外来体组成,覆盖在北美大陆被动陆缘岩石之上。北美大陆基底出露在 莫纳希杂岩的 Frenchman 顶部 (FC)和 Thor-Odin 穹窿 (TO)处,为一个侵蚀双重构造,但可能在拉勒米事件中被完好地迁移 到了南部。注意:大量小型超基性侵入体被切的岩石,在这里都认为是踢马圈 (Kicking Horse Rim)西侧的卢比亚超级地体的 一部分,通常认为其代表了面向西的北美寒武纪大陆架边缘的一部分。他们很可能标志着卢比亚-北美缝合带,就如 Burke 等 (2003)、特别是 Johnston 等 (2003)所建议的一般模型中那样。温德米尔-贝尔特南北两侧的外来体,来自更远的外侧,位于 该带且逆冲推覆在其之上,这样构成卢比亚-北美缝合带东界的限制性标志。F—Fang 岩株;FB—弗尼盆地;IR—冰河杂岩; K—Kettle-Grand Forks 穹窿;Kx—Kuskanax 岩基;N—Nelson 岩基;O—Okanagan 穹窿;PR—Priest 河杂岩;



图 7.2 外来体岩石逆冲推覆到白垩纪前渊沉积岩石之上 国家冰川公园中贝尔特-珀塞尔超群岩石沿路易斯(Lewis)逆冲断层向东逆冲推覆到白垩纪前渊沉积岩石之上。 根据本文所描述的模型,路易斯(Lewis)逆冲断层代表了北美和卢比亚超级地体之间的缝合线

## 7.7 库特尼地体

该地体的岩石(图 2.5 和图 7.1)都发生了变质,并构造不整合地覆盖在更东面的温德米尔和珀 塞尔(Purcell)超群外来体之上,主要由各种各样下一中古生代沉积岩岩石组成,包括含古杯动物 门的大理岩——典型的卡斯尔地台、而非北美克拉通的剖面——被奥陶纪—泥盆纪深成岩体侵入。有 这样一个群,称为拉尔多(Lardeau)群,其岩石变质成了石英岩、片岩和片麻岩,并在密西西比阶 的含玄武岩米尔福德(Milford)群沉积之前发生了褶皱作用(Read and Wheeler, 1975; Klepacki, 1985; Klepacki and Wheeler, 1985; Roback, 1993; Paradis et al., 2006)。Smith and Gehrels (1992b) 在内华达州的罗伯茨山外来体中发现了与拉尔多(Lardeau)群相似的岩石。Paradis 等(2006)将老 鹰湾(Eagle Bay)组合中的泥盆纪—密西西比期的火山岩和深成岩岩石阐述为建立在北美西部的弧 地体,该地体内的其他岩石组合,如新元古界马贼溪(Horsethief Creek)群、始寒武统哈米尔(Hamill)群、含古杯动物门的Badshot组都位于珀塞尔(Purcell)主断层西侧,并不认为其覆盖在北美 基底之上,包括更老的变形——向西大型平卧褶皱变形在内(Ross et al., 1985; Brown and Lane, 1988; Simony, 1992; Ferri and Schiarizza, 2006),且不出现珀塞尔(Purcell)断层东侧的岩石,推 测其至少被向东运移了200 km (Price and Mountjoy, 1970),并被晚古生代到白垩纪的各种各样的深 成岩体所侵入(Okulitch et al., 1975; Parrish, 1992; Crowley and Brown, 1994; Colpron et al., 1998)。

位于莫纳希山(Monashee)杂岩东翼的塞尔扣克山(Selkirk)扇形构造(图 2.5 和图 7.1)是一个侵蚀窗,可能出露双重构造:即库特尼地体之下的北美基底与盖层岩石。在变形之前和/或变形过程中,187~173 Ma发生了一次深成岩体侵入活动,并在173~168 Ma时期岩石从7~3 kb的快速剥露之前(Colpron et al., 1996)。类似的是 Scrip 推覆体,向西倾的等斜构造恰好位于其北侧,存在一个走向上长 50~60 km 的倒转岩片,并且很可能是在同一时间形成的(Raeside and Simony, 1983)。正如前文所述,Höy(1977)记录和报道了再往南的相似构造,尽管其年龄约束程度低,明显在178~

164 Ma 之间已经形成(Read and Wheeler, 1975)。Colpron 等(1996, 1998)认为这些事件发生在外部和邻近北美冒地槽的地层中(Colpron and Price, 1995),但在北美克拉通地体内没有地质记录,如 在该时期的主要岩浆作用、褶皱、增厚和剥露等变形或沉积作用的证据。

## 7.8 育空-塔纳纳地体

该地体的基岩出露较差,广泛分布残积碎屑,Colpron 等(2006)将该地体(图 2.5)的岩石划 分成了四个组合:①山顶积雪(Snowcap)组合,代表了该地体中最老的岩石,由多期变形和角闪岩 相变质沉积岩组成,并被晚泥盆世—早密西西比期深成岩体所侵入;②芬雷森(Finlayson)组合,由 晚泥盆世—早密西西比期变质沉积岩和变质火山岩组成;③中密西西比期—早二叠世中基性火山岩和 火山碎屑岩组合,不整合地覆盖在山顶积雪和芬雷森组合之上;④中—晚二叠世钙碱性火山岩及其相 关的克朗代克(Klondike)深成岩体组合。不管如何创建杂岩形成模型——根据大量强烈变形和变质 的火山岩地球化学数据和从变质沉积岩中获得的非均一碎屑锆石剖面——表明前三个岩石组合的岩浆 活动与一个岛弧有关,该岛弧发育在北美西部边缘向东倾的俯冲带之上(Piercey et al., 2006; Piercey and Colpron, 2009),但在地体内还未发现克拉通基底,沉积、岩浆和变形作用与北美被动大 陆边缘的关系也不明确。

## 7.9 斯莱德山地体

该地体的岩石以构造岩片的形式产出,将卡斯尔地台北部的 Quesnellia 和育空-塔纳纳 (Yukon-Tanana) 地体的岩石与库特尼地体南部的岩石 (图 2.5)分开。它们出露的长度超过加拿大分区,并在不同地方具有不同的名称 (Harms, 1986; Struik, 1987; Schiarizza and Preto, 1987; Ferri, 1997)。在北部,它们沿白北鲑 (Inconnu) 逆冲断层向东逆冲到卡斯尔地台的岩石之上 (Murphy et al., 2006; Piercey et al., 2012)。外来体内部复杂,但包括石炭系—早二叠世被肢解的盆地相沉积岩,以及随着晚泥盆世到二叠纪蛇绿岩物质一起的玄武岩 (Nelson, 1993; Roback et al., 1994; Murphy et al., 2006)。Nelson (1993) 论证指出逆冲带中的岩石代表了崩塌洋盆的一个剖面,该洋盆位于卡斯尔下部板块和上冲板块之上的育空-塔纳纳 (Yukon-Tanana) 弧之间。Beranek and Mortensen (2011) 最近的研究表明碰撞时间有可能在 260~253 Ma 之间,并在育空-塔纳纳 (Yukon-Tanana) 弧内发现了几个二叠纪的深成岩体。

## 7.10 Stikinia 和 Quesnellia

在不列颠哥伦比亚省南部,Quesnel 地体由覆盖在三叠纪和晚古生代岩石之上的晚三叠世—早侏 罗世火山岩和沉积岩组成,它们自身又覆盖在 372 Ma 的片麻岩之上(Beatty et al., 2006; Simony et al., 2006)。在该区内,上古生界哈珀牧场(Harper Ranch)群由晚泥盆世—晚密西西比期岛弧岩套 和上覆一个含麦克劳德(McCloud)动物群的二叠纪碳酸盐岩台地组成(Beatty et al., 2006)。晚三叠 世—早侏罗世(227~210 Ma)的尼古拉(Nicola)群位于更老的岩石之上,由各种各样的斜辉石斑 状碱性熔岩、相关碎屑岩、火山碎屑岩和沉积岩夹层组成,与出露在莫哈韦沙漠到育空(Yukon)地 区的岩石具有相似的组分和年龄(Mortimer, 1986, 1987; Monger, 1989; Monger and McMillan, 1989)。Miller(1978)在Quesnellia发现了与美国西部相似的三叠纪碱性深成岩。Quesnellia内外壳 岩层序(supracrustal sequence)的上部含有 204~187 Ma 的罗斯兰(Rossland)群,由不同批次的稀 少碳酸盐岩、由细到粗的碎屑岩,一个以钙碱性基性到中基性火山熔岩流和碎屑岩、以火山碎屑熔岩 流为主的厚层堆积,以及相关的外源碎屑岩和伴生的基性-中基性成分的深成岩体等组成(Tipper, 1984; Andrew and Höy, 1990, 1991; Höy and Dunne, 1997)。

不列颠哥伦比亚省再往北, Stikine 和 Quesnel 地体都含有上三叠统塔克拉(Takla)群的玄武质-安山质火山岩和沉积岩,上覆粗粒沉积岩、英安质和安山质熔岩、碎屑岩、少量早侏罗世黑泽尔顿 (Hazleton)群玄武岩和流纹岩(Monger and Church, 1977),代表了弧内伸展盆地(Thorkelson et al., 1995)。Dostal等(1999)的研究表明 Stikina 和 Quesnellia 两个地体的塔克拉(Takla)群的岩石具有 相似的年龄、组分和岩相,并认为它们形成于一个连续的弧地体。在此之前,也有学者发现了两地体 的相似性,并认为 Stikinia 和 Quesnellia 是在早到中侏罗世期间,围绕着两个地体之间的卡什溪地体, 连同育空-塔纳纳(Yukon-Tanana)地体一同发生了褶皱(Nelson and Mihalynuk, 1993; Mihalynuk et al., 1994);或者 Stikinia 在中一晚侏罗世期间被迫从哥伦比亚湾地区向北逃逸(Wernicke and Klepacki, 1988)。

#### 7.11 卡什溪地体

卡什溪(Cache Creek)地体(图 2.5) 与位于 Stikinia 和 Quesnellia 之间的地理分界线呈断层接 触,它被认为代表了密西西比期到早侏罗世的增生杂岩,包含玄武质海山最上部的碎片以及上覆上二 叠统以含古地中海(Tethvan)动物群为特征的沉积岩(Monger and Ross, 1971)、混杂岩带、含放射 虫硅质岩及凝灰质板岩、硬砂岩、砾岩、页岩、基性-超基性岩浆岩带和蛇绿岩组合,以及长达 75 km, 宽 40 km 的礁碳酸盐岩岩块 (Gabrielse, 1991; Johnston and Borel, 2007)。托尔阶的硅质岩 又受到40 Ar/39 Ar 变质年龄为 174 Ma 的蓝片岩相变质作用叠加, 171 Ma 被剥露提供的碎片邻近 Whitehorse 凹陷 (Thorkelson et al., 1995; Mihalynuk et al., 2004)。卡什溪的岩石受到向南西倾褶皱和北东 倾逆冲断层的影响而变形,局部褶皱向北东倾,其底部为将卡什溪岩石推覆到 Stikinia 之上的向东倾 的逆冲断层 (Struik et al., 2001)。在 Stikinia-卡什溪接触带附近的一个变形深成岩体的年龄为 219 Ma; 一个切断逆冲断层的深成岩体给出了 161 Ma 的年龄; 另一个切过逆冲断层的深成岩体, 位于超 基性岩块底部,其锆石 U-Pb 年龄为 166±2 Ma;一个变形后的深成岩体切割了卡什溪变形岩石,其 U -Pb 年龄为 172 Ma (Mihalynuk et al., 1992; Ash et al., 1993; Struik et al., 2001)。这些年龄与 Stikinia 内的最年轻的侏罗纪深成岩套很好地吻合(MacIntyre et al., 2001);在 Stikinia 出现的巴柔阶 (Bajocian) 燧石卵石砾岩显然来自卡什溪地体 (English and Johnston, 2005); 整体位于 Stikinia 之上的鲍尔斯(Bowser)盆地北部出现了从东部搬运过来的巴柔阶的碎片(Ricketts et al., 1992)。卡什溪地体的东部边界是一系列白垩纪—古近纪—新近纪右旋走滑断层或倾斜走滑断层 (Gabrielse, 1985; Struik et al., 2001)

不整合覆盖在 Stikinia 地体和部分卡什溪地体之上的是鲍尔斯和 Sustut 两个盆地的沉积岩 (Evenchick and Thorkelson, 2005; Ricketts, 2008)。鲍尔斯盆地(图 2.5)沉积了大于6000 m 的中侏 罗统到中白垩统海相和非海相碎屑岩,而 Sustut 盆地含大于 2000 m 的非海相碎屑岩,年龄范围从阿 普弟期—阿布儿期到坎帕期 (Evenchick et al., 2007)。

## 7.12 三叠纪超覆序列

加拿大科迪勒拉山系内的几个地体都有砾岩层,与邻近地体呈超覆接触关系和/或明显包含来自 它们的碎屑,表明在当时它们离得很近。卡斯尔地台的岩石被育空-塔纳纳(Yukon-Tanana)地体二 叠纪的岩石推覆逆掩,在覆盖缝合带的含蓝片岩和榴辉岩碎屑的三叠纪砾岩沉积之前(Murphy et al., 2006; Johnston and Borel, 2007)。这些关系和 U-Pb 年代学数据表明育空-塔纳纳、塞尔温盆地、卡 斯尔地体和斯莱德山大洋板片是在三叠纪拼合在一起的(Beranek and Mortensen, 2007, 2011)。

#### 7.13 海岸山脉深成杂岩

海岸深成杂岩(图 2.5)是一个由离散和复合深成岩体构成的宽 50~175 km 的岩带,岩性范围由

辉长岩到花岗岩,从华盛顿州一直延伸到阿拉斯加 (Brew and Morrell, 1983; Armstrong, 1988; Mahoney et al., 2009),长约1700 km。与半岛地区的岩基类似,并列分布着两个不同年龄的深成岩带: 西部带,位于之前拼合的亚历山大-兰格利亚复合地体:东部带,位于拼合的 Stikinia-Yukon Tanana 地体: 在两者之间以 Gravina-Dezadeash-Nutzotin-Gambier 带为界(图 7.3),由一个中侏罗统—中白 垩统浊流沉积盆地残余、局部含流纹岩(177-168 Ma)、安山质和玄武质熔岩组成(Berg et al., 1972; Rubin and Saleeby, 1991, 1992; Haeussler, 1992; Journeay and Friedman, 1993; Crawford et al., 2000; Manuszak et al., 2007; Gehrels et al., 2009), 在阿拉斯加, 为晚侏罗世—上白垩统卡希尔 特纳盆地浊积岩(Ridgway et al., 2002; Kalbas et al., 2007)。西部带深成岩体侵位明显有间断,分别 是 177~162 Ma、157~142 Ma 和 118~100 Ma; 而东部深成岩是连续地从 180 Ma 到 110 Ma (Gehrels et al., 2009)。一个在100~90 Ma 之间发育的向西倾的逆冲断层(Haeussler, 1992; Rubin et al., 1990), 将东部带的深变质岩石推覆到 Gravina-Dezadeash-Nutzotin 带的浅变质岩石之上,南部的情况可能也 类似,形成了一个向上变质程度逐渐变深的逆冲推覆体 (Journeav and Friedman, 1993; Crawford et al., 2000; McClelland and Mattinson, 2000)。在变形期间, 一个年龄为 100~90 Ma 的含绿帘石和榍石 大晶体的独特岩套侵入到大于 25 km 的深度(Gehrel et al., 2009)。在该带的南部, 北纬 49°30'附 近,每个 92 Ma 之前的深成岩体都发生了明显的变质和褶皱作用(图 7.4 和图 7.5);而 90~84 Ma 的 岩体则没有明显的变质 (Brown et al., 2000; Brown and McClelland, 2000), 但发生了明显的变形和 褶皱 (图 7.6)。晚白垩世—古新世的岩浆作用一般集中在一个线性的窄带里,至少 1000 km 长,位 于海岸带剪切带-海岸山脉地区大型线状构造的东侧,约70 Ma时在北部陡倾的层状英云闪长岩岩基 侵位,并在 60~50 Ma 时侵位到东南部,伴随着 76~67 Ma 的带状岩体的侵位,从边缘到中心,从闪 长岩逐渐过渡到英云闪长岩和花岗闪长岩 (Crawford et al., 1987; Rusmore et al., 2001; Gehrel et al., 2009; Mahoney 等 (2009)。在该区的中南部发育一个向东倾的逆冲带,并且从大约 87 Ma 一直活跃 到 68 Ma (Rusmore and Woodsworth, 1991; Evenchick et al., 2007)。



图 7.3 Nutzotin-Dezadeash-Gravina-Gambier 盆地地层柱状图

(据 Ricketts (2008) 修改)

显示地质单元和年龄的相似性。这些盆地发育在海岛(Insular)与山间(Intermontane)的超地块之间, 在 105~100 Ma 转换挤压事件期间变形



图 7.4 不列颠哥伦比亚省南部海岸深成杂岩中被褶皱的深成岩体 (据 Brown and McClelland (2000); Brown et al. (2000)修改)









7.4



图 7.6 Skuzzum 侵入体边缘陡倾的岩床 内部构造、岩床陡倾和与侵入体的接触关系与该区海岸带深成杂岩的较老岩体相似,表明都发 生了褶皱作用。Ned Brown 摄

## 7.14 白垩纪超覆层序

在不列颠哥伦比亚省中南部,海岛(Insular)和山间(Intermontane)超地体在早白垩世(95~85 Ma)就拼合在一起,两个沉积序列分别称为 Powell Creek 组和 Silverquick 组(图 2.5),构成一个火山-沉积超覆层序不整合地覆盖在两个超级地体之上(Garver, 1992; Mahoney et al., 1992; Schiarizza et al., 1997; Riesterer et al., 2001; Haskin et al., 2003)。超覆层序的古地磁研究包括火山熔岩流内的 26 个点和沉积岩中的 54 个点,连同一些正接触点、砾岩和倾斜测试相结合一起,为地磁存在误差倾向的地区提供一个可靠和健全的记录,获得一个 39.5°±2.2°的古纬度,预计那时候北美位于古纬度约 20.3°±2.7°以南(Enkin et al., 2003; Enkin et al., 2006b; Enkin, 2006; Kent and Irving, 2010)。这些数据与前文所述的从位于育空(Yukon)地区的海岸深成杂岩北端的卡马克斯(Carmacks)火山岩中获得的 70 Ma 古磁极数据相吻合(图 2.5)(Wynne et al., 1998; Johnston et al., 1996; Enkin et al., 2006a)。

## 7.15 晚白垩世—古近纪岩浆活动

在加拿大西部的海岸山脉中, 侏罗纪到晚白垩世的英云闪长岩-花岗闪长岩深成岩体在角闪岩-变粒岩相的条件下发生了变形和变质成为片麻岩, 通常认为其构成了科迪勒拉型岩浆弧的中下地壳, 在 65~60 Ma 时被快速剥露(Armstrong, 1988; Hollister, 1982; van der Heyden, 1992; Crawford et al., 1999)。在 60 Ma 之前发生的伸展作用导致了至少 15 km 的构造剥蚀,并且伴随着大量多种来源 的晚白垩世—古近纪侵入活动高峰(图 5.12)(Hollister and Andronicos, 2000, 2006; Hollister et al., 2008; Mahoney et al., 2009; Andronicos et al., 2003)。强烈隆起的线性轨迹和深成岩浆作用与主断层 或深部地壳边缘相吻合,已被两个岩石探针(LITHOPROBE)深反射地震成像证实(Cook et al., 1992, 2004)。恰好在晚白垩世弧岩浆作用结束之前,中部片麻杂岩在约 60 Ma 时快速剥露,以及大 量的晚白垩世—古近纪岩浆活动,使得 Hildebrand (2009)确定这些岩石应该被称为俯冲板片断离岩 浆作用(slab-failure magmatism),他认为岩浆流穿过俯冲板块被撕裂的狭长缝或许可以解释沿着西 部边缘带高度集中和拉长分布的晚白垩世英云闪长岩体(如 Barker and Arth, 1990)。

#### 7.16 北喀斯喀特

在华盛顿州的北喀斯喀特 (North Cascades) (图 2.5 和图 7.7) 有一个沿南南东方向断续分布的 海岸深成杂岩,此外还有一个与约 100 Ma 变质作用相关的向西倾的逆冲断层带 (Misch, 1966; Mattinson, 1972; McGroder, 1991; Miller et al., 2009),在约 68~59 Ma 之间有一个更年轻的古新世深成 岩体侵位 (Miller et al., 1989)。96~91 Ma 的斯图尔特山 (Mount Stuart) 岩基 (图 7.1) 形成于较老 的逆冲之后,但明显在另外两期变形之前,因为它包裹了具有组成不均一的片理和线理的褶皱鼻,而 该片理和线理穿切了内部接触带,如在混合区域 (Paterson and Miller, 1998; Matzel et al., 2006)。 该区的其他侵入岩套,如 Napeequa 杂岩和喀斯喀特河岩套 (图 7.7) 是在约 88~76 Ma 发生的变形 (Matzel et al., 2004),位于北喀斯喀特 (Cascades)的岩石还有 Swakane 片麻岩 (图 4.1,图 5.12, 图 5.19 和图 7.7),在组分、结构和时代上都与加利福尼亚州南部的 Orocopia–Pelona 片岩相似,它们 所含最年轻的碎屑锆石为约 73 Ma,在 5 Ma 内沉积物中变质作用发生的 P−T 条件为 9~12 kbar 和 640 ~740℃ (Matzel et al., 2004)。在 Swakane–Napeequa 接触带的一个片理化花岗闪长岩测定的年龄为 84±1 Ma (Hurlow, 1992)。

在喀斯喀特(Cascades)结晶核的内部和西侧,横跨直溪-弗雷泽河(Straight Creek-Fraser River)断层到圣胡安群岛(San Juan Islands)西侧的是一个叠瓦状海蚀柱(图7.7),从110 Ma之后形成的各地质单元来看,似乎由在96~90 Ma之前侵入的深成岩体组成(Brandon et al., 1988; Brown and Dragovich, 2003; Brown and Gehrels, 2007; Brown et al., 2007)。堆积体中的单元包括片岩和大量的混杂岩带,其中一些含蓝片岩相变质作用,170~160 Ma 的菲达尔戈(Fidalgo)蛇绿混杂岩



图 7.7 华盛顿州西北部地质简图

展示了圣胡安岛(San Juan)-喀斯喀特(Cascades)北西逆冲系统和喀斯喀特西部结晶核及周围地区的基础地质。 CB—Chilliwack 岩基; GHB—金 Horn 岩基; MSB—斯图尔特山(Stuart)岩基; SB—史诺柯米(Snoqualmie)岩基; 据 Brown and Dragovich (2003)修改,深成岩体年龄据 Miller 等(2009)

(Garver, 1988),其可能与东部 161 Ma 的英戈尔斯(Ingalls)蛇绿混杂岩有关(Miller, 1985; Miller et al., 2003; MacDonald et al., 2008),观景峰(Lookout Mountain)组(MacDonald, 2006)中所含的最年轻碎屑锆石为 160 Ma(MacDonald et al., 2003, 2008),高磷的二叠纪—三叠纪维德(Vedder) 杂岩(Armstrong et al., 1983)和含辉石片麻状黄紫菀(Yellow Aster)杂岩(Misch, 1966; Mattinson, 1972),其与育空-塔纳纳(Yukon-Tanana)和嘘飞(Shoo Fly)地体的岩石具有相似的历史(Brown and Gehrels, 2007)。仅西部的岩石与弗朗西斯科的特征和年龄相似(Brown, 1986),它们都是在110~80 Ma之间发生变形,含有年龄范围从160~144 Ma 的较老深变质岩块、130~120 Ma的蓝片岩岩片、164~163 Ma 的辉长岩-英云闪长岩、167±5 Ma 的蛇绿岩和年龄变化范围在114~90 Ma 的混杂岩母岩(Brown and Gehrels, 2007)。古地磁和生物证据均表明这些岩石在 75 Ma 时位于南部(Brown et al., 2007)。

## 7.17 爱达荷岩基

岩石或多或少地与喀斯喀特(Cascades)产出相关,向东南延伸跨过左旋的路易斯(Lewis)断层和克拉克(Clark)线性构造以及奥罗菲诺断层(Armstrong et al., 1977; Wallace et al., 1990; Mc-Clelland and Oldow, 2007)。中白垩世爱达荷岩基的亚特兰大(Atlanta)和比特鲁特(Bitterroot)岩瓣(图 2.5和图 4.1)侵入到各种岩石中,并沿萨蒙(Salmon)河缝合带东侧分布,包括中元古界的贝尔特(Belt)超群,一个向西倾的逆冲带,以及较年轻的爱达荷西部剪切带,该剪切带通常认为是大陆尺度的横向运动标志(McClelland et al., 2000)。沿岩基的西部边缘发生了各种变形作用,在某些地方,年龄在118±5~105±1.5 Ma的含绿帘石英云闪长岩到花岗质正片麻岩被92→90 Ma的含绿帘72

石英云闪长片岩所切割(Taubeneck, 1971; Hyndman, 1983; Manduca et al., 1993; Giorgis et al., 2008)。往北,奥罗菲诺断层的南侧(其可能为一个剥蚀双重构造)一个向西倾逆冲断层的构造窗中含新一中元古界岩石,并且出露一组94~86 Ma 的副片麻岩和94~73 Ma 的正片麻岩,两者明显都不跨过西部的缝合带和断层,少量薄层状超基性岩变形时间不晚于 68~61 Ma(Lund et al., 2008)。岩基其他地方为 98→80 Ma 之间的偏铝质英云闪长岩、黑云母花岗闪长岩、钾长石斑状花岗闪长岩和花岗岩(Kiilsgaard et al., 2001; Gaschnig et al., 2010)。亚特兰大(Atlanta)岩瓣的主体岩石为弱过铝质、年龄为 83~67 Ma 的深成侵入杂岩,深成杂岩由以黑云母花岗闪长岩为核心的白云母-黑云母花岗岩组成,所有都被淡色花岗岩脉切断(Kiilsgaard and Lewis, 1985; Lewis et al., 1987; Gaschnig et al., 2010),而比特鲁特(Bitterroot)岩瓣并不含在亚特兰大(Atlanta)岩瓣中发现的较老的偏铝质岩石,由年龄为 74~69 Ma 的较年轻的偏铝质岩体组成;然而,大部分的岩瓣是由年龄从 66~53 Ma 之间的过铝质岩体组成,在成分上比南部同类岩石更不均匀,并且伴生基性岩脉、岩席和小型侵入体,这些是亚特兰大(Atlanta)岩瓣所不具有的(Hyndman, 1984; Hyndman and Foster, 1988; Foster and Hyndman, 1990; Foster and Fanning, 1997; Gaschnig et al., 2010)。

博尔德(Boulder) 岩基,位于东侧并与埃尔克霍恩(Elkhorn)火山岩有关(图 4.1),是一个由小型基性-中基性深成岩体、大量花岗岩及晚期淡色岩体组成的复合岩基,年龄从 78→66 Ma,均为 钾质和钠质系列(Hamilton and Myers, 1967; Smedes et al., 1973; Tilling, 1973, 1974; Johnson et al., 2004)。在比尤特(Butte),这期岩浆作用形成了斑岩型铜矿化(Lund et al., 2002; Dilles et al., 2003),随后该区成为既有火山作用又有侵入作用的 53~43 Ma 查利斯(Challis) 岩浆活动区域(McIntyre et al., 1982; Johnson et al., 1988; Moye et al., 1988)。

#### 7.18 蓝山地体

这是三个不同地体的组合(图 2.5 和图 4.1),一个超覆层序位于俄勒冈东部-爱达荷西部的萨蒙 河(Salmon)缝合带的西侧,由于在地质和动物化石方面与其具有相似性,通常认为其与北部的岩 石有对应关系。这些地体有:位于东南端的老费里(Olds Ferry)地体,是最东南端的群,下部由中 晚三叠世的基性到中基性火山岩组成,上部由流纹英安质-流纹质熔岩和碎屑岩、火山碎屑质砂岩和 砾岩组成;贝克(Baker)地体,由发生剪切的二叠纪到早侏罗世硅质岩和硅质黏土岩组成,并含泥 盆纪—三叠纪的灰岩、蛇纹岩、基性和超基性岩石的岩块,局部发生蓝片岩相变质作用;瓦罗瓦 (Wallowa)地体,主要由二叠纪到早侏罗世的火山岩和沉积岩组成;Izee 超覆层序,是一个三叠纪一 侏罗纪沉积岩组成的厚层超覆层序,不整合地覆盖在其它三个地体上(Avé Lallemant, 1995; Dorsey and LaMaskin, 2008)。侏罗纪—早白垩世的深成岩体将这些地体缝合在一起。最初这些地体为南北 向,但在其缝合过程中转为现在的北东-南西向(Housen, 2007)。贝克(Baker)地体位于均被称为 代表弧的老费里和瓦洛瓦地体之间,两侧均受逆冲断层约束,倾向朝外形成了双向俯冲带(Avé Lallemant, 1995),据此 Dorsey 和 LaMaskin (2007)认为其属于摩鹿加群岛(Moluccan)的海洋型弧-弧碰撞。

一些学者认识到蓝山(Blue Mt.,位于美国俄勒冈州东北、华盛顿州东南的山脉)的火山弧与 Stikinia 和 Quesnellia 具有相似性,并认为其与瓦洛瓦和老费里这两个弧具有相关性,卡什溪的增生岩 石与贝克地体相似(Mortimer, 1986; Stanley and Senowbari-Daryan, 1986);但是,贝克地体内有两 个次级地体:格林霍恩(Greenhorn)次级地体是一个以蛇纹石为基质的混杂岩,其中只含麦克劳德 亲和的蜒类化石,而 Bourne 次级地体以细粒硅质黏土岩为主,含古地中海的和麦克劳德的蜒类化石 (Ferns and Brooks, 1995; Schwartz et al., 2011)。从 Izee 地体中获得的古纬度数据表明其在晚三叠世 时位于低纬度地区,并在中侏罗世时迁移到较高纬度地区(Pessagno, 2006)。

Ochoco 残余盆地覆盖在贝克(Baker)和 Izee 地体最西部的露头之上,由中至晚白垩世的厚层海相沉积组成,主要为泥岩、砂岩和砾岩(Dorsey and Legnegan, 2007)。从霍恩布鲁克盆地超覆层序中收集的碎屑锆石数据表明该盆地不会老于森诺曼期(Kochelek and Surpless, 2009),并很可能与霍

恩布鲁克盆地一样是同一盆地的一部分(Surpless et al., 2009)。从古地磁研究中获得的古纬度数据 表明盆地的岩石(即一般而言蓝山的岩石)在森诺曼期位于1200~1700 km 更南的地区(32°±7°) (Housen and Dorsey, 2005),但在始新世就已到其现在的位置,那时被克拉洛(Clarno)火山岩上超 (Grommé et al., 1986)。这些岩石的倾斜校正方向在本质上与霍恩布鲁克盆地的岩石是一样的,支持 这两个盆地有联系的结论(Housen and Dorsey, 2005)。

# 第8章 阿拉斯加地区

## 8.1 布鲁克斯山脉、阿拉斯加北坡和山弧

布鲁克斯(Brooks)山脉和阿拉斯加北坡的地质情况非常复杂,但在最通常情况下可以将其分为 北极阿拉斯加-北坡地体和 Angayucham 地体(图 2.5)。在晚侏罗世—早白垩世布鲁克(Brookian) 造山期期间,当北极阿拉斯加地体及其被动边缘的岩石部分地俯冲到 Angayucham 地体之下,两者都 发生了变形和变质作用(Moore et al., 1994)。北极阿拉斯加地体的地层分为三个主要的岩群:前密 西西比纪岩石、埃尔斯米尔和布鲁克(Lerand, 1973; Moore et al., 1994; Handschy, 1998)。前密西 西比纪序列由各种新元古界岩石组成,其中一些岩石的岩性与温德米尔超群非常相似(Moore and Bird, 2010),但与加拿大北部阿蒙森(Amundsen)盆地的岩石年龄不同(Macdonald et al., 2009), 上覆下古生界碎屑状的硅质沉积岩及火山岩。古生界和更老的岩石被泥盆纪深成岩体所侵入,北部在 被埃尔斯米尔序列的岩石覆盖之前发生了变形和变质作用,并主要呈面向南(现在的坐标)的被动 陆缘,基本上由密西西比纪非海相砾岩及其上覆密西西比纪—三叠纪碎屑质-碳酸盐沉积岩组成 (Brosgé et al., 1962; Martin, 1970; Brosgé and Tailleur, 1971; Mull, 1982; Hubbard et al., 1987; Mayfield et al., 1988; Moore et al., 1994)。富兰克林阶和埃尔斯米尔序列两者的岩石都在布鲁克 (Brookian)造山期间与它们的基底分离、褶皱并向北逆冲。

与科迪勒拉造山带其他地方相似,在中生代—古近纪—新近纪期间的被动边缘上发育一个山前盆地,称为科尔维尔湖(Colville)盆地,其岩石总体上都属于布鲁克序列,该序列所含的沉积岩与两个阶段的变形有关:早阶段变形与布鲁克造山运动有关,以含阿普弟期(Aptian)岩石、向北长距离逆冲岩片侵位为特征;晚阶段变形,以麦斯里希特期—森诺曼期(Maastrichtian-Cenomanian)为主,包括小位移和褶皱的逆冲断层(Mull, 1985; Moore et al., 1994)。

恩迪科特山(Endicott)和德龙山(DeLong)外来岩体(图 2.5)是结构上最低的外来岩体,并 在布鲁克造山运动时期就位(Moore et al., 1994)。虽然受到许多内部逆冲断层和褶皱的影响,但还 是可以看出外来岩体内部的地层与北坡非常相似,除了它们并不含有比泥盆系顶部更老的岩石,原因 是该部分可能被基底拆离构造截去顶部,没有明显的证据表明其与亚密西西比阶不整合接触(Moore et al., 1994; Handschy, 1998)。

布鲁克斯山脉的南半部主要由哈蒙德(Hammond)和科尔德富特(Coldfoot)次级地体组成(图 2.5),两者在构造上都呈叠瓦状,但不稳定地变形,可能为元古界到下古生界的碳酸盐岩、片岩和 千枚岩岩石组合,容易变质成绿片岩、蓝片岩和角闪岩,并被上元古代和泥盆纪的深成岩体所侵入 (Dillon et al., 1980, 1987; Karl et al., 1989; Till, 1989)。一个通常达到蓝片岩相的窄条带状元古代 千枚岩、板岩和砂岩连同多斑点的基性侵入体和混杂岩一起分布在布鲁克斯山脉的南部边缘, Moore 等(1994)将其统称为斯莱特溪(Slate Creek)次级地体。苏华德半岛(Seward Peninsula,美国阿拉 斯加州西部)内也有相似的岩石,但与其并无必然联系,被称为诺梅(Nome)杂岩(Till et al., 2010),其中的碎屑锆石 U-Pb 数据出现了新元古代和古生代两个峰值,这表明苏华德半岛的岩石相 对于北美来说属于外来体(Amato et al., 2009)。

在布鲁克斯山脉,处于构造上最高位置的外来体出现在 Angayuchan 地体内,露头位于斯莱特溪带的南部,并作为飞来峰覆盖在山脉的顶部 (Mull, 1982; Patton et al., 1994; Moore et al., 1994)。 一般将其岩石划分为两个结构单元:下部组合由绿岩、枕状玄武岩为主以及晚泥盆世到早侏罗世硅质 岩组成;上部组合由中侏罗世辉长岩和超基性岩组成(Zimmerman and Soustek, 1979; Pallister and Carlson, 1988; Patton and Box, 1989),并被 170 Ma 的斜长花岗岩和 163 Ma 的细晶岩脉所切割(Moore et al., 1994)。从两个不同时代的组合可以得出以下模型:在侏罗纪期间的 154 Ma 左右,下部的火山岩组合被逆冲到上部的蛇绿岩组合之下,这一事件发生在北极阿拉斯加之下的复合外来体在凡兰吟阶(Valanginian)期间就位之前(Boak et al., 1987; Mayfield et al., 1988)。

Tozitna 地体(图 2.5)位于南部并逆冲到红宝石地体之上,Silberling等(1992)认为及是一个单独的地体;然而,它与Angayucham 地体在完全相同的逆冲构造序列中具有同样年龄的岩石和相同的结构顺序,这使得一些学者(Patton et al., 1989, 1994)认为两者之前为同一地体。

同样,加拿大西部的内陆盆地,在侏罗纪和早白垩世的充填沉积层相对较薄并主要由浊积岩组成,Moore等(1994)称之为原科尔维尔盆地,而较年轻的科尔维尔盆地充填沉积层更厚、颗粒更粗(Bird and Molenaar, 1992)。科尔维尔(Colvile)是一个造山带前渊盆地,也是从布鲁克斯山脉内碰撞造山带向南的主要沉积场所(现在的坐标)。主要发育在另一个 Beaufortian 沉积序列之上,后者的物源来自美亚(Amerasian)盆地隆起端的巴罗(Barrow)拱形,该盆地在早白垩世时向北开口(Grantz and May, 1982; Grantz et al., 1990)。

在布鲁克斯山脉的中部和西部,高度变形的贝利亚斯阶到凡兰吟阶(Berriasian to Valenginian) 浊积岩局部伴随重力滑动沉积、砾岩透镜体和薄层雏蛤属(Buchia)贝壳灰岩层,被统称为 Okpikruak组,该组覆盖在恩迪科特(Endicott)和德龙山外来体之上(Bird and Molenaar, 1987, 1992),一般认为其代表了初始碰撞产生的一部分向东脱落的碎片。更北部分相关岩石的一个单元由 含燧石的卵石页岩、石英岩鹅卵石、铁矿石结核、磨砂石英颗粒组成,显然是在侏罗纪—豪特里维期 沉积下来的,面向南的地台外部在美亚盆地(Amerasian basin)张开期间隆升和剥蚀提供了沉积物源 (Macquaker et al., 1999)。Okpikruak组的岩石并不含与布鲁克斯山脉的外来体岩石相匹配的碎屑锆 石,所以其可能在逆冲作用发生之前就沉积覆盖在了边缘之上,或者如果其比较年轻,由碎片组成的 物源则在纵向上经历相当大的距离搬运过来(T. Moore, 2011,私人通讯)。这与科尔维尔盆地在巴 列姆阶—阿普弟阶(Barremian-Aptian)时期急剧向下倾斜的证据相符(Cole et al., 1997),并与逆冲 到利斯本山(Lisburne)西部相吻合(Moore et al., 2002)。不整合覆盖在Okpikruak组岩石和 Endicott 和 DeLong 山外来体风化岩石之上的是厚达 3400 m 的阿普弟阶—森诺曼阶砂岩、砾岩、泥岩和煤层, 统称为堡寨(Fortress)组(Mull, 1985; Siok, 1989; Crowder, 1989; Bird and Molenaar, 1992)。

沿 Angayuchan 地体南侧分布的是科伯克-科尤库克(Kobuk-Koyukuk)盆地的岩石(图 2.5), 为一个 U 型盆地,含有 5~8 km 厚的中到上白垩统陆源碎屑沉积岩(Nilsen, 1989),与西侧的苏华德 半岛、北侧的布鲁克斯山脉和东侧的红宝石地背斜(Patton and Box, 1989; Patton et al., 1994)三面 接壤,可能来自于变质程度相似的元古界和古生代大陆地体。占据盆地中心部分的是科尤库克地体, 其下部岩层由年龄不详的玄武质-安山质熔岩和超基性岩石组成,并被英云闪长岩-奥长花岗岩所侵 入,所有都被上部岩层不整合覆盖,上部岩层由以贝利亚斯期-凡兰吟期为主——局部为年轻的阿普 弟期——浅海到深海相火山碎屑岩、玄武碎屑岩、枕状斑状斜长玄武岩和英安质熔岩组成,成分范围 从拉斑玄武岩到橄榄安粗岩(Patton and Box, 1989; Patton et al., 1994, 2009)。

往东,伏于 Angayucham-Tozitna 地体之下的是红宝石地体(Patton et al., 1989)。该地体(图 2.5)是一个北东走向的线状带,由前寒武纪—古生代泥质片岩、石英岩、变质基性岩、大理岩和正片麻岩组成,岩石发生了各种程度不同的变质作用,如绿片岩相、蓝片岩相和角闪岩相等,并被泥盆纪深成岩体所切割,共同组成了一个中白垩世之前北东走向的隆升核心(Patton et al., 1994)。基于相似的岩石类型、变质程度、产出泥盆纪深成岩体以及构造背景,大多数学者倾向于认为红宝石和北极阿拉斯加地体曾经构成了一个连续的内陆带(例如, Carey, 1958; Tailleur, 1980; Box, 1984; Grantz et al., 1991; Roeske et al., 1995; Johnston, 2001)。

中一晚白垩世深成岩体产出在两个带 (图 2.5): 一条为向西延伸的带,出现在育空-科伯克-科 尤库克盆地内和苏华德半岛;另一条出现在红宝石地体,延伸长度超过地体范围 (Patton and Box, 1989; Miller, 1989; Arth et al., 1989a; Patton et al., 1987, 2009; Till et al., 2010)。Arth 等 (1989b) 认为该深成岩体是老大陆地幔和老大陆地壳混合重熔的产物,岩体走向变化很大,成分范围从钙碱性到超钾碱性岩。红宝石带中的岩石均为钙碱性,沿走向变化也很大,其中西南部的物源可能来源于老的大陆地壳,而东北部很可能来源于洋壳岩石的重熔,年轻的地壳受到老地壳的混染 (Arth et al., 1989a)。该深成岩体明显切割了红宝石-Angayucham 逆冲推覆体 (Patton et al., 2009), 其年龄为112~99 Ma (Miller, 1989),因此相对于布鲁克造山运动明显属于后碰撞岩石。

位于红宝石地体南部的是偏远的 Farewell 地体 (图 2.5), 在一些划分方案中该地体是由尼克松 福克(Nixon Fork)、迪林格(Dillinger)、Minchumina 和 Mystic 等次级地体组成的复合地体 (Silberling et al., 1992),在另一些方案中是由两个截然不同的沉积序列:白山(White Mountain)和 Mystic 组成,以及其他 (Decher et al., 1994)。尼克松福克次级地体 (图 2.5) 由一个向南东倾的前 寒武变质基底构成,岩性由绿片岩、绿岩和小型长英质深成岩石等组成(Patton et al., 1980),岩石 年龄范围从至少 1265 ± 50 Ma→约 850 Ma 不等 (Dillon et al., 1985; McClelland et al., 1999), 全部被 约5km 厚的奥陶系—上泥盆统碳酸盐岩沉积所覆盖(Patton et al., 1994)。迪林格次级地体(图 2.5) 由近3km 厚的被复杂地错断和褶皱的古生代浊积岩/少量绿岩、黑色页岩和燧石,薄层状灰岩、砂岩 和角砾岩组成,被认为是代表了一个地台的盆地相和尼克松福克次级地体的斜坡相岩石 (Churkin and Carter, 1996: Decker et al., 1994)。Mystic 次级地体的密西西比纪—二叠纪到早白垩世陆源碎屑岩岩 石或层序,不整合地覆盖在迪林格和尼克松福克次级地体的较老岩石之上(Decker et al., 1994; Bundtzen et al., 1997)。Minchumina 次级地体(图 2.5)由晚元古代的深水碳酸盐岩、燧石、硅质黏 土岩和石英岩和可能代表尼克松福克-迪林格层序离岸沉积的早古生代沉积岩组成(Patton et al., 1994)。尼克松福克的古生代岩石现在看来更有可能是地台相岩石,与迪林格层序的盆地相岩石相互 贯穿;因此 Decker 等(1994)将其归为白山序列。晚二叠世的40 Ar/39 Ar 坪年龄表明存在一个晚古生 代变形事件, Bradley 等(2006)将其命名为布朗斯福克造山运动(Browns Fork orogeny)。

阿拉斯加地体明显弯曲环绕在科伯克-科尤库克盆地周围(图 2.5),自从 Carey (1955, 1958) 将其划分到马蹄形山系的一部分时就引起了地质学家们的关注。Johnston (2001, 2008)采用马蹄形 山系的观点将带状大陆解释为向北的移动,而 Box (1985)则引用不规则边缘,或海角模型来解释地 体的这种弯曲现象。

## 8.2 兰格利亚复合地体

兰格利亚是一个位于加拿大最西部和阿拉斯加南部的已被肢解的复合超级地体(图 2.5),其跨 越加拿大和阿拉斯加两个分区。通常认为其包含三个地体:亚历山大、兰格尔山脉(Wrangell)和半 岛(Peninsular)(Jones et al., 1977; Nokleberg et al., 1994, 2000)。兰格尔地体以 230~225 Ma 的溢 流玄武岩为主,与温哥华和昆沙罗特湾(Queen Charlotte)岛的 Karmutsen 组相似,其中完好地保存 了约 6 km 的高 Ti 玄武岩和苦橄岩、枕状玄武岩、玻质碎屑岩和顶部的陆上熔岩,在阿拉斯加和育空 地区的尼古拉(Nikolai)组,保存有 1~3.5 km 的陆上高 Ti 玄武岩为主与高 Ti 枕状玄武岩(Lassiter et al., 1995; Greene et al., 2008, 2009, 2010)。在温哥华岛,玄武岩位于叠瓦状逆冲推覆体堆积最 上部岩席的底部(Monger and Journeay, 1994),并被地台相灰岩覆盖,与阿拉斯加中南部相关的岩石 类似(Carlisle and Susuki, 1974; Jones et al., 1977; Yorath et al., 1999)。地台相灰岩上覆三叠系顶 部和侏罗系页岩、灰岩和硅质黏土岩夹层,也被其间的侏罗系博南扎弧火山岩覆盖(Caruthers and Stanley, 2008; Nixon et al., 2006)。

温哥华 202~165 Ma 的博南扎弧和夏洛特皇后岛(Queen Charlotte islands,位于加拿大西部海岸外)由厚达 2500 m 的熔岩互层、凝灰岩和角砾岩互层组成,其中角砾岩成分从玄武岩 到流纹岩(DeBari et al., 1999)。西海岸结晶杂岩和片状岛侵入岩两组侵入岩体通常被认为 代表了从 190→169 Ma 的岛弧深成岩的根部(Isachsen, 1987; DeBari et al., 1999; Canil et al., 2010)。

被认为是博南扎弧一部分的岩石,出露在不列颠哥伦比亚省南部的大陆和附近岛屿,它们被统称为鲍恩岛群(Bowen Island Group)(Friedman et al., 1990)。该群地层发育紧密褶皱和面理化,形成于一个被定年为154 Ma的深成岩体侵位之前(Friedman and Armstrong, 1995)。实际上,这种变形范围可能更加广泛(McClelland and Gehrels, 1990; Monger, 1991)。

亚历山大地体 (Berg et al., 1972) 下伏于阿拉斯加东南部大部分地区、育空地区和不列颠哥伦 比亚省的部分地区 (图 2.5), 主要由一系列混合的和各种变形的岩石组成, 覆盖在下古生界弧地体 之上, 其本身又被晚古生代和三叠纪的碳酸盐岩和碎屑岩夹层与数量不定的玄武质熔岩和角砾岩所覆 盖 (Jones et al., 1972; Gehrels and Saleeby, 1987; Gehrels, 1990; Gehrels et al., 1996)。根据不同的 和有特色的巨形动物群和碎屑锆石数据, 该地体的岩石明显是来自波罗的海或者西伯利亚 (Bazard et al., 1995; Gehrels et al., 1996; Soja and Antoshkina, 1997; Blodgett et al., 2002, 2010; Pedder, 2006)。一个测定为 302 Ma 宾夕法尼亚纪的深成岩体侵入到兰格尔山脉和亚历山大地体中, 为这两个 地体的拼合提供了最小的年龄约束 (Gardner et al., 1988)。

虽然博南扎弧代表了加拿大南部的地体,但半岛地体主要位于阿拉斯加(Nokleberg et al.,2000)。该地体在阿拉斯加中南部绵延1000多千米,以塔尔基特纳-博南扎弧为主,被称为边界山脉(Border Ranges)超基性和基性杂岩,位于弧岩石和边界山脉断层之间,(Smart et al.,1996; Pavlis and Roeske, 2007),把半岛地体从楚加奇(Chugach)增生杂岩中分离出来,认为其代表了塔尔基特纳弧的基底(Burns, 1985; DeBari and Coleman, 1989; Kusky et al., 2007; Farris, 2009)。

塔尔基特纳弧在斜截面上很好地出露,展示了从地幔构造岩到上覆沉积岩的一个被肢解弧的剖面 (Greene et al., 2006; Hacker et al., 2008)。在楚加奇和塔尔基特纳山脉,保留有约7km的陆上和海 底熔岩、角砾岩、凝灰岩、熔凝灰岩和火山碎屑沉积(Clift et al., 2005),并被拉长的石英闪长岩和 奥长岩深成岩体所切割(Rioux et al., 2007)。塔尔基特纳弧火山作用的主要阶段发生在 202~175 Ma 之间,在北部地区也有几个年龄在 190~153 Ma 之间的深成岩体侵入(Rioux et al., 2007; Hacker et al., 2011)。再往西到科迪亚克岛(Kodiak,在阿拉斯加南部)和阿拉斯加半岛是一个分布广泛的侏 罗纪岩基,以石英闪长岩和英云闪长岩为主,最老的年龄 213 Ma 出现在科迪亚克岛,较年轻一组出 现在半岛北部,年龄范围从 184~164 Ma(Rioux et al., 2010)。因此,无论是东部还是西部的岩浆作 用都是向北逐渐变年轻。岛弧内岩石的同位素也和碎屑锆石和继承锆石总数一样系统变化,岛弧内沿 走向往东部岩石更年轻、在西南部岩石演化更强,表明基底形成的地质背景不同(Clift et al., 2005; Greene et al., 2006; Rioux et al., 2007; Amato et al., 2007)。

在阿拉斯加山脉东部有几个线性带状深成岩体:晚侏罗世—早白垩世,钙碱性的奇提纳 (Chitina) 岩体侵入到超级地体南部的岩石中;奇萨纳弧位于兰格利亚更北一点,由钙碱性深成岩和 安山质熔岩组成;白垩纪末—古近纪的克卢恩(Kluane)深成岩岩套,侵入到兰格利亚和育空-塔纳 纳地体中(Trop and Ridgway, 2007)。Hildebrand(2009)认为晚白垩世—古近纪的深成岩体是俯冲 板片断离岩浆作用形成的广泛分布的岩浆带一部分。

晚侏罗世期间,沿兰格利亚南缘分布的岩石逐渐向北逆冲,沿着向南倾的逆冲断层同时产生的牛 津期到蒂托期(Oxfordian to Tithonian)碎屑在 Nutzotzin-兰格尔山脉盆地沉积,不整合覆盖在地台相 碳酸盐岩之上,并且向上粒度逐渐变粗,从海相泥岩、砂岩到砾岩(Trop et al., 2002; Manuszak et al., 2007),该盆地通常被认为代表了塔尔基特纳弧的一个弧后盆地(Trop and Ridgway, 2007; Manuszak et al., 2007)。另一个牛津阶到蒂托阶盆地相充填序列,由巨型扇三角洲杂岩组成,具有从 逆断层陡坡脱落的粗巨砾的碎片,位于塔尔基特纳弧南侧,它被称为纳克内克(Naknek)组(Trop et al., 2005)。古地磁和动物化石群数据表明,自三叠纪以来,兰格利亚的超级地体相对于北美克拉 通向北至少移动了 30°,并且在约 52 Ma 时到达现在的纬度(Jones et al., 1977; Hillhouse, 1977; Hillhouse et al., 1985; Irving et al., 1996; Pedder, 2006)。

## 8.3 阿拉斯加中部的造山盆地

沿兰格利亚的北部和东部边界分布着两个复理石盆地,分别为晚侏罗世—白垩纪的卡希尔特纳 (Kahiltna) 盆地(图 2.5 和图 8.1)以及前文与加拿大西部海岸深成杂岩—起讨论的具有相似年龄的 格拉维纳岛(Gravina)-Nutzotin+Dezeadesh(?) 盆地(图 2.5)。在阿拉斯加山脉南部,在 Valenginian—森诺曼阶(Cenomanian)卡希尔特纳盆地内发育一个近 6 km 长的海相碎屑岩带(Kalbas et al., 2007),盆地的岩石沿走向出露近 800 km,位于兰格利亚南部和育空-塔纳纳(Yukon-Tanana)和 Farewell 地体北部之间,但该地区被德纳里峰(Denali,美国阿拉斯加州中南部山脉,即麦金利山) 断层一分为二(图 2.5)。详细的沉积学研究,如相分析、碎屑锆石测年和碎屑统计结果表明,在阿 普第期—阿尔必期,盆地内轴向向西倾的海底扇杂岩之上至少充填了 4 km 厚的沉积物,物源主要来 自盆地两侧(Kalbas et al., 2007)。由于盆地关闭的同时育空-塔纳纳上盘板块的岩石明显后退,导 致在 110~90 Ma 之间盆地内沉降停止、岩石变形并且向南逆冲(Ridgway et al., 2002)。当岩石局部 变质为含蓝晶石组合且在 74 Ma 被同构造英云闪长岩席所侵入时,显然碰撞带在坎帕期—麦斯里希特 期被重新激活(Davidson et al., 1992; Ridgway et al., 2002),这次变形与坎帕期—麦斯里希特期坎特 韦尔(Cantwell) 盆地的演化一致(图 2.5 和图 8.1),在卡希尔特纳岩石 84~65 Ma向北逆冲和褶皱 期间,位于其北部的一个逆冲顶盆地正好形成,在其上部还包含一个 60~54 Ma 玄武岩-流纹岩序列 (Cole et al., 1999; Ridgway et al., 2002; Trop and Ridgway, 2007)。



图 8.1 阿拉斯加中南部地质简图 (据 Trop and Ridgway (2007)修改) 该图表示了本文所讨论的一些地质特征的分布

## 8.4 楚加奇增生杂岩

另一组岩石,一般被认为代表了沿阿拉斯加南部分布的中生代—古近纪—新近纪增生杂岩露头,这些岩石统称为楚加奇增生杂岩,或者南部边缘复合地体(图 2.5),该杂岩被边界山脉(Border Ranges)断层从兰格利亚的弧地体中分隔出来(Smart et al., 1996; Pavlis and Roeske, 2007),像弗朗西斯科—样,向远离断层的方向逐渐年轻,在外侧和临海地区变质程度逐渐降低(Plafker et al., 1994; Roeske et al., 2003; Pavlis and Roeske, 2007)。不连续的、以 200±10 Ma 断层为界的蓝片岩岩片出露在临近边界山脉断层的北部(Roeske et al., 1989)。较年轻的混杂岩露头出现在南端根据其出露位置,被称为乌亚克(Uyak)或者麦克汉(McHugh),由内部含复杂条带的泥岩和硬砂岩组成,含蛇绿岩岩石的岩块和碎片,如角岩、变质玄武岩、辉长岩和不太常见的超基性岩块(Kusky and Bradley, 1999)。最近的碎屑锆石研究揭示,麦克休(McHugh)杂岩在两个不同的时间沉积:一个不老于牛津阶(157~146 Ma),另一个最大年龄介于91~84 Ma 之间(Amato and Pavlis, 2010)。更远的外侧位于楚加奇复理石中,为南缘复合地体的最大部分,该复理石的沉积发生在 85 Ma 之后,但主要在麦斯里希特阶期间的68~67 Ma(Sample and Reid, 2003; Kochelek and Amato, 2010; Kochelek et al., 2011),因此,在增生楔中可能存在一个长达 15 Ma 的沉积间断。

根据古地磁、同位素和物源区资料,复理石可能来自海岸深成杂岩(Farmer et al., 1993; Sample and Reid, 2003; Kochelek et al., 2011),现在大部分位于不列颠哥伦比亚省,但在沉积时则位于更远的南方。这一发现获得了 Roeske 等(2003)的支持,他对比了兰格尔山脉的西海岸侵入岩套中一个颇具特色的 170 Ma 侵入岩套,该岩套位于温哥华岛西部,结果表明岩套可能早在 85 Ma 就已开始向 北迁移,但有点连续地从 70~51 Ma。

威廉王子(Prince William)地体位于更远的外海,为古新世—始新世复理石带(图 2.5),层间 夹玄武质熔岩,称为幽灵石(Ghost Rocks)组,位于现今 55°N 的科迪亚克(Kodiak)岛(Moore et al., 1983)。前人的工作(Plumley et al., 1983)和最近完成的研究(Housen et al., 2008; Roeske et al., 2009)都表明该熔岩的古纬度从 48°迁移到 41°。一个侵入岩套,称为盛纳克-巴拉诺夫(Sanak-Baranof)深成岩带(Hudson et al., 1979)侵入到 61 Ma 的增生岩西部的岩石中,直到 51 Ma 时才停止了向东运移,被广泛认为是洋脊俯冲的产物(Bradley et al., 2003b)。

## 第9章 相关问题讨论

自 Monger 和 Ross(1971)在加拿大科迪勒拉山脉内发现了两种不同的蜓类化石种群并提出某些 地块可能存在远距离迁移的观点后,众多学者都尝试揭示北美科迪勒拉的构造拼合,其中最具影响力 的可能是 Coney 等 (1980), 他提出科迪勒拉地体存在巨大差异, 有一段增生历史。美国西部特别重 要的文章为 Burchflel 和 Davis (1972, 1975; Davis et al., 1978) 早期所著。另一个早期的工作是由 Monger 等(1982)完成的,他们尝试将加拿大科迪勒拉中不同地体拼合成更大的超级地体,这项工 作在那个年代是相当复杂的,他们提出,两个先前已经拼合起来的超级地体在不同时期与北美陆缘发 生了对接拼合。在他们的模式图中:东斯莱德山、Quesnel、卡什溪以及 Stikine 地体被归入山间的 (Intermontane) 超级地体内, 该超级地体在侏罗纪与北美陆缘发生了对接; 亚历山大、兰格尔山脉和 Gravina-Nuztotin 地体被归并入海岛的(Insular)超级地体,它们在白垩纪期间到达。其他较小地体, 如桥河 (Bridge River) 地体, 捕获山间的超级地体与海岛的超级地体之间的或楚加奇地体外侧的地 体,可能已分别抵达。他们还认为,奥米尼卡带及海岸(Coast)深成杂岩带为两个独立的深成岩带, 那里有很多深成岩体侵位于先前已经拼合的地体中。Oldow 等(1989)、Saleeby (1983)、Saleeby 和 Busby-Spera (1992)及 Coney 和 Evenchick (1994) 作了很精辟、适时的综述, 但他们采用了基于 "科迪勒拉模型"作为讨论的出发点,因此与本文提出的分析有很大差别。本文与先前成果最根本的 区别在于:一般而言,老文献中提出的增生模型认为外来地体是逐次拼接到北美大陆之上,而我认 为,多数地体在近海组合为一个巨大的带状大陆块,之后再与北美大陆碰撞,其碰撞始于约 124 Ma 的塞维尔造山运动 (Sevier orogeny), 并最终在约 80~75 Ma 的拉勒米造山运动 (Laramide orogeny) 期间结束。由 Moores (1970) 最早提出的侏罗纪—白垩纪模型非常有远见, 他提出: 科迪勒拉是由 向西俯冲带之上的陆-弧碰撞作用形成的,这与 Mattauer 等(1983)、Templeman-Kluitt(1979)及 Chamberlain (Chamberlain and Lambert, 1985; Lambert and Chamberlain, 1988) 的更局部模型一样。本 文的读者会发现 Stephen Johnston (2001, 2008) 独立提出的观点与我的相近,是富有卓见并值得仔细 研读的。

下文是依据前面章节的资料对科迪勒拉拼合过程中的主要事件总结。然而,尽管该总结来源于许 多确凿的信息,但仍有很多不确定因素,因此本文也绝不能被认为是完全正确。但愿它能作为指南, 使将来重点问题和地区的研究更加集中。

## 9.1 大盆地和加拿大分区古生代事件

在早密西西比世期间,下古生界燧石-页岩序列和在罗伯茨山(Roberts Mt.)外来体中(图9.1) 零星分布的碱性玄武岩经历了复杂的断层错动和褶皱,由于它们在安特勒造山运动期间,就已位于安 特勒地台西缘之上(Roberts et al., 1958; Nilsen and Stewart, 1980; Miller et al., 1992; Poole et al., 1992)。外来体的东边形成了一个充填沉积粗粒陆源碎屑的前陆盆地(Gehrels and Dickinson, 2000), 但是它们没有到达北美地台,距离北美地台不到100 km。某些学者(Finney and Perry, 1991; Miller et al., 1992; Poole et al., 1992)认为罗伯茨山外来体的岩石仅仅是北美大陆边缘大陆架外的细粒浅 海沉积;其他学者,如最著名的Ketner(1968, 1977, 1991)认为:大量的粗粒古生代沉积物、低成 熟度要求其西侧存在一个物源区,因为在这段时期的大部分时间里,北美大陆周边都是碳酸盐岩海 滩,并被陆表海覆盖。Wright和Wyld(2006)非常精辟地总结了沉积学方面的各种观点以及最新的 碎屑锆石研究成果,认为罗伯茨山岩石是亲冈瓦纳古陆的更大地体家族的一部分,这个家族包括嘘飞 杂岩、怀里卡-崔尼蒂杂岩和亚历山大地体。泥盆纪期间罗伯茨山从 Ligerian 洋迁移至古泛大洋 (Panthalassic)(也见: Grove et al., 2008b)。无论罗伯茨山外来体最初的物源是什么,弧火山作用及 深成侵入作用的缺失表明,其最西边部分(包括推测的晚古生带弧)在它增生到安特勒地台上之后、 宝山外来体抵达其西部之前就被移除了(图 2.5),一种可能是如今出现在加拿大库特尼地体南部的 岛弧岩石,该岩石与罗伯茨山外来体的岩石具有很多相似性,并且在老鹰湾(Eagle Bay)岩石组合 内包含一套泥盆纪—密西西比纪弧岩石(Schiarizza 和 Preto, 1987; Smith 和 Gehrels, 1992b; Paradis et al., 2006);另一种可能是育空-塔纳纳(Yukon-Tanana)地体,它同样含有泥盆纪—密西西比纪 弧岩石(Colpron et al., 2006)。Burchfiel 和 Royden(1991)认为,如果安特勒造山运动表现为亚平宁式(Apennine-type)的碰撞,那么缺失火山弧就不是问题,在亚平宁式碰撞中的岛弧被强烈肢解并 主动下沉,但是考虑到沿科迪勒拉边缘地体的移动性,有必要对科迪勒拉造山带内一些从别处迁移过 来的岛弧进行研究。

宝山外来体的岩石是复杂变形的海相沉积岩和火山岩,年龄范围在泥盆纪末至二叠纪之间(Silberling and Roberts, 1962; Miller et al., 1992),在大多数模型中,它们在晚二叠世—早三叠世索诺玛(Sonoman)造山运动期间,沿向西倾的逆冲断层就位于罗伯茨山外来体之上(图9.1)(Burchfiel et al., 1992; Miller et al., 1992),这里有两个突出问题:①该区没有二叠纪弧岩石;②正如 Wright 和 Wyld (2006)所指出的,没有任何证据表明东部存在前渊。现在位于加拿大和阿拉斯加分区的育空-塔纳纳地体中确实存在密西西比纪和二叠纪的弧岩石(Colpron et al., 2006),该地体有可能代表着宝山外来体的弧部分。所有碰撞模型都难以解释前渊的缺失,于是 Wright 和 Wyld (2006)提出了一个宝山盆地模型。

总之,联合起来的宝山-罗伯茨山外来体的岩石已结合在一起,并被面朝西的三叠纪碳酸盐台地 覆盖(Wyld et al., 2001)。

如本书前面章节所述,兰格利亚和亚历山大地体被年龄为 309±5 Ma 的深成岩体连接到了一起 (Gardner et al., 1988),因此,这些地体在晚古生代和中生代期间就在彼此靠得很近的地方演化。

育空-塔纳纳(Yukon-Tanana)和斯莱德山(Slide Mountain)地体,以及卡斯尔地台-塞尔温盆 地地块的拼合是加拿大科迪勒拉分区的重要地质事件之一,由Templeman-Kluitt(1979)提出,他指 出北美西部边缘向西俯冲到含弧的育空-塔纳纳地体之下,具有斯莱德山在两者之间捕获的大洋岩石 及增生楔岩石。育空-塔纳纳地体由泥盆纪变质岩及多期变形的基底岩石组成,被晚泥盆世—早密西 西比世深成岩体所切割,并被密西西比纪—二叠纪岛弧岩石所覆盖(Mortensen and Jilson, 1985; Colpron et al., 2006)。在260~253 Ma(乐平世,晚二叠世)期间安科尼(Inconnu)逆冲断层一次失 败的俯冲尝试中,卡斯尔地台的西部边缘被拉拽到育空-塔纳纳地体之下(图 9.1)(Murphy et al., 2006; Beranek and Mortensen, 2011),大洋性质的斯莱德山地体的岩石被错开,并呈构造接触关系位 于两个地体之间(Nelson, 1993)。

斯莱德山地体的岩石含有巨型蜓类化石,这些化石已知仅来源于少数几个地区,例如:华盛顿州 的水壶瀑布(Kettle Falls)、克拉马斯(Kiamaths)东部、墨西哥的索诺拉省(Sonora)、以及德克萨 斯州西部冒地槽内的一个原地沉积区域,所有这些都表明斯莱德山地体的岩石现在所处的位置较其起 源的温暖海水区要更北得多(Carter et al., 1992),古地磁数据也表明其至少向北移运了2000 km,支 持化石资料(Richards et al., 1993)。

塞尔温盆地的新元古界-古生代岩石通常被认为代表了北美被动陆缘的细粒滨外陆棚沉积,随着 卡斯尔地台的演化,北美被动陆缘上有一个晚志留世—中泥盆世抬升的边缘(Gabrielse et al., 1973; Cecile, 1982; Gordey and Anderson, 1993),但它们与罗伯茨山外来体的岩石具有很多共同之处,包 括:都含有燧石细粒沉积岩及在整个剖面零星出露碱性玄武岩、广泛分布重晶石层和局部出现沉积-喷气矿床(Goodfellow et al., 1995)。盆地的岩石是外来体,在晚侏罗世后、104 Ma之前作为一个巨 型复合外来体,沿着道森(Dawson)及破脑壳(Broken Skull)断层运移到北美陆缘,并与其下盘岩 石完全不同(Mair et al., 2006; Gordey and Anderson, 1993)。往南,盆地的岩石被褶皱成向南倒转 82



科迪勒拉山系碰撞前、碰撞过程及后碰撞(特别是板片断离)岩浆作用年龄图解;根据现有资料总结, 没有地体连续增生到北美大陆西部的证据

的推覆体,且被泥盆纪碳酸岩侵入,这个事件在北美地台以东地区前所未闻 (McLeish et al., 2010; McLeish and Johnston, 2011)。

## 9.2 晚三叠世到中侏罗世弧和碰撞

科迪勒拉地区普遍发育晚三叠世至中侏罗世弧地体,有些发育于洋壳之上、有些发育在古老的硅 铝质基底上、还有些兼而有之。这些地体包括塔尔基特纳、库特尼、Stikinian、Quesnellian、黑石, 几个克拉马斯弧和内华达山脉弧,以及鲜为人知的索诺兰沙漠弧。它们每一部分都是独特的,但有时 也会有很多相似之处,表明先前它们曾经相连在一起。在露头很差或研究程度低的地区,最年轻的岩 浆岩年龄大致与增生作用的时间相当。

## 9.2.1 塔尔基特纳弧

在阿拉斯加中南部和西南部,塔尔基特纳(Talkeetna)弧出露在兰格利亚(Wrangellian)超级地体的半岛(Peninsular)地体内(图2.5),沿走向延伸超过1000km。塔尔基特纳弧岩浆作用的主要阶段出现在202~175Ma之间,该区北部深成岩侵入发生在190~153Ma之间(Rioux et al., 2007)。在西部的科迪亚克(Kodiak)岛及阿拉斯加半岛,科迪亚克岛的侏罗纪岩基年龄为213Ma,而在阿拉斯加半岛为184~164Ma(Rioux et al., 2010)。根据同位素数据及捕获锆石资料,弧的基底沿走向及倾向是不同的(Clift et al., 2005; Greene et al., 2006; Rioux et al., 2007; Amato et al., 2007)。在兰格利亚超级地体内,弧的北侧是与牛津阶-蒂托阶Nutzotzin-兰格尔山(Wrangell)盆地一同发育的向南倾逆断层,盆地不整合于碳酸盐台地之上,向上粒度逐渐变粗,由海相泥岩及砂岩转变为砾岩(Trop et al., 2002; Manuszak et al., 2007),通常认为这个盆地代表了塔尔基特纳弧的弧后盆地(Clift et al., 2005; Trop and Ridgway, 2007; Manuszak et al., 2007)。但考虑到弧已经被逆冲推覆到碳酸盐台地之上,我认为,与之相反,它

是一个典型的前渊盆地,随后成为约 170 Ma 发生的塔尔基特纳弧与兰格利亚之间碰撞相关的前陆褶皱-逆冲断层带。兰格利亚地壳之上弧的就位可以解释下列现象:缺失切割兰格利亚的老于 170 Ma 的深成 岩体;岛弧构造强烈减薄,正如由地质温压计及 Ar 测年所推断的那样,在 10 Ma 的碰撞时间内,从原 来的 25~28 km 减薄至 7 km (Hacker et al., 2008; Rioux et al., 2007);弧岩浆作用年龄向北逐渐变年轻; 牛津阶-蒂托阶 (Tithonian)发育的纳克内克 (Naknek)盆地沿位于弧南边的向北倾逆断层形成悬崖, 盆地被粗碎屑岩和岩屑充填 (Trop et al., 2005),而该区后碰撞深成岩体 (<170 Ma,最年轻的为 153 Ma) (Rioux et al., 2007)可能代表着板片断离岩浆作用。

#### 9.2.2 博南扎弧

与阿拉斯加塔尔基特纳弧相比一般认为它们是相关的,不列颠哥伦比亚省南部博南扎弧(Bonanza Arc)在约202~165 Ma之间发育在兰格利亚地壳之上,由2500m厚的熔岩夹层、凝灰岩及角砾岩组成,成分范围由玄武岩至流纹岩,并被190~169 Ma的深成岩体切割(Isachsen,1987; DeBari et al.,1999; Canil et al.,2010)。该深成岩成分范围从主要是辉长岩-闪长岩到英云闪长岩和花岗闪长岩,它们明显侵入到兰格利亚之下的弧岩石中(Nixonet et al.,2011a,2011b,2011c,2011d)。虽然有这些关系,目前还不清楚为何弧岩浆作用在约165 Ma终止。

## 9.2.3 Quesnellia-库特尼-贝尔特-珀塞尔-温德米尔

这三个地体在侏罗纪时期由于相互碰撞而拼合在一起,并且其岩浆作用和增生作用在时间上表现 出非常明显的向东迁移特征 (图 9.1)。三个地体中位于最西部的 Quesnellia 地体含有大量晚三叠世— 早侏罗世火山岩,最年轻的可以到托尔期(Tipper, 1984),它也有大量 212~204 Ma的深成岩体 (Armstrong, 1988)。在北部,卡斯尔地台岩石在 186 Ma 时被向西倾的逆冲断层拖曳至 Quesnellia 地 体的东缘之下 (Nixon et al., 1993); 而在南部, 库特尼地体的西缘于 187~185 Ma 被拖曳至 Quesnellia 地体之下,形成一个向东的褶皱-逆冲断层带 (Murphy et al., 1995; Colpron et al., 1996, 1998)。早 期向西南的构造,如 Scrip 推覆体,被约 173~168 Ma 的向北东的褶皱和逆冲断层所叠加,并在 173 Ma 被同构造的 Kuskanax 岩基侵入(图 7.1)(Parrish and Wheeler, 1983)。第二阶段的变形与贝尔特 - 珀塞尔- 温德米尔 (BPW) 地块向西俯冲到库特尼地体之下的事件相符; 在图 7.1 中可以容易地看 到深成岩岩石年龄的变化,其中 Ouesnellia 含有大量早侏罗世深成岩体,而库特尼和 PBW 则没有: Ouesnellia 和库特尼均有大量的晚侏罗世深成岩体, 而 PBW 只有少数该年龄的小岩株, 和一个更年轻 的 115~90 Ma 深成岩体,称为贝永(Bayonne,美国新泽西州东北部海港) 岩套(Logan, 2002)。-个更年轻的岩套(白垩纪末—古近纪深成岩体)切割了上述这三个地体,它被 Hildebrand(2009)认 为是在拉勒米造山运动期间与北美板块俯冲失败有关的俯冲板片断离岩浆作用。年龄为157~153 Ma 的含白云母和石榴子石侵入岩以及 159 Ma 纳尔逊 (Nelson) 岩基 (图 7.1) 的黑云母花岗岩显然是 后碰撞的(Armstrong, 1988; Sevigny and Parrish, 1993)。两者很可能代表了与库特尼-PBW碰撞有 关的板片断离岩浆作用,而年龄变化在167~162 Ma范围的不同程度变形的深成岩体 (Ghosh, 1995),则可能是与库特尼板块俯冲失败(断离)有关的俯冲板片断离体。由于两次碰撞之间的时间 较短,要区分出哪些侵入体与俯冲有关、哪些与板片断离有关是非常困难的。库特尼地体双向汇聚的 扇形构造特征 (Wheeler, 1963; Brown and Tippett, 1978; Price, 1986; Brown and Lane, 1988; Colpron et al., 1996, 1998) 可能与 Quesnellia 作为狭窄的地体楔入有关, 它发生在第二次碰撞期间, 在 Quesnellia 和 BPW 地块之间被捕获。

有趣的是库特尼-BPW的碰撞和大盆地内陆带的变形年龄相似,大盆地的变形、变质和大型伏卧 推覆体的形成发生在165~160 Ma (Snoke and Miller, 1988; Miller and Hoisch, 1995; Zamudio and Atkinson, 1995; Camilleri et al., 1997; McGrew et al., 2000),被 Thorman (2011)命名为艾尔克造山运 动 (Elko orogeny)。紧靠内陆带东部的岩石为与温德米尔相似的厚层新元古界碎屑岩序列,这似乎表 明库特尼地体的碎片可能被捕获进大盆地内陆带的碰撞带。

#### 9.2.4 黑石弧

内华达西北部三叠纪末-早侏罗世的黑石弧(Black Rock Arc)地体(Quinn et al., 1997; Wyld, 2000)是该弧在 163~160 Ma 期间与面向西的、以碳酸盐岩为主的卢比亚被动陆缘西部碰撞的上部板块,碰撞带以向东的 Luning-Fencemaker 褶皱-逆冲断层带为标志(Wyld et al., 2003; Wyld and Wright, 2009),在该弧和卢比亚边缘之间出现了一个已崩塌的碎屑盆地(Burke and Silberling, 1973; Speed, 1978)。基于指状交错关系和产出超镁铁质到花岗质侵入体、玄武质枕状熔岩和中性熔岩、角砾岩和凝灰岩(Dilek and Moores, 1995),这些盆地相岩石明显已接近弧的性质。因此,盆地岩石可能代表了一个已变形的弧前盆地。据我所知,目前还没有约束该弧和卢比亚超级地体西部被动边缘之间大洋宽度的直接数据。洋内的俯冲明显向西到黑石弧下(图 9.2),如果如公认的那样,该弧与内华达山区相关,那么就强烈地表明该俯冲也是向西到那个侏罗纪弧之下的。

#### 9.2.5 克拉马斯弧

克拉马斯(Klamath)山脉内有几个侏罗纪弧地体。最东端的三叠纪—中侏罗世弧属于发育在古 老基底之上的雷丁(Redding)次级地体,这里被看作是巨大弧地体的一部分,该巨大弧地体从内华 达延伸穿过内华达山脉岩基直到墨西哥,这将在随后章节中讨论。

克拉马斯山地块内 177~168 Ma 的海福克弧 (Hayfork arc) 是海福克复合地体的一部分 (图 3.9),由一个二叠纪—三叠纪混杂岩的东部带和一个弧杂岩西部带组成,其中弧杂岩包括火山岩岩石和 170 Ma 的深成岩体,它们在 169~164 Ma 期间增生到北福克地体西缘 (Wright and Fahan, 1988),尽管现在海福克和北福克地体之间的断层为向东倾,这使大多数研究者认为海福克弧是向东俯冲至北福克之下的 (Harper and Wright, 1984; Irwin, 1981, 2003; Irwin and Wooden, 2001),但我认为该断层是倒转的或是较年轻的,因为断层东部没有恰当年龄的弧岩浆作用,这就表明该俯冲是向西到海福克弧之下。在该假说中,海福克地体代表着面朝东的弧增生楔状杂岩体,被更东面的克拉马斯山脉地体逆冲覆盖。正如 Wright and Wyld (1994)的研究结果所显示,弧形成于由三叠纪增生杂岩构成的基底之上,这表明:由于俯冲板片向东回撤,弧已向东迁移覆盖在自身增生杂岩的更老部分之上。

位于海福克弧西部(图 3.9)、年龄为 164~162 Ma 的约瑟芬(Josephine)蛇绿岩(Harper et al.,1994),可能代表了弧后盆地的洋壳,弧后盆地开启时间明显在更东边的海福克-北福克碰撞(Wyld and Wright, 1988)期间或之前;以同样的方式,168~161 Ma 的海岸山脉蛇绿岩似乎已经形成,在斯马特维尔弧与内华达前陆西部碰撞之前到 159 Ma 期间。在西部,可能存在的残余弧被称为罗格谷地体(Rogue Valley),形成于晚侏罗世,由厚层的安山质火山碎屑岩、角砾岩和熔岩流序列与加利斯(Galice)复理石夹层组成(Garcia,1979,1982),这两者都老于 150 Ma,该年龄是切过这两套地层的英安质岩墙群的年龄(Saleeby, 1984)。这种弧碎片可能是在内华达山脉最西部发现的更大的弧的一部分,以与马里波萨(Mariposa)组互层的火山岩为代表(Bogen, 1985; Snow and Ernst, 2008; Ernst et al., 2009b),马里波萨组被认为与加利斯组等同。

艾恩赛德(Ironside) 岩基(Charleton, 1979) 和相关的 171~168 Ma 基性-中性深成岩体(图 3.9) 侵入于海福克地体,但缺乏地壳组分的输入,西海福克后期逆冲覆盖在东海福克混杂岩之上、 具有中央地体(Central terranes)的海福克早期增生,代表了解释不清的深成岩浆作用的一次脉动 (Wright and Fahan, 1988; Barnes et al., 2006a)。如果将海福克弧作为与东面岩石碰撞时的上盘,就 可以容易地将深成岩体形成解释为由于下降板块斜坡沉积物量的增加及其在碰撞前的脱水作用,这将 导致地幔楔内熔融大量增加,并增加弧内岩浆的通量,这是 Hildebrand (2009)设想的产生科迪勒拉 型岩基的总体机制,但在这种情况下,下盘的边部是年轻的,所以,边界带的宽度很可能比长期存在 的、裂开的北美西部边缘窄得多;这样,岩浆作用脉动的时间也要短得多。

几套后碰撞侏罗纪深成岩体侵入到克拉马斯(Klamaths)岩石中。最老的岩套年龄范围为162~156 Ma(图3.9),它们侵入到海福克和更东面的地体中,结束于克拉马斯弧西部岛弧地体增生之前

(Harper and Wright, 1984; Harper et al., 1994)。一个较年轻的岩套年龄为151~144 Ma (图 19),是 由一个原始的、富含幔源 H<sub>2</sub>O 玄武岩岩浆流上涌到地壳中形成,该岩套侵入到几个以前已缝合的地 体中 (Allen and Barnes, 2006; Barnes et al., 2006b)。这两个岩套都难以用现有的模型解释,但用典 型的俯冲板片断离岩浆作用就容易理解,因为这两组岩套明显滞后于它们各自的碰撞时间,且并不局 限在板片上盘弧中,而是侵入穿过几个地体。这个原因以及板片断离岩浆弧一些更普遍的特性将在后 续章节讨论。一个白垩纪最早期 (142 ~ 136 Ma)的岩套侵入到克拉马斯中心部分的岩石中 (图 3.9),也出现在内华达前陆西部 (图 5.7 和 5.8) (Irwin and Wooden, 2001; Irwin, 2003),尽管它们 代表了岩浆作用的一个短期脉动,并很可能与俯冲板片断离有关,但这是可能的,它们是向东倾俯冲 形成弗朗西斯科杂岩的初始产物。

#### 9.2.6 结合湖-斯莱特溪弧

内华达山脉前陆西部地区,斯莱特溪-结合湖(Slate Creek-Lake Combie)弧岩石的年龄从210~172 Ma,它被逆冲到大致同时代的小提琴溪杂岩之上(图 5.8),小提琴溪杂岩由含蛇纹石化基质和蛇绿混杂岩组成,被三叠纪—早侏罗世枕状玄武岩、火山碎屑、泥质板岩和放射虫硅质岩覆盖(Edelman et al., 1989a; Fagan et al., 2001; Moores and Day, 1984; Day and Bickford, 2004)。两个杂岩体的接触带被年龄为167 Ma 的斯凯尔斯(Scales)深成岩体侵入,提供了变形事件的最小年龄约束(Day and Bickford, 2004),角闪石冷却年龄的范围为156~152 Ma(Fagan et al., 2001),被该地区—组159~150 Ma的后构造深成岩体所切割(Day and Bickford, 2004)。与斯莱特溪-结合湖同期的与弧相关的火山岩和侵入岩没有在东部出现,它们与海福克非常相似,这些岩石通常是相关的(Fagan et al., 2001; Irwin, 2003),俯冲可能是直接向西到该弧之下。在这种情况下,小提琴溪杂岩代表一个与斯马特维尔-斯莱特溪-结合湖弧有关的已垮塌的弧前盆地-增生楔杂岩,侵入于该区的159~150 Ma 深成岩体群(图 5.7 和 5.8)可能代表了俯冲板片断离岩浆作用(slab-failure magmatism)产物。

## 9.2.7 斯马特维尔弧

内华达山前陆带最外侧地体包含有厚层的侏罗纪弧岩石序列,位于小块洋壳之上(Menzies et al., 1980; Bogen, 1985; Day et al., 1985)。杂岩中最年轻火山岩的年龄为159 Ma (Saleeby, 1981),只是年龄比尤巴河(Yuba Rivers)深成岩体老(Saleeby et al., 1989),它被斯马特维尔和斯莱特溪-结合湖带之间的逆冲断层切割,但向西为变质岩石(图 5.7 和 5.8);因此,斯马特维尔地块与东部岩石的碰撞发生在159 Ma,而弧内的岩浆作用一直持续到碰撞时间。斯莱特溪(Slate Creek)杂岩的剥露作用在156 Ma 时已经开始(Fagan et al., 2001)。

自 Moores (1970) 以后,大多数研究者已认可了面朝东的斯马特维尔洋岛弧,该弧不受向西倾的 俯冲带影响,而与前陆地体西部边缘发生碰撞(Dickinson et al., 1996a, 1996b; Moores and Day, 1984; Godfrey and Dilek, 2000)。因此,在159 Ma 碰撞前的某个时期,任何东部的岩浆作用都不可能 由向东的俯冲产生,除非还有另外一个俯冲带。因此,有些研究者(Schweickert and Cowan, 1975; Schweickert, 1978; Schweickert et al., 1984; Ingersoll and Schweickert, 1986; Dickinson et al., 1996b; Ingersoll, 2008) 提出摩鹿加群岛海型模型(Moluccan sea-type models),即大陆弧往东、大洋弧往西 的双极东-西双向俯冲导致了两者之间的碰撞。这个类似情形,最近在老费里(Olds Ferry)与瓦罗瓦 (Wallowa) 弧在159~154 Ma 的碰撞中也被提出,瓦罗瓦弧位于俄勒冈州的蓝山北部(Schwartz et al., 2011)。然而,还有另一种可能性:沿着黑石-内华达地块东部边缘向西俯冲,导致有两条向西 倾的俯冲带(图 9.2)。

在这两个模型中,海岸山脉蛇绿岩都很容易地被解释为一个形成于斯马特维尔杂岩西部的蛇绿岩,在159 Ma 斯马特维尔弧上盘对接之前,形成于俯冲带之上的弧后盆地背景 (Shervais, 2001; Shervais et al., 2004, 2005)。在168~161 Ma 期间,它可能已经在与陆地有一定距离的某个近赤道纬度地区形成,且是弧后背景,这也符合 Hopson 等 (2008)的开放海模型,该模型主要基于存在有远



图 9.2 板块模型卡通图 图解本文所提出的构造模型。注意:在170~160 Ma 这张图的西部没有标注出斯莱特溪-结合湖弧 (Slate Creek-Lake Combie),因为它们与斯马特维尔弧特征基本一致

洋深海火山沉积物薄层和晚侏罗世(152~144 Ma)沉积物间断和/或角砾岩化压顶石层序(capstone sequence)。他们(Hopson et al., 2008)认为蛇绿岩必须离美洲大陆足够近才能接受通过信风传送的来自东北部的火山灰,但Hildebrand(1988)的研究显示,火山灰一般是通过对流层风从西向东传输的而不是像信风那样的低层风,因此,可能曾有另一个位于蛇绿岩西边的弧,就象约瑟芬蛇绿岩的情况那样。

位于蛇绿岩顶部的独特角砾岩,年龄明显小于154±5 Ma (Blake et al., 1987),当斯马特维尔弧 西部的洋底弯折和断裂之时,可能是由碰撞进行期间的变形作用所造成。这与喜马拉雅山印度-欧亚 大陆碰撞期间印度洋海底变形的方式很相似(Beekman et al., 1996)。Hopson 等(2008)认为角砾岩 的成因是转换断层带中的研磨作用,这可能部分正确,但碰撞期间也能产生角砾岩,因为在印度洋的 海底变形显然不仅仅导致普遍的逆断层贯穿弯曲海底区域(Weissel et al., 1980; Zuber, 1987),而且 也可使老的转换断层重新活化(Bull, 1990; Bull and Scrutton, 1990, 1992)。

在内华达山前陆内, 斯马特维尔及相关的地体显然相对较薄, 并在大陆地壳的顶部就位, 因为该 区被一套 140 Ma 的大型花岗质深成岩体所切割 (Saleeby et al., 1989; Irwin and Wooden, 2001; Day and Bickford, 2004; 图 5.7, 图 5.8),在那些增生较早的地区,有另一套年龄在约 160 Ma 的深成岩体 侵入那些地体 (Edelman et al., 1989)。外来地体是薄的、并且其下有大陆地壳,这些结论已被产出 的深成岩体本身、根据地震反射数据建立的地球物理模型,以及重力-磁力数据和地震数据的层析成 像反演等证据支持,要求中-上地壳之下的下地壳密度较低,而中-上地壳基性-超基性板片被认为是 大型蛇绿岩板片的一部分——可能将海岸山脉蛇绿岩和斯马特维尔地块之间联系起来 (Stanley et al., 1998; Godfrey and Dilek, 2000)。在变形的和后构造侵入体中的前寒武纪继承锆石,如尤巴河 (Yuba Rivers)、鹿溪 (Deer Creek) 和斯马特维尔深成岩体等 (图 5.7),为构造岩片下存在古老大陆地壳 提供了额外的证据 (Day and Bickford, 2004)。需要注意的是,人们对位于前陆带之下的大陆地壳源 区还完全不了解,源区亦位于 0.706 等值线的西侧,该线一般被认为是古老克拉通地壳的西部边界 (Moore, 1959; Kistler and Peterman, 1973; Chen and Tilton, 1991)。

对位于大陆地壳顶部的西部洋岛弧地体的认识极其重要,因为这很可能打败摩鹿加群岛海模型(Moluccan sea-type model),该模型展示了向东和向西的相向俯冲带出现在斯马特维尔和内华达山区地块之间。在摩鹿加群岛海模型的双向俯冲汇聚情形下,两个弧都没有附着在俯冲板片上,所以,除了板块运动,没有其他驱动力能使一个弧俯冲到另一个弧下面,相反,这些弧-弧碰撞是软的(Pubellier et al., 1991)。因此,这样的模型最合理的似乎是有两个向西倾的俯冲带,一个在内华达山地块东部,另一个在西部;内华达西部的岩石支持这样的模型,因为在那里,Luning-Fencemaker冲断层带代表了一个面向西的地台边缘,恰好在 163 Ma 之后向西俯冲到三叠纪—中侏罗世弧地体之下,这也是洪堡特杂岩体变形前的年龄(Wyld et al., 2001; Dilek and Moores, 1995)。

因此, 斯马特维尔和斯莱特溪-结合湖增生事件的结果是:碰撞前的任何岩浆作用与更年轻的白 垩纪内华达岩浆作用之间都不存在关系, 因为两者起源于完全不同的俯冲带。约 158~145 Ma 之间的 岩浆作用的类型和成分与其之前和之后的都有很大的不同, 这期间的岩浆作用是过渡性质的岩浆事 件,包括 500~600 km 长、年龄为 148 Ma 的独立岩墙群和各种小深成岩体, 如那些出现在戈达德 (Goddard)的顶垂体和内华达州欧文斯山 (Owens Mountain)的 152~148 Ma 同构造岩墙 (Wolf and Saleeby, 1995)和莫哈韦沙漠地区可能为 155~150 Ma 的闪长质深成岩体 (Miller and Glazner, 1995), 当前模型对这些现象鲜有解释。

#### 9.2.8 内华达山区侏罗纪弧

尽管大部分已被更年轻的走滑断层和正断层错断和扭曲,且被白垩纪—古近纪—新近纪深成岩体 所掩盖,侏罗纪内华达弧可能是一个连续弧的一部分,该弧从克拉马斯东部的雷丁次地体和内华达山 脉北部向南-南东向穿过白印优山脉、莫哈韦和索诺兰沙漠北部一直到墨西哥湾 (Busby-Spera, 1988; Barton et al., 1988; Busby et al., 2002; Tosdal et al., 1989; Mauel et al., 2011)。基于断断续续 的露头和同位素定年,这个弧在三叠纪到 160 Ma 明显是活动的,在 160 Ma 弧发生了变形并且弧岩浆 作用停止 (图 9.3), 随后的岩浆作用通常是双峰式和碱性的, 也许最好视之为是俯冲板片断离岩浆 作用,这部分内容将在后面讨论。尽管对内华达山和莫哈韦沙漠地区的侏罗纪弧已有大量研究,但亚 利桑那州西南索诺兰沙漠地区和加州东南端可能是这个弧出露最好和研究得最透彻的部分。正如变形 和未变形火成岩套的 U-Pb 定年结果所显示 (Tosdal et al., 1989), 在那里, 早—中侏罗世基特峰-特 里戈峰-卡戈穆沙舒(Cargo Muchacho)超单元的深成岩及其火山沉积围岩(菲涅尔峡谷序列)的变 质和变形作用就发生在 158 Ma 前。这些学者还证实早侏罗世低级变质岩和主要是富碱的阿蒂西亚 (Artesia) 序列位于已变形和变质的菲涅尔-基特峰岩石之上,并被主要为富碱的双峰式柯维亚 (Ko Vaya) 深成岩套所切割 (图 6.2)。类似的, 在圆顶石山 (图 5.11), 164 Ma 的花岗闪长岩和变形作 用被 161~158 Ma 未变形的淡色花岗岩切割(Boettcher et al., 2002)。沿着弧从内华达北部一直到墨 西哥索诺兰沙漠,所有弧都可以发现这两个时期的岩浆作用、弧和板片断离的残留部分(图2.5), 因此有一个连贯的事件:从三叠纪到侏罗纪弧岩浆作用、碰撞以及在约 160 Ma 弧岩浆作用停止,紧 接着出现板片断离岩浆作用、剥露。



图 9.3 克拉马斯-黑石-内华达弧等地区岩浆岩年龄测定汇总简图 如文中所述,在西雅里塔-卢比亚 160 Ma 碰撞期间,沿弧走向到处可见遭受的变形与弧岩浆作用停止一致。 碰撞后不久,新的后碰撞岩浆作用迅速开始,并被认为是俯冲板片断离岩浆作用。数据来源见正文

这个弧最令人困惑的问题是它的俯冲极性,通常认为是向西的(Burchfi el and Davis, 1972, 1975; Schweickert and Cowan, 1975; Busby-Spera, 1988; Ingersoll, 2008)。弧及其变形作用时间与内华达西北部黑石弧(Black Rock arc)相同,黑石弧是在 160 Ma 时与更东边的碳酸盐岩被动陆缘碰撞时的上盘(Dilek et al., 1988)。根据岩浆岩年龄和碰撞终止的时间,它可能代表了同一个弧向北的延续(Barton et al., 1988),如果是这样,那么就意味着侏罗纪内华达山区弧是向东俯冲,而不是人们普遍认为的向西俯冲。目前已在内华达识别出弧和下盘东部之间的碰撞带,被称为 Luning-Fence-maker 褶皱-逆冲断层带(Wyld et al., 2003; Wyld and Wright, 2009),而对其更南部却知之甚少,该区被称为内华达山区东部褶皱-逆冲断层带(Stevens et al., 1998),然而,这个褶皱-逆冲断层带的碎片却是已知的,它从鞍袋湖地区往东南穿过印优山脉-莫哈韦沙漠地区,其特点是向南西倾、北东倒转的逆冲断层,逆冲断层形成于侏罗纪弧岩浆作用主要阶段之后,但早于 148 Ma 的独立岩脉群(Dunne et al., 1978; Dunne, 1986; Walker et al., 1990b; Gerber et al., 1995; Howard et al., 1995; Stevens et al., 1998; Miller and Walker, 2002; Martin et al., 2002; Dunne and Walker, 2004; Stone et al., 2009)。在亚利桑那州的玛丽亚(Maria)褶皱-逆冲断层带中也有保存(图 4.1),该区 175~160 Ma 的火山岩发生变形并向北逆冲,时间发生在 154 Ma 的火山岩喷发和沉积之前(Reynolds et al., 1986; Richard et al., 1987; Spencer et al., 2011)。

评估是否可能有向东的俯冲非常困难,因为那时唯一的向东倾俯冲带在内华达中部的盆地内,而 那里有两个向西的碰撞,即在167 Ma和159 Ma分别发生在斯莱特溪-结合湖和斯马特维尔弧向西的 碰撞事件,即便有这样一个盆地,也不知道其宽度,因为它在随后的105~100 Ma内华达山脉东部和 西部之间挤压转换事件(后续分章节讨论)中关闭了,而且甚至也不知道盆地中是否存在过大洋岩 石圈。一个不太有说服力但是或许很重要的证据是,对内华达山区前陆西部内南北走向的接触带观 察,而侏罗纪弧是南东-北西走向的,这样在莫哈韦-索诺拉和内华达山区南部之间就产生了明显的 离散。总体而言,现有资料支持这一假说:侏罗纪黑石-内华达形成了一个连续的面朝东的弧地体, 它在160 Ma时因碰撞而关闭,它与卢比亚超级地体西部边缘碰撞并向东逆冲覆盖其上,随后俯冲板 片断离产生了双峰式碱性岩浆作用群。

侏罗纪卡梅尔 (Carmel) 组位于犹他州南部的科罗拉多高原,该组泥石流沉积和火山颗粒内出现 火山灰凝灰岩的巨砾,而砂粒出现在细粒沉积中 (Chapman, 1989, 1993; Blakey and Parnell, 1995), 这些组分很可能是从菲尼克斯 (phoenix) 断层南部的侏罗纪火山弧地体中向北流出的。鉴于断层可能经过了 1500 km 的左旋滑动 (Hildebrand, 2009),最初的源区可能就位于现今的东南部,或许在 Oaxaquia 内 (图 2.5)。Oaxaquia 有侏罗纪弧岩石,形成于以格林威尔期 (Grenvillian) 基底为主的背景,弧岩石大部分被二叠纪—三叠纪沉积岩所覆盖 (Ruiz et al., 1988; Keppie et al., 2001, 2003; Solari et al., 2003; Ortega-Obregón et al., 2003; Centeno-García and Silva-Romo, 1997; Jones et al., 1995; Barboza-Gudiño et al., 2004),这与科罗拉多高原南部侏罗纪岩石中的碎屑锆石分布剖面相当 吻合 (Dickinson and Gehrels, 2009; Mauel et al., 2011)。

#### 9.3 俯冲板片断离

有许多地质学家对板片破坏或断离的概念不熟悉,因此需要有一个简短的回顾。板片破坏,或有 时被称为板块断离或拆沉,是俯冲的必然结果,而科迪勒拉地质学家们慢慢才接受了它的影响。早在 1981年,地震学家们已认识到当大块大陆块体的边缘部分俯冲时,它们的浮力会导致俯冲板片不能 抵达,或非常靠近大陆架前缘(McKenzie,1969;Osada and Abe,1981;McCaffrey et al.,1985;Welc and Lay,1987),这是因为阻止大陆岩石圈俯冲的浮力与大洋岩石圈的向下拉力一样大(Cloos, 1993)。最终,大洋岩石圈的较大密度导致板块下盘在最弱处被撕裂并下沉到地幔中。当俯冲板块破 坏、板块下盘挣脱了大洋板块的束缚时,部分已俯冲的大陆边缘岩石会在浮力作用下上升。 Hildebrand(2009)认为,因为撕裂通常是穿时的,板块断离期间,板块仍在汇聚,虽然速度缓慢。 汇聚和上升共同作用导致强烈的摩擦阻力使板块上盘和下盘之间接触面拓宽,挤压板块下盘地壳产生 厚皮构造,例如褶皱和贯穿地壳的逆冲断层。在新几内亚非常年轻的弧-陆碰撞带中,Mapenduma 背 斜(厚皮背斜逆冲构造)就是碰撞减弱阶段发育在碰撞前陆带的(cloo et al.,2005)。美国西部的 拉勒米厚皮构造在构造背景和变形时间上与新几内亚的相似性,使 Hildebrand(2009)确信它们的形 成方式相同。已知在安第斯(Andes)山脉南部的潘佩阿纳斯山脉(Sierras Pampeanas)和阿根廷的 麦哲伦(Magellanic)前渊(Fosdick et al.,2011; Cristillini et al.,2004),Wopmay 造山带(Hoffman et al.,1988)以及其他加拿大造山带内(Hoffman,1989)也有类似构造。

板片断离的另一个重要特征是相关的岩浆作用。当一个俯冲板片遭到撕裂和破坏时,通过撕裂通 道软流圈(物质)可上涌高达100 km,并绝热熔融产生岩浆,这些岩浆可以侵入到板片破裂处的下 盘(Hildebrand and Bowring, 1999),或在撕裂处上方的上盘(McDowell et al., 1996; Housh and Mc-Mahon, 2000; McMahon, 2000a, 2000b; Chung et al., 2003),或两者兼而有之。决定板片断离岩浆作 用模式最重要的因素或许是拉伸率和实际断离速度(cloo et al., 2005)。因为地幔和地壳的不均一, 由此产生的岩浆可能在时间上和空间上变化性很大(Housh and McMahon, 2000),并且,从单纯的软 流圈熔融到复杂的地壳熔融都有,当然也包括两者的混合。

通常情况下,板片断离岩浆作用发生在碰撞期和随后的弧关闭时期。如果岩浆侵入到板块上盘,可能会在老的弧顶部或沿其旁边形成一条线性岩浆带,它们与较老的岩浆作用在时间上是连续的,因此很容易混淆。岩浆也可能侵入到前渊和/或已缩短的板块下盘被动边缘的岩石中,或两者兼而有之(Hoffman, 1987; Hildebrand and Bowring, 1999; Hildebrand et al., 2010a)。

## 9.4 侏罗纪俯冲板片断离岩浆作用

大部分侏罗纪弧岩浆作用在 160 Ma 之后停止,随后独立岩墙群侵位于内华达山脉中-东部、并向 南到莫哈韦沙漠(James, 1989; Carl and Glazner, 2002; Hopson et al., 2008);独立岩岩墙群由 Moore 和 Hopson (1961)首次识别出, Chen 和 Moore (1979)最早对其定年。现有模型难以解释这些岩墙群 (Dickinson and Lawton, 2001),但由于它们只是比 Luning-Fencemaker 碰撞晚了几个百万年,它们可 能代表了与向西俯冲的板片断离有关的岩浆作用,这可能归因于卢比亚大陆地壳的俯冲难度(图 9.2)。板块断离很容易解释这个长的、线形的和非常狭窄的岩墙群,以及它局部的双峰式特征(Mc-Manus and Clemens-Knott, 1997)和一般呈碱性(Karish et al., 1987)的特点。在内华达山脉的同一 90 地区内,有一些90 Ma的白垩纪的岩脉,它或多或少与独立岩墙群共线的,但这两期岩浆作用之间有 60 Ma的间隔,表明它们是不相关的,较年轻的岩岩脉可能只是与个别内华达山深成岩体有关 (Cole-man et al., 2000)。

比 160 Ma 年轻的岩浆作用明显晚于沿弧线的变形作用,以戈达德顶垂体、莫哈韦沙漠中部和亚 利桑那州南部等地为代表(图 9.3)。前面已讨论过的双峰式碱性深成岩线性带,在莫哈韦还没有命 名,但在亚利桑那州南部被称为柯维亚(Ko Vaya),与岩脉群年龄大致相同,也可能与俯冲板片断 离有关,因为它们不是典型的与俯冲相关的岩浆岩。沿着岩脉群,它们的出现范围超过了大部分已经 封闭弧的长度。岩脉群和双峰深成岩套总体上已被认为是连续的侏罗纪弧岩浆作用的一部分,但是主 要时期不包括其中的一段弧内挤压时期。依据成分确定板片断离岩浆作用是困难的,而岩浆带长条线 性分布的特征及其奇特的双峰式组成等表明它不是典型的弧岩浆作用。然而,它出现在 160 Ma 的变 形和弧的关闭之后,就在著名的科迪勒拉型内华达山区岩浆作用前,这使其成为一个研究俯冲板片断 离岩浆作用的最佳对象。

## 9.5 蓝山山脉拼合

在蓝山拼合体内有老费里(Olds Ferry)和瓦罗瓦(Wallowa)两个弧地体,以及它们夹持的贝克(Baker)地体增生楔,在增生岩石逆冲在每个弧之上时,以软碰撞的方式连接起来(Avé Lallemant, 1995; Dorsey and LaMaskin, 2008)。碰撞事件介于159 Ma(最年轻的变形沉积岩年龄)和154 Ma之间(154 Ma 为金甲虫(Goldbug)深成岩体年龄,它切过贝克地体内的Greenhorne-Bourne 隐伏地体的断层接触面)(Schwartz et al., 2011)。虽然不知道该拼合地体在侏罗纪末时位于何处,但到晚白垩世时它作为Ochoco盆地属于卢比亚拼合体的一部分,它不整合在更老的岩石之上,在更南部是霍恩布鲁克盆地的一部分(Surpless et al., 2009)。关于晚侏罗世—早白垩世时它的位置,一条线索来自麦克劳德(McCloud)动物群,该动物群出现在贝克地体增生楔岩石中,也出现在侏罗纪内华达弧西部地块内的二叠纪—侏罗纪增生杂岩中,并可能延伸到更北。两个三叠纪—侏罗纪弧地体与更北面的Stikine和Quesnelliam弧和更南面的克拉马斯-内华达山区带类似,所以,总的来说有理由将它们包括在这些大范围的岩石之内,它们在150 Ma加入到卢比亚地体。蓝山超级地体和贝尔特-珀塞尔-温德米尔湖地块在爱达荷州东部的接触带为一大型的走滑断层,推测可能是存在于Quesnellia西边的Intra-Quesnellia断层的一部分(Irving et al., 1995)。

## 9.6 侏罗纪卢比亚联合大陆

从以上分析能够清楚地看到,中侏罗世到晚侏罗世是卢比亚超级地体主要拼合时间。由不同类型的三叠纪—侏罗纪弧组成的西部地块构成了西部超级地体,我称其为西雅里塔(Sierrita),它在大约160 Ma 拼合到被称为原卢比亚(Proto-Rubia)的东部,形成卢比亚带状大陆,或称为巨型地体(megaterrane)。那时,除了少量游离在西部的地块,几乎所有的地体都已拼合到卢比亚带状大陆。俯冲通常向西,导致更多东边的地体被拉到更西边的地体之下。在斯马特维尔和罗格谷(Rogue Valley)弧以及它们的边缘海盆地岩石圈板片增生后,沿着卢比亚西部边界开始了新的向东俯冲。

## 9.7 卢比亚西部边界俯冲

在159 Ma之后的某个时期,在斯马特维尔增生地块的西部开始了新的向东俯冲,与此次俯冲相关的新岩浆作用出现在比老的侏罗纪俯冲带更往西的内华达山区前陆 (Saleeby, 1981),岩浆作用向西跃迁显然代表了海沟位置从增生弧东侧向斯马特维尔增生地块西侧的跳跃。一套年龄为142~140 Ma 的深成岩体切割了内华达山区前陆所有地体和克拉马斯 (Hietanen, 1973; Irwin, 2003),但是,据我所知,内华达岩基中只有少数岩体可能与这次俯冲有关;如果是这样的话,那么在接下来的15 Ma 里岩浆作用将是稀少乃至缺失的。

最老的向东俯冲的明确证据就出现在弗朗西斯科俯冲杂岩内,在那里,最东边的带----南福克山

片岩中的碎屑锆石年龄年轻至131 Ma (Dumitru et al., 2010)。这不同于普遍接受的想法,就是试图 测定初始俯冲的时间,通过测定中央带 (Central belt) 泥质混杂物中最古老的高级变质块体,尽管这 些块体周围已被蛇纹岩壳包围 (Cloos, 1986),且表明它们在混杂并入之前就已经被改造了。高度变 质块体可能源自东部更老的带,并在物质迁移过程被分散开 (Jayko, 2009)。不管怎样,如果弗朗西 斯科杂岩最初形成于内华达山区前陆西部,它应该不老于159 Ma,并且在131 Ma 之后明显活动,表 明其沿着卢比亚超级地体西部边缘主动地向东俯冲。

如上所述, 白垩纪内华达山脉中最老的深成岩体是 125 Ma 沿岩基西部分布的基性-中性岩体 (Saleeby and Sharp, 1980; Stern et al., 1981; Bateman, 1992; Clemens-Knott and Saleeby, 1999), 因为 其间有 15 Ma 的岩浆作用间隔, 故它们与向东的俯冲没有简单的相关关系。这就对内华达山脉岩浆起 源于洋壳, 向东俯冲到内华达山脉之下的模型提出了挑战。在后续讨论中我们将看到, 内华达山脉岩 浆作用可能与向东俯冲根本就没有关系, 反而是向西俯冲引起的 (Hildebrand, 2009)。另一种可能性 是, 也许在那个时候弗朗西斯科杂岩、海岸山脉蛇绿岩和大谷地群还没有与内华达山毗邻 (Wright and Wyld, 2007)。

## 9.8 北极阿拉斯加-Angayucham 碰撞

晚侏罗世—早白垩世期间,北极阿拉斯加(Arctic Alaska)开始试图俯冲到科尤库克(Koyukuk) 弧之下,由于北极阿拉斯加的北坡(North Slope)隐伏地体的岩石不同于北美的,这个碰撞事件被认为发生在卢比亚超级地体内,而那时它与北美是分开的。北极阿拉斯加在泥盆纪埃尔斯米尔(Ellesmerian)造山运动时期经历了强烈变形和变质作用,其上覆盖的早密西西比世前渊序列,如罗伯茨山 外来体,是比北美大陆西部典型得多的卢比亚超级地体内的各种地体,北美西部完全没有泥盆纪变形 或来自西部的造山带碎片脱落。

在一般的碰撞模式中,当 Angayucham 洋关闭、科尤库克地体与北极阿拉斯加碰撞时,阿拉斯加的晚侏罗世—早白垩世布鲁克(Brookian)造山运动开始发生,科尤库克地体代表了一个板块上盘的弧地体,在侏罗纪末-尼奥科姆期(Neocomian)期间与北极阿拉斯加碰撞(Roeder and Mull, 1978; Box, 1985; Box and Patton, 1987; Mull, 1985; Moore et al., 1994)。碰撞的准确时间目前尚不清楚,困难在于如何解释晚侏罗世—尼奥科姆期Okpikruak组的强烈变形,该组岩石位于恩迪科特(Endicott)和德隆(DeLong)山外来体中部和布鲁克斯山脉西部,被认为代表了南移来的外来体碎片,由此可以确定碰撞时间(Moore et al., 1994),主要问题是Okpikruak组的岩石中没有与布鲁克斯山脉外来体岩石相匹配的碎屑锆石图谱(Moore, 2011,个人交流),所以,Okpikruak可能是碰撞前的岩石。较年轻的碰撞年龄能较好地与以下证据吻合:科尔维尔(Colville)盆地在巴列姆阶—阿普第阶(Barremian-Aptian)时期突然向下挠曲(Cole et al., 1997),并明显与利斯本(Lisburne)山的向西逆冲一致(Moore et al., 2002);也与科尤库克弧最上部的巴列姆期—阿普第期岩浆作用(Box and Patton, 1989)和 3400 m 厚的巴列姆阶—阿普第阶(Aptian-Cenomanian)砂岩、砾岩、泥岩和前 渊煤层相一致(Mull, 1985; Siok, 1989; Crowder, 1989; Bird and Molenaar, 1992)。推断与隆升有关的冷却事件分别发生在 100±5 Ma, 60±4 Ma 和 24 ±3 Ma (O'Sullivan et al., 1997)。

可能至少有部分与这个碰撞作用相重叠的是加拿大盆地的开启,如沿着北极阿拉斯加北部边缘(当前坐标)侏罗纪裂谷盆地的形成,并最终在欧特里夫期(Hauterivian)形成新洋壳(Grantz and May, 1982)。有些学者基于大量的古地磁学资料提出,在盆地打开期间,北极阿拉斯加裂开并沿着加拿大北极群岛北部旋转,因而是共轭边缘(Grantz and May, 1982; Ziegler, 1988; Embry, 1990; Plafker and Berg, 1994; Grantz et al., 2011);但 Lane (1997)提出了几条证据质疑共轭模型的可行性,如泥盆纪变形的年龄不同、裂开-漂移转换的时间不同,以及如果按照旋转 66°复原,那么俄罗斯大陆架与加拿大北极群岛将有 600 km 重叠;麦克唐纳等(2009)对比了新元古界地层得出了同样的结论。古地磁数据显示,在130 Ma 时北极阿拉斯加逆时针方向旋转了 105°+49°/-43°,位于现今位置的 12°±5°南(Halgedahl Jarrard, 1987),这似乎很难与那时美亚盆地(Amerasian Basin)的开始92

相一致,除非当时有一个很长的、向南倾的裂谷手臂状海湾。最近,Helwig等(2011)解释了从麦 肯齐三角洲(Mackenzie River delta)到班克斯岛(Banks Island)波弗特海(Beaufort)的地震剖面, 得到了一个已经死亡的扩张中心的影像,它可能代表了这个裂谷的北部。

这两个区域的岩体带成分变化很大,从钙碱性到碱性岩体,它们侵入到育空-科尤库克盆地-苏 华德半岛和红宝石地体中(Patton and Box, 1989; Miller, 1989; Arth et al., 1989a; Patton et al., 1987, 2009; Till et al., 2010),明显切割了(Miller, 1989)红宝石-Angayucham package 逆冲断层 (Patton et al., 2009),它们的年龄为112~99 Ma,所以,对于形成布鲁克造山运动而言显然是后碰撞 的。我认为这些关系表明:它们是在北坡(North Slope)地体试图俯冲期间,北极阿拉斯加板片被破 坏时形成的板片断离深成岩体,就如最早 Wartes (2006)所提出的那样。

许多以前的研究者已经发现,阿拉斯加地体与整个科迪勒拉山脉中的地体群类似。总的来说, Farewell 地体有与塞尔温盆地和罗伯茨山外来体相似的序列:寒武系至泥盆系深水沉积岩和泥盆系至 宾夕法尼亚系的碳酸盐,泥盆系和三叠系含磷酸盐黑色页岩、重晶石和砂岩,以及各种各样的辉长岩 岩床和枕状玄武岩(Bundtzen and Gilbert, 1983; Bradley et al., 2006)。这些岩石在岩性、年龄和成矿 作用方面与库特尼地体内普遍发育的岩石也很相似(Smith and Gehrels, 1992a, 1992b; Colpron and Price, 1995),而且,在很多方面,如存在大量重晶石和锌矿床(Kelley and Jennings, 2004),也与北 极阿拉斯加的岩石相似。时间和岩性上显著的相似性暗示这些地区的岩石可能起源于同一盆地 (Turner et al., 1989),并在与北美大陆碰撞之前和/或期间沿卢比亚分布。Bradley等(2007)通过U -Pb年龄和化石研究认为,从新元古界晚期到泥盆纪,Farewell、基尔巴克(Kilbuck)和北极阿拉斯 加地体是一个位于西伯利亚和劳伦古大陆(Siberia and Laurentia)之间的微陆块部分,但也有可能, 只有地体中最古老部分是西伯利亚起源的,而年轻的有不同的演化历史(Malkowski et al., 2010)。 阿拉斯加-育空边界地区的育空-塔纳纳高地地体(Dusel-Bacon et al., 2006)也是类似的,但正如 Bradley等(2006)所指出的那样,侏罗纪—白垩纪历史有很大不同,因此,它很可能后来会与 Farewell 地体并列。

所有这些地体中都存在上古生界的变形,然而,北美被动陆缘的岩石中却没有。这些碰撞事件 (其中可能有几个)通常在相邻盆地具有同时期的沉积,很可能代表了前渊充填沉积,如奥克尔-伍 德(Oquirrh-Wood)流域盆地(Geslin, 1998),某些碳酸盐杂岩有可能代表了缝合线(McLeish and Johnston, 2011)。

一个可能的联系是 Angayucham 和海岸山脉蛇绿岩的年龄类似,这点经常被忽略(Moore, 2010, 私人通信)。海岸山脉蛇绿岩与克拉马斯的约瑟芬和北喀斯喀特山的英格尔斯(Ingalls)蛇绿岩的年龄相同(MacDonald et al., 2008),更加令人难以置信的是整个科迪勒拉山脉中的蛇绿岩年龄都相同。 所有的蛇绿岩都在 170 Ma 和 160 Ma 之间形成,并分布在每个地区的一些最外侧单元,这就表明它们形成于同一个古海洋。在科尤库克弧内分散出露着的中侏罗世—二叠纪火山岩和侵入岩,位于 Ang-ayucham 岩石的南侧(Box and Patton, 1989),所处外侧位置与罗格(Rogue)弧到约瑟芬蛇绿岩的 类似。

## 9.9 北美西部塞维尔褶皱冲断层带和早期碰撞

西部内陆盆地的形成通常被解释为前渊(Price, 1973; Kauffman, 1977; Jordan 1981; Beaumont, 1981),是对西部的塞维尔褶皱-逆冲断层带逆冲加载的挠曲响应,是对大盆地分区塞维尔褶皱-逆冲断层带内逆冲年龄的最好估计。因为由岩石圈的弹性挠曲产生的盆地沉降开始与加载是同时发生的(Turcotte and Schubert, 1982),可以通过对前渊底部的地层测年来确定盆地挠曲半波长内逆冲的开始时间。盆地最老沉积物位于124 Ma 的砾石顶部(图 5.3),该砾石是北美地台骑压在隆起的外缘处时形成,最古老沉积物提供了开始逆冲的最大年龄估计,就是阿普第期(Aptian)。这个年龄与其他方法估计的该带初始逆冲时间吻合(Heller et al., 1986; Heller and Paola, 1989)。

对于这个年龄,如今的问题是那时紧邻的内陆带中没有变质作用,也没有地壳增厚,如前所述,

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已知该区两期变形和增厚的时间分别为侏罗纪和晚白垩世。它们都与塞维尔逆冲主要阶段的年龄不一致,即不小于 94 Ma,甚至可能不小于中阿尔必期。要解决这个问题需要向北往加拿大分区去寻找,在那里,Bourgeau 逆断层是最西部而且可能是最老的,它冲断并切割了北美地台的岩石(Larson et al., 2006; Price,出版中)。在正东的向斜下盘和构造上的冲断层下方是森诺曼-桑托阶的阿尔伯塔群 沉积岩(Leckie and Smith, 1992),这表明影响该区北美地台岩石的逆冲作用发生在桑托期或之后;西面可见很多细粒碳酸盐岩和碎屑物,与地台岩石大不相同,背向褶皱的逆冲断层被<sup>40</sup> Ar / <sup>39</sup> Ar 年龄为 108 Ma 的侵入体所切割,所以那里的变形比东边的逆冲老约 25 Ma (Larson et al., 2006)。就在Bourgeau 逆冲断层西边其间有一大群早古生代火山口(图 7.1)(Pell, 1994),表明在那里可能出现 过重要的缝合带(Burke et al., 2003; Johnston et al., 2003)。大部分贝尔特-珀塞尔和温德米尔也位 于这个构造带的西部,同样,马塞尔温盆地巨大的逆冲岩席也在 104 Ma 以前就位 (Gordey and Anderson, 1993; Mair et al., 2006)。

如上所述,除了科迪勒拉岩基内 105~100 Ma 的转换挤压碰撞,卢比亚内的大部分地体都是侏罗 纪拼合的,因此其他部分获得的古地磁数据可以为塞维尔事件期间卢比亚超级地体与北美碰撞后的迁 移提供约束。Kent 和 Irving (2010)最近构建了一个新的北美综合磁极移曲线,其结果显示从三叠纪 到早白垩世,卢比亚比北美向北移动得更缓慢,所以两者之间的剪切是左旋的;但从白垩纪末至始新 世,卢比亚与北美之间的剪切是右旋的。因此,在卢比亚与北美大盆地地块碰撞后,它以及卢比亚超 级地体剩下部分相对于北美继续向南移动 (Kent and Irving, 2010)。这些左旋运动可能持续到约 80 Ma,直至北美开始向南移动,卢比亚与北美之间的相对运动才转变为右旋。现今的工作支持 Ave Lallemant 和 Oldow (1988)最早提出的"三叠纪—中侏罗世外来地体沿北美边缘左旋迁移,接着右旋迁 移"的观点。

鉴于加拿大分区和大盆地分区的变形作用在时间上有很大不同(不匹配),我认为大盆地地区 "缺失"的碰撞者现在位于加拿大主山脊(Main Ranges)西部的岩石内,在塞维尔造山运动期间最早 与北美西部碰撞的正是加拿大分区。随后,它们相对于北美向南移动,只是在 80 Ma 后返回,这与莫 纳西(Monashee)杂岩体内原地岩的变质作用很好地吻合(图 7.1)。该区在约 125 Ma 时开始增厚, 并在约 60 Ma(Parrish, 1995)时快速剥露,并与贝尔特--珀塞尔外来体(Sears, 2001)和前陆盆地 (Price and Mountjoy, 1970; Ross et al., 2005)的剥蚀时间相同,这也与加拿大分区内奥米尼卡弧的科 迪勒拉岩浆作用在约 105~100 Ma 时结束(Hart et al., 2004)非常吻合。

## 9.10 塞维尔碰撞与弗朗西斯杂岩和大谷地群的响应

在北美大陆试图俯冲至卢比亚超级地体之下期间,沿碰撞带东部边缘有大量的沉积作用证据,因为这些证据在西部的内陆盆地保存得很好,然而,很少有人关注到卢比亚西部边缘沉积作用的影响,而在碰撞带内已抬升地区的物质可能往西以及往东流失。事实上,Dumitru 等(2010)注意到弗朗西斯杂岩体一个主要变化是在123 Ma 由非增生类型到增生类型,并指出:基于前人的工作(Clift and Vannucchi, 2004; Scholl and Von Huene, 2007, 2010; Von Huene et al., 2009),对变化到增生状态最好的解释是海沟的浊流沉积。Dumitru 等(2010)还指出,弗朗西斯沉积物的变化与大谷地群岩石在约125 Ma(图 5.16)时主要的岩相变化(Ingersoll, 1983)和主要的不连续(Constenius et al., 2000)一致,该主要的不连续以断层、翘曲和侵蚀为标志。这两个事件的时间如此接近,反映了北美和卢比亚超级地体东部边缘之间的碰撞开始,我认为北美西部边缘的试图俯冲使整个卢比亚产生变形、隆起和侵蚀作用,并使大量的沉积物被剥蚀,由河流向西搬运到弧前和海沟。由于那个时期在中纬度地区,北美西部边缘大致呈南-北向展布(Kent and Irving, 2010),来自太平洋盆地向东移动的极地风暴加剧了卢比亚内已抬升区域的剥蚀效应(Hoffman and Grotzinger, 1993),产生了大量的沉积物。这种情况和印度与欧亚大陆碰撞类似,那里大量的沉积物向南流进印度洋,并被印度东部和西部向北倾的海沟淹没(White and Louden, 1982; Kopp et al., 2000, 2001)。

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#### 9.11 板块因碰撞而重组

125~120 Ma 时发生了几个重要事件,影响到造山带的大部分地区:①薄皮逆冲作用在塞维尔褶皱-逆冲断层带中启动;②科迪勒拉山系岩浆作用开始;③大谷地群内的沉积岩变形;④弗朗西斯增生杂岩体被沉积物完全覆盖;⑤太平洋板块开始向北漂移。我将所有这些事件与北美克拉通的试图俯冲联系起来,克拉通的被动边缘盖层位于卢比亚超级地体之下(图9.2)。

当北美与卢比亚在约 125 Ma 碰撞时,这两个板块之间的汇聚速度可能开始大幅地减小,相邻板块的板块边界——连同它们的运动矢量——很可能被重新组织,正如印度板块与亚洲板块碰撞过程中一样(Copley et al., 2010)。因此,古大洋的关闭和卢比亚超级地体的增生对于沿着太平洋盆地内的西部边界可能有重大影响。事实上,正如 Dumitru 等(2010)指出那样,这个时间在误差范围内与太平洋板块磁极移轨迹(APWP)的一个主要尖端(由南至北迁移)相对应(Beaman et al., 2007; Sager, 2007)。

#### 9.12 大盆地分区俯冲拆沉岩浆作用和阶梯状断层

卢比亚带状大陆与北美在大盆地分区碰撞后,弧岩浆作用停止。在奥米尼卡带内这发生在约105~100 Ma (Hart et al., 2004; Johnston, 2008),在加拿大分区东部的弧地体,根据变形时间,被认为已和卢比亚分区碰撞。接着以加拿大南部塞维尔逆冲为特征的巨大外来体就位,根据那里的逆冲结束,得知塞维尔逆冲作用明显在108 Ma 时终止 (Larson et al., 2006);而根据犹他州的资料,大盆地分区的逆冲在约105 Ma 终止 (Lawton et al., 2007),几个年龄为96~90 Ma 线状排列的深成岩套侵位于上覆的卢比亚超级地体之中,其中包括加拿大北部墓碑-钨-梅奥 (Tombstone-Tungsten-Mayo)岩套中的数百个小型矿化深成岩体 (图 5.12)和阿拉斯加东部利文古德 (Livengood)和费尔班克斯-萨彻 (Fairbanks-Salcha)岩套 (Hart et al., 2004, 2005; Reifenstuhl et al., 1997a, 1997b; Newberry et al., 1990, 1996)。在恢复了被 Tintina 断层分离的约 430 km 后,深成岩体带沿着走向延伸可达1000多千米 (Gabrielse et al., 2006)。深成岩体在成分上变化很大,但以碱性岩为主,岩石组合为黑云母花岗岩,二长花岗岩和石英二长岩,与白钨矿砂卡岩、铜、铋和金矿化有关 (Hart et al., 2004, 2005)。

在大盆地分区内,俯冲板片断离意味着那里必定是北美板块下盘位于这个分区的两边,并且是(STEP)断层(STEP=俯冲-转换边界传递断层,Govers and Wortel,2005)的克拉通延续的阶梯,因为那时在阿拉斯加和加拿大南-索诺兰沙漠分区的俯冲板片并没有被破坏。有证据支持在北美和加拿大分区之间有一个撕裂出现在前陆,那里路易斯和克拉克线性构造两边的沉积作用很不相同,沿着南边分布的要厚得多(Wallace et al.,1990)。根据这些沉积物,该区由辫状、交织断层组成,断层活动时间从约100~78 Ma:基本上从塞维尔到拉勒米变形期,往南分布着菲尼克斯断层,前渊沉积显然没有延续到这个断层的南部。

#### 9.13 白垩纪科迪勒拉岩基

白垩纪科迪勒拉岩基,以内华达岩基岩浆作用为典型代表,但也以遍布于科迪勒拉山系的其它岩体为代表(图2.5和5.12),它们一直被认为是向东俯冲到北美板块西部边缘之下的产物(Hamilton, 1969a, 1969b)。然而,这一解释存在很多问题和复杂性,值得进一步进行更详细的研究。基于Ducea(2001)岩基代表异常高的岩浆通量的观点,Hildebrand(2009)提出,岩基起源于北美被动陆缘向外延伸部分向西俯冲期间沉积物的脱水作用,而DeCelles等(2009)试图将其解释为近400km的北美中-下地壳的俯冲和熔融作用,以及其后发生的大密度熔融残余的拆沉。这两种模式都需要在内华达岩基下面有向西俯冲,以拉动具浮力的大陆边缘冲向产生岩浆的地幔深处,虽然DeCelles等(2009)并没有提出将大洋板片作为驱动力,而是提出了将克拉通与向东俯冲的大洋岩石圈之间的强烈耦合作用作为驱动力。

通常情况下这个假设似乎是合理的:由于过程相同,无论是大洋壳下还是大陆地壳之下的玄武岩 通量都应该大致相同。首先,我们来看看大洋弧中岩浆通量的最佳估算,然后评估后碰撞时期(即 100 Ma之后)科迪勒拉岩基(这是研究最好的部分)的近似通量。尽管它们没有得到很好的约束, 对到达大洋弧壳底部的玄武质岩浆量的估计变化很大,每千米弧长在1~100 km<sup>3</sup>/Myr (Marsh, 1979; Reymer and Schubert, 1984; Crisp, 1984; Taira et al., 1998; Holbrook et al., 1999; Larter et al., 2001; Dimalanta et al., 2002; Scholl and von Huene, 2009; Stem and Scholl, 2010; Schmidt and Jagoutz, 2012)。在半岛山脉岩基的大陆地壳内,Silver和 Chappell (1988)估计东部后碰撞时期 La Posta 岩体的通量约为每千米弧长75 km<sup>3</sup>/Myr,其依据是侵入体平均厚度为20 km。对于海岸(Coast) 深成杂岩的后碰撞侵入体(100~80 Ma),Gehrels等(2009)估算,其岩浆通量为每千米弧长40~50 km<sup>3</sup>/Myr。在内华达岩基东部,内华达山脊(Sierran Crest)岩浆岩出现在50 km宽的带上,侵位时间 在15 Ma左右(Coleman and Glazner, 1988),如果假设岩浆岩有10 km厚,计算出来的岩浆通量速率 为每千米弧长33 km<sup>3</sup>/Ma/km;如果假定它们有20 km厚,岩浆通量速率将翻倍即每千米的弧长66 km<sup>3</sup>/Ma/km。虽然大洋和大陆岩浆通量估算的方法不怎么稳定,但是白垩纪科迪勒拉岩基可能的岩 浆通量速率与估算的大洋弧通量速率一致。

与 Hildebrand (2009) 提出的碰撞模型不同, DeCelles 等 (2009) 提出了一个关于科迪勒拉岩基 起源的模型,包括在弧后背景下由大陆岩石圈下部俯冲,以及被软流圈熔融该部分岩石圈而产生再循 环岩浆。即使背景为弧后,基于几个推理,他们的模型也是站不住脚的。

首先,该假说是假定所有三叠纪—侏罗纪的岩浆作用都与同一个俯冲过程有关,正如我们已经见 到的,这种可能性不大,这是因为他们的模型没能考虑到大多数的变形事件。其次,因为他们的弧后 模式需要内华达岩基东部的薄皮缩短量达到约 3~400 km (DeCelles, 2004; DeCelles et al., 2009), 需要去除大约 3~400 km 上述地区的克拉通基底,以达到地壳平衡。一些学者 (Ducea, 2001; De-Celles, 2004; Ducea and Barton, 2007; DeCelles et al., 2009) 提出,这 3~400 km 的克拉通地壳消失 在内华达岩基之下,在那里它们被部分熔融形成了内华达岩浆和致密的熔融残余,然后沉入地幔 (图 9.4)。在他们的模型中极端的和难以解释的是克拉通地壳要俯冲 300~400 km, 而没有附加上大 洋岩石圈以抵消和克服克拉通地壳的浮力,即使是有可能的,该模型也存在严重的质量平衡和空间方 面的问题。考虑到内华达岩基有 100 km 宽, 那么一条约 3~400 km 宽、30~40 km 厚的条带意味着要 有足够多的北美地壳被底侵,从而使内华达山脉地壳加厚至 120~150 km,考虑到剥蚀量,这似乎太 多了。融熔地壳和使熔融残余拆沉不会有助于解决问题,因为大陆地壳二氧化硅平均值为约61% (Rudnick and Gao, 2003), 根据 Ducea (2002) 估算, 内华达山脉地壳上层 30 km 含二氧化硅约 65%, 二者只有6%的差异, 这意味着由于必须熔融几乎等体积的克拉通地壳, 才能产生具有内华达 岩基总成分的相同地块,这里几乎没有残留物。这样的话,熔融作用没有去掉地壳,而只是将其向上 分配了,即使你仅仅想熔融中、下地壳,它的组分与内达华岩基相比将会有大约11%的差异,这样 仍然几乎没有残留物落下或拆沉进入到地幔,而这一切都还没有涉及到熔融所有这些地壳所必需的能 源,它们源自何方?

根据调查,科迪勒拉岩基在形成时间上与塞维尔褶皱-逆冲断带内发生的薄皮逆冲作用时间一 致,Hildebrand (2009)认为它们形成于北美被动陆缘向西未遂的俯冲过程中外侧的大陆隆起楔形体 内沉积物的脱水作用。然而,岩基是复式体的认识暗示了可能存在另外的解释,所有这些都很复杂, 并没有得到很好地约束,这在很大程度上是因为对不同时期内华达岩基和弗朗西斯科杂岩体之间的相 对位置,以及两者相对于北美克拉通所在的位置等知之甚少。此外,向东俯冲到内华达岩体西部之下 的开始时间也可能比先前认为的年轻很多。首先,缺乏早于125 Ma 弧岩浆作用的证据,老于侏罗纪 末一早白垩世的岩浆作用仅限于短期脉动岩浆作用,这些短暂的岩浆脉动事件更像是与内华达西部前 陆的碰撞事件有关,包括159~150 Ma 的岩浆脉动和142~140 Ma 之间短时爆发的岩浆活动,前面一 期可能与斯马特维尔碰撞之后发生的板片断离有关,而后面一期,所有模型都没能很好地解释。在西 部弧地体内最早的侵入体明显是位于内华达山脉最西端发现的环状杂岩体,年龄为125~120 Ma

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#### 图 9.4 科迪勒拉造山带的整体演化过程图

在图中,其他学者(DeCelles et al., 2009)试图解释科迪勒拉型岩基的起源以及白垩纪—古近纪—新近纪科迪勒拉造山带的整体演化过程。在弧后背景下的北美大陆,发生着如图所示的向东的俯冲作用,在这种条件下,北美克拉通数百公里的地壳发生了向西的逆冲推覆,从而形成了科迪勒拉造山带。在这个模型中,3~400 km宽的克拉通地壳消失在内华达山脉之下,克拉通地壳熔化形成内达华山脉岩浆和致密的熔融残余,然后熔融残余沉入地幔。这种模式具有质量平衡和空间问题,因为内达华山脉宽约100 km,这样一个约3~400 km宽、30~40 km厚的地壳剥离意味着足够的北美地壳向西迁移去加厚内达华山脉至120~150 km。熔化地壳和拆沉熔融残余并不能帮助解决问题,因为一般的大陆地壳包含有大约61%的二氧化硅,内达华山脉地壳上部的30 km含有约65%的二氧化硅,两者只有6%的差异,这意味着由于几乎等体积的克拉通地壳必须熔化去和内达华山脉的主体成分—起产生相同的质量组成,这将几乎不产生熔融残余。而且熔融作用不会去除地壳,但只会向上散发运移。这个模型还无法解释为什么岩浆作用沿整个山脉在80 Ma的时候终止

(Clemens-Knott and Saleeby, 1999)。其他类似年龄的岩体如 123 Ma 的沃德山(Ward Mountain)奥长 花岗岩和 121~105 Ma 的贝斯湖(Bass Lake)石英闪长岩出现在更北的约塞米蒂(Yosemite)国家公园正西侧(Lackey et al., 2012a, 2012b)。

一个较年轻的内华达山脉岩浆活动起始年龄可能与弗朗西斯科杂岩的开始年龄更相匹配。虽然基于混杂岩带内高级变质岩块的年龄,弗朗西斯科俯冲作用的开始时间通常被约束在 160±10 Ma 左右,但可能直到 130 Ma 之前不久,弗朗西斯科也没有形成,因为 130 Ma 是该杂岩内最老碎屑沉积岩的最大年龄——最东部和构造上高度连贯的蓝片岩含有 131 Ma 的碎屑锆石 (Dumitru et al., 2010)。高级变质岩块明显是多旋回的,因为岩块被阳起石和绿泥石包裹,很可能曾经是蛇纹岩,但仍然漂浮在陆源碎屑混杂岩中 (Coleman and Lanphere, 1971; Cloos, 1986)。因此,这个岩块相对于弗朗西斯科混杂岩完全是外来的,而弗朗西斯科开始俯冲的时间比一般认为的要晚得多。因此,如果能够证明弗朗西斯科混杂岩在 125 Ma 左右就在内华达山脉附近 (它们两个当时所处的位置还不是很清楚,Jayko and Blake, 1993),那么向东的俯冲可能就只是在 130 Ma 之前才开始。

这个假说主要的问题是,在100 Ma 的碰撞之前,弗朗西斯科可能不与内达华山脉相邻,这个碰 撞使其西半部和东半部连在了一起 (Jayko and Blake, 1993)。考虑到内华达山脉西部 125~100 Ma 的 深成岩体由西向东变年轻 (Lackey et al., 2012a, 2012b),很有可能在碰撞之前没有向东的俯冲,而 可能是向西俯冲至西部区域之下,在这种情况下,俯冲和板片后撤可能出现在西部和东部地块之间的 盆地中,且由西向东表现出年龄逐渐年轻的趋势。这个假说受到两个时期岩石发生不同程度的变形支 持,即大于 100 Ma 的内华达山脉内岩石发生了强烈变形 (Bateman et al., 1983a; Wood, 1997),而 125~100 Ma 大谷地群内岩石只发生了弱得多的变形 (Constenius et al., 2000)。

## 9.14 岩基内中白垩世转换挤压变形

正如前面章节所描述,关于海岸深成杂岩、内华达岩基、爱达荷岩基以及半岛山脉岩基的岩石在 约 100 Ma 发生变形时,与之相对应的事件我们知之甚少。大型断层、处处可见的走滑断层和其它地 方的逆冲断层等明显主导着构造样式,并把岩基分为两个部分,早有公认横穿接触带有不同的基底, 并展示出年龄、地球化学特征和同位素等方面的变化(图 9.5)。



图 9.5 科迪勒拉岩基岩和 110~100 Ma 转换挤压事件之间的关系 简图展示了文中提及的各种科迪勒拉岩基岩地体和 110~100 Ma 压扭事件之间的关系。 SN—内华达山脉

在加利福尼亚州南部和下加利福尼亚的半岛山脉(Peninsular Ranges) 岩基内,阿尔图斯(Alisitos)弧西部(可能在年轻地壳上喷发并穿过)与更东边的大陆弧地体之间著名的边界(Gastil et al., 1975,1990)形成于约114 Ma(约114 Ma是阿尔图斯地体中变形深成岩的年龄)和约98 Ma(约98 Ma是 La Posta深成岩最老的构造期后年龄)之间(Johnson et al., 1999a; Kimbrough et al., 2001)。 变形成因的模型包括在向东倾俯冲带之上的白垩纪碰撞(Gastil et al., 1981; Wetmore et al., 2003)、 东倾俯冲带之上的边缘盆地塌陷(Busby et al., 1998),以及向西倾的俯冲带之上的碰撞(Dickinson and Lawton, 2001)。根据白垩纪变质火山岩和正片麻岩刚好出现在主殉道(MainMartir)断层东侧, Johnson等(1999)认为在这两个弧地体之间有一次碰撞,但这些岩石的变质和变形程度都很高,以 致于它们都成为阿尔图斯地块的一部分。

鉴于阿尔图斯弧产出在碰撞带西侧,而且它的变质程度较东部岩石要低很多,可以合理地假设碰

撞的极性是向西倾的,卡沃尔卡-科尔特斯(Cortes)地体组成的东部带的前锋带部分俯冲在弧地体之下,缝合带截面上变质程度的差别支持着这个基本概念。向东倾的主殉道(Main Martir)逆冲断层可能是一个背冲断层,因阿尔图斯弧向上拱起而形成,并压在东部地块之上,而更东部的向西倾冲断了地块之间的主缝合带。

如前所述,后碰撞的 La Posta 深成岩代表一个从 99→92 Ma 的短暂岩浆脉冲 (Silver and Chappell, 1988; Walawender et al., 1990; Kimbrough et al., 2001),在 5~20 km 的深度侵入到高绿片岩相到角闪岩相的围岩,在很多地方发生了混合岩化 (Gastil et al., 1975; Todd et al., 1988, 2003; Grove, 1993; Rothstein, 1997, 2003)。这个岩体在剥露期间就位,即地下 10 km 深处的岩石通过拆离断层作用和崩塌作用被带至地表,这一剥露作用在时间上与盆地内一次脉动式粗碎屑沉积作用一致,该盆地位于岩体的西侧,沉积的时间为早森诺曼期到晚白垩世土仑期,含有 100~90 Ma 的碎屑 锆石 (George and Dokka, 1994; Lovera et al., 1999; Kimbrough et al., 2001)。虽然普遍接受的观点是 La Posta 侵入体与俯冲作用有关,但是岩浆作用与紧随碰撞作用之后的剥露作用一致,因此我认为存在这种可能,即 La Posta 深成岩由板片断离岩浆作用所形成,与阿尔图斯弧之下向西俯冲的板片破裂有关。一旦离开大洋岩石圈的束缚,东部板块由于浮力将快速上升,然后软流圈熔体上升至上覆地壳,形成 La Posta 侵入体。

在海岸深成杂岩体内,在可能的左行转换挤压时期之后 (Monger et al., 1994; Chardon et al., 1999; Chardon, 2003; Hampton et al., 2007; Gehrels et al., 2009), 一个广泛的变形事件影响了杂岩 体岩石和 Gravina-Dezadeash-Nutzotin-Gambier 带的岩石。这两个地块已有 U-Pb 年龄很好地限定,东 部为 190~110 Ma、西部为 160~100 Ma、见图 5.14。一个形成于 100~90 Ma 的向西逆冲断层带 (Haeussler, 1992; Rubin et al., 1990) 将东带的高级变质岩逆冲到 Gravina-Dezadeash-Nutzotin 带的 低级变质岩和可能的等同物之上,如更南边的 Gambier 群 (图 7.3),从而形成一个向上变质程度更 高的冲断堆积(Lynch, 1992, 1995; Journeay and Friedman, 1993; Crawford et al., 2000; McClelland and Mattinson, 2000)。同样,阿拉斯加南部卡希尔特纳浊流沉积盆地的岩石,位于兰格利亚与卢比 亚的更内侧部分之间,其在向南汇聚的构造作用下在 115~110 Ma 之间开始变形 (现在的坐标系) (Pavlis, 1982; Wallace et al., 1989; Ridgway 等人, 2002; Trop and Ridgway, 2007; Hampton et al., 2007),在这明显代表了一个由南到北穿时的盆地闭合(Kalbas et al., 2007)。Stikinia 和兰格利亚相 似的三叠纪古地磁极表明,至少根据那个时期的纬度它们相隔并不远(Kent and Irving,2010),可能 间隔两者的盆地还不是很宽,虽然目前约束还很少。半岛山脉岩基和海岸山深成杂岩体在变形时间上 的相似表明,这两者以前可能是相连的;但只有盆地相岩石的碎屑锆石年龄谱才能显示出不同的 (或者相近的) 源区,因为它们多少有些不同(Kapp and Gehrels, 1998; Manuszak et al., 2005; Manuszak et al., 2007; Hampton et al., 2010; Alsieben et al., 2011)

爱达荷西部剪切带内或附近的岩石,沿着爱达荷岩基的亚特兰大舌形体(Atlanta lobe)西缘分布(图 2.5 和 4.1),这些岩石遭受了右旋剪切强烈变形,时间在 105 Ma 之后(105 Ma 为小古斯克里克(Little Goose Creek)杂岩体中正片麻岩的岩浆岩年龄)和 90 Ma 之前(90 Ma 为横切较老组构的淡色伟晶岩的年龄)(Manduca et al., 1993; Giorgis et al., 2008)。这个变形带很可能是海岸深成杂岩体内部变形的延续,但在 90 Ma 之后和 70 Ma 之前,沿奥罗菲诺剪切带以左旋方式将爱达荷岩石水平错断(McClelland and Oldow, 2007)。

如前所述,根据地球化学、磁化率、年龄、放射性和稳定同位素、围岩来源以及基底类型,内华达山脉也可以分为较老的西半部和较年轻的东半部两部分(Nokleberg, 1983; ChenandTilton, 1991; Batemanetal, 1991; Kistler, 1990, 1993; Saleeby et al., 2008; Lackey et al., 2008, 2012a, 2012b; Chapman et al., 2012)。内华达岩基的岩石有约 100 Ma 变形事件的证据,这一变形事件晚于众所周知的内华达岩基内沉积岩和火山岩围岩的年龄(Bateman et al., 1983a; Bateman, 1992; Memeti et al., 2010a; Wood, 1997; Saleeby et al., 1990),然而,早于或者部分早于同时在 98~85 Ma 的内华达山脊岩浆作用事件的深成岩体(Greene and Schweickert, 1995; Coleman and Glazner, 1998; Davis et al.,
2012)。深成岩体的侵位与糜棱状剪切带的形成、90~87 Ma 冷却速率快速增大、以及大谷地群中土伦阶的沉积物增加(Mansfield, 1979; Renne et al., 1993; Tobisch et al., 1995)在时间上同步。

在内达华岩基南部,105→102 Ma 变火山岩、变沉积岩和深成岩岩石发生等斜褶皱(图9.6)穿 透性变形并向西逆冲(Wood,1997; Saleeby et al.,2008),年龄为约115~100 Ma 的正片麻岩和深成 岩体被剥露出来,从98 Ma 所处的9→10 kb 压力深度,抬升至95 Ma 时所处的4 kb 压力深度(Pickett and Saleeby, 1993; Wood and Saleeby, 1998; Saleeby et al.,2007)。Christe 最近的工作(2011)证 实,在内华达岩基北部,位于汝拉山地块内 Trail 组的127 Ma 流纹质凝灰岩,被逆掩倒转在129~127 Ma 的泰勒斯维尔(Taylorsville)冲断层之下,这次变形的最小年龄还没有得到有效约束。同样地, 沿内华达山脉西边的弗雷斯诺(Fresno)正北(图5.7),121~105 Ma 的贝斯湖英云闪长岩的褶皱 (Bateman et al.,1983a; Lackey et al.,2012a,2012b)是等斜的,并局部发生倒转。另外两个褶皱时 期重新褶皱了更老的等倾线(图9.7)。在内华达山脉中部,戈达德顶垂体(图5.7)的岩石也在这 个时期发生变形(Tobisch et al.,2000)。内华达山脉105~100 Ma 的强烈变形、并伴随着向西汇聚的 平卧褶皱和逆冲断层,而大谷地群内的岩石缺乏变形,两者的状况是不协调的,上述差异支持了 Wright和Wyld(2007)关于大谷地群是远距离迁移而来的模型,或者直到变形结束所有大谷地群的 岩石才开始沉积,支持这一观点的证据有:大谷地群内多数沉积岩中的碎屑锆石年龄以小于100 Ma 的为主(Surpless et al.,2006)。

Memeti 等 (2010a) 对位于侏罗纪弧岩石之上的几个顶垂体中变形的变质沉积岩石进行了碎屑锆 石测年,获得 U-Pb 年龄为 100±4 Ma。由于变质沉积岩被 101~95 Ma 的深成岩体穿切 (Memeti, 2010a),这个单元定年非常精确。这些岩石是一套分布更广泛的岩套的一部分,如前文所述,已知 变形的阿尔布阶火山岩和沉积岩来自遍及内华达山脉中部的顶垂体。这些岩石比 Gravina-Dezadeash-Nutzotin 盆地岩石略微年轻,同样位于中侏罗世弧岩石之上,而且在沉积之后不久就遭遇变形;因此 内华达山脉的例子可能代表了两个内华达地块之间的盆地相岩石。正如 Memeti 等 (2010a) 提出,雪



图 9.6 特哈查比山脉东部地质简图

显示了约 100 Ma 特哈查比(Tehachapi)侵入杂岩体(TIC)的岩席状深成岩体呈等倾伏卧褶皱,在构造上位于布莱克本峡谷(Blackburn Canyon)拆离断层之下,92 Ma 的深成岩体岩石逆冲到晚白垩世(?)Witnet 组之上



图 9.7 内达华岩基西部边界的地质简图

显示了沿内达华山岩基西部边界的地质概况,图示说明褶皱的几个世代,这些褶皱影响了 100 Ma 之前的侵入体和 它们的围岩。据 Bateman 等(1983a)修改,U-Pb 锆石年龄据 Lackey (2012a, 2012b)。注意早期的 F1 等倾线 湖断层在 145~102 Ma 期间或 87 Ma 的某个时期是活动的,活动时间取决于它的精确位置,现在位于 雪湖顶垂体之内的岩石可能来自死亡谷,它们在弧内发生转换挤压作用时期向北迁移,在两边的后续 盆地关闭之后被捕获。

这刚好略早于95~84 Ma 圆顶 (Domelands) 侵入岩套的深成岩体侵位 (Saleeby et al., 2008),该 侵入岩套是内华达山脊南部的对等岩套 (Coleman and Glazner, 1998)。另外,在岩基更远的南部,显示有有限的走滑运动,约在100~90 Ma 之间在原克姆峡谷 (proto-Kern Canyon) 剪切带上发生了约 25 km 的缩短,而在 90~80 Ma 之间,沿断层发生了另外 10 km 的右旋滑移 (Nadin and Saleeby, 2008)。

海岸深成杂岩体、内华达岩基以及半岛山脉岩基之间的相似关系,表明东部后构造(100~85 Ma)深成岩体可能都是板块断离的结果,而不是一般认为的板块俯冲。岩基的东带有约50 km 宽,并在盆地闭合和碰撞之后紧跟着发生短暂的岩浆脉动。在内华达山脉,东部成带(?)的深成杂岩,如图奥勒米和惠特尼山侵入岩系列侵位时间很短,但很强烈,这次爆发的岩浆活动被称为内华达山脊岩浆活动事件,发生在98~86 Ma之间(Coleman and Glazner, 1998; Davis et al., 2012)。在下加利福尼亚,复合带状的 La Posta 深成岩岩套在 98~92 Ma 期间被侵入(Kimbrough et al., 2001)。在上述两个例子中,在快速剥露过程中或紧随其后深成岩体明显就已经就位,并与西边的厚层森诺曼阶—土伦阶碎屑沉积序列同时发生(Mansfield, 1979; Surpless et al., 2006; Kimbrough et al., 2001)。

因为所有已定年的盖层(表壳岩系)岩石都早于碰撞事件并且发生了变形,对于后碰撞的内华 达山脉东部深成岩体的俯冲起源问题之一,就是几乎没有或很少有火山碎屑岩,如火山灰流、碎屑流 和富凝灰质的表生碎屑岩,作为特色产出于邻近岩浆弧。如在喀斯喀特(Fiske et al., 1963; Smith, 1985, 1991),尽管存在几种可能性——岩基是后来被从相邻地区构造运移来的,或者火山岩已被剥 蚀掉——板块断离模型可能代表着对火山碎屑缺失的一个简单的解释,因为岩浆房的岩浆没有被排放 到地表,因此那里没有火山。

约在 100~95 Ma 期间,内华达山脉南部岩石剥露速率从 9→4 kbs (Saleeby et al., 2007),伴随沿 主要断层的走滑分量,所有这些地区都明显有持续的岩浆作用(图 9.4)。正如长期以来被公认 (Fitch, 1972; Jarrard, 1986; Oldow et al., 1989; McCaffrey, 1992, 2009),这样的断层是斜向汇聚 板块机制下走滑运动分量轨迹,以苏门答腊西部的 Semangko (Barisan Mountain-巴里桑山脉)断层为 例,在那里它们跟随岛弧前缘 (Van Bemmelen, 1949; Westerveld, 1953),并将应力分配在不同的岩 石圈地块上。在北美西部的例子中,缝合带两侧的基底不同,与西部弧的变形和关闭相结合表明发生 了碰撞,但在不同情况下都有有力证据表明至少最后的相对运动是陡倾斜的。总的来说,这次约 100 Ma 事件出现在卢比亚超级地体西缘的一部分与北美发生对接的塞维尔事件之后,这次事件范围比通 常认为的更为广泛。

#### 9.15 内华达岩基褶皱

五十年前,Bateman 和 Wahrhaftig (1966) 推测,内华达山脉岩基的围岩(它们的"框架"岩石)形成了一个大的向斜,并推测(在那个板块构造理论出现之前的时代)地壳的向下挠曲导致轴向带的熔融和生成的岩浆上涌从而形成岩基。随着板块构造学的出现,大多数学者(如果不是全部的话)仍将岩基的形成与俯冲过程相关联,并认为区域变形发生在晚侏罗世,而且是由更西侧的碰撞事件引起的(Nokleberg and Kistler, 1980; Schweickert, 1981; Tobisch et al., 2000),或者是由底辟上升和随后发生的围岩向下沉所形成的(Moore, 1963; Hamilton and Myers, 1967; Saleeby et al., 1978a; Bateman, 1992; Pitcher, 1993; Saleeby and Busby – Spera, 1993; Saleeby et al., 1990; Paterson et al., 1996; Saleeby, 1999; Paterson and Fans, 2008)。尽管整个岩基的中白垩世变火山岩和变沉积岩已发生褶皱(Saleeby and Busby – Spera, 1986; Saleeby et al., 1990; Fiske 和 Tobisch, 1994),事实上,整个岩基年龄范围低至 96 Ma 的岩石(Memeti et al., 2010a)都发生了区域性褶皱,轴向为 NNW(北北西);另一组褶皱的走向差不多是正东的(Peck, 1980; Nokleberg and Kistler, 102

1980; Nokleberg, 1981, 1983; Wood, 1997; Saleeby et al., 1990, 2008), 因此所有 96 Ma 以前的深 成岩体必定也发生褶皱了。

内华达山脉最明显的问题是:有多少比 96 Ma 更晚发生的褶皱作用?褶皱发生有几期?这些都是 很难确切回答的问题,因为没有已知的小于 96 Ma 的表壳岩,所以由于它们的缺失,古水平面就未 知。然而,由于白垩纪侵入深成岩体在二维形状明显是相同的(不考虑年龄),这貌似可信的,即它 们所有的都发生褶皱了。

在整个内华达山脉岩基中,广义上的单一地层单元沿着侵入接触面分布达数千米,暗示着顶板和 底板最初是水平或者缓倾的。深成岩体接触带与框架岩石的层理普遍一致,即使是较年轻的深成岩体 也是如此,例如北穹隆 (North Dome)的花岗岩体 (图 9.8),如果发生褶皱后再侵入,这种接触带 将难以再现。这获得大量单个深成岩体之间围岩的协调地质景观支持 (Bartley et al., 2002; Grasse et al., 2001b; Bartley et al., 2012)。重要的是要记住这一点,虽然岩体的顶板和底板可能与围岩层理一 致,但是大而显著地脱离通常的岩席状,也许在外侧边缘,或者岩体上升到不同地层水平的某段有可 能提供分歧广泛和不协调的地图模式。通过这种方式,原来陡峭的部分甚至可能被更年轻的褶皱倒 转,并且在缺少含有可以沿着接触面顶部测定特征的围岩时,原始的三维形状可能是不清晰的和难以 解析的。甚至更有趣且难以解决的例子是:这些接触面在倒转的伏卧褶皱的翼部发生了旋转——一个 真实的可能性是,已知褶皱翼部的倒转发生在内华达山脉变形的白垩纪围岩中 (Peck, 1980; Wood, 1997)。

尽管没有大量的围岩,难以圈定出褶皱的轮廓,仍然有可能确定整个内华达山脉的许多发生了褶皱的深成岩体。例如熊穹窿(Bear Dome)(图9.9)的侏罗纪花岗岩(Jurassic Granite)很明显地被褶皱,因为面理和接触面向外倾斜,围岩环绕着它。邻近的94~91 Ma 拉马克花岗闪长岩内面理所限定的褶皱如图9.9所示。对图中向斜走向的东南部分进行了详细研究(Davis et al., 2012),揭示出该侵入体单元以及接触面和面理明确清晰地定义了一个向斜,如图9.10所示。在这个地区,可以核实它是一个由黑云母团块和板状包体确定的面理,并且被褶皱了,原始面理大致是水平的,平行于原来的席状侵入体的外部接触面,推测这些单元沿着走向向西北继续延伸。穿过褶皱,即走向大约向东的背斜,以几乎没有围岩的线性条带为特征。

在区域图中(Bateman, 1992),这样的线性条带密度交替出现在围岩中更为普遍。围岩密度更大的带很可能代表了向斜,那些区域深成岩的顶部岩石已广泛出露,而那些几乎没有带的围岩很可能代表了背斜,出露在深成岩体的核部区域。

对褶皱模型的另一个支持来自对内华达山脉深成岩体内面理的详细研究。内华达山脉深成岩体的 很多面理长期以来都被认为是岩浆成因的(Bateman et al., 1963; Bateman and Wahrhaftig, 196; Bateman, 1992),但最近的认识表明,很多深成岩体以显著的复式面理为特征,它们交切陡峭的内部和 外部接触带(Coleman et al., 2005; Zak and Paterson, 2005, 2009, Zak et al., 2007; McNulty et al., 1996; de Saint Blanquat and Tikoff, 1997; Tikoff and de SainBlanquat, 1997),这就使得大部分面理是 在侵位期间或者侵位后不久由黏性流动形成的岩浆特征这一假设很难被接受。

Zak 等(2007)的详细研究结果表明,图奥勒米侵入系列岩石中有四种不同的面理,并认为2种 较年轻的面理(N-S和NW-SE)明显横切了内接触面,它们和侵入后的应力有关(图9.11)。最 近,其他学者的详细研究已经证实了在图奥勒米侵入系列的深成岩体内的东向延伸面理(Economos et al., 2005; Johnson and Miller, 2009年),它可能是向东延伸的褶皱轴向平面,这些面理也可能提 供一个可靠的时间约束,因为图奥勒米侵入系列的从向北西到向西面理(第4类,Zak et al., 2007) 影响了所有地质单元,包括最年轻的和最中心的单元——约翰逊(Johnson)斑岩,而北到北西向面 理(第3类,Zak et al., 2007)影响了除约翰逊花岗斑岩之外的其它所有单元。对这一观察结果需要 提醒的是,在露头上的他们可以观察到两个面理的相对时序,北西向的组构(他们的第4类)比更 向北的组构(他们的第3类)要晚。这一矛盾最简单的解释是只有向北的组构在该区域内没有被识 别出来,或者可能是约翰逊斑岩在变形事件中没有完全固结,以至于没能留下那一期的组构的记录,



图 9.8 内达华山脉深成岩体和围岩之间关系地质简图

显示了在加利福尼亚州国王峡谷(Kings Canyon)地区内华达山脉的深成岩体和围岩的关系。注意深成岩体从北穹隆(North Dome)花岗闪长岩向北深成岩体明显堆叠,沿途深成岩体明显环绕着它。我认为这些关系,深成岩体和 它们的围岩之间一起的整体性质,说明所有这些岩体的顶板和底板是一起褶皱的。地质资料据 Moore (1978), Moore and Nokle berg (1992)和 Grasse (2001)



图 9.9 戈达德顶垂体 (pendant) 北端地质概况示意图 显示了戈达德顶垂体 (pendant) 北端的地质概况,同时还显示了被褶皱的熊穹隆 (Bear Dome) 侏罗纪花岗岩和白垩纪深成 岩体内的其他褶皱。地质资料据 Bateman (1965a, 1965b), Bateman and Moore (1965), Lockwood and Lydon (1975)。

但是记录了第二次变形事件。如果这是正确的,那么北-南组构是形成于约翰逊斑岩冷却过程中,锆石 U-Pb 年龄为 85~84 Ma (Coleman, 2005),而横向的面理必定更年轻。

对 98 Ma 的 Jackass 湖深成岩体进行详细研究 (Stern et al., 1981; McNulty et al., 1996), 该岩体 位于图奥勒米侵入岩系列的正南方, Krueger (2005) 和 Pignotta 等 (2010) 发现已发生褶皱的深成 岩墙, 轴面平行于围岩和包体中的变质面理轴。Pignotta 等 (2010) 也证实深成岩体中主要为北西向 面理和陡倾线理, 与变火山岩围岩中的基本相同, 说明它们形成于岩体侵位之后。

内华达山脉向斜的东部,在白印优山脉(图 2.5)内为一系列大褶皱,包括白山-因约郡复背斜(Ross,1967; Morgan and Law,1998),而且,尽管那里构造错综复杂,例如不有同时代的断层和褶皱,但两组主要褶皱(图 9.12)形成了一个盆地和穹窿型样式(Ramsay,1967; Thiessen and Means, 1980)。在这个地区,细致而详尽的野外研究已显示有几期的褶皱,包括垂直和翻转的 NW-SE 和 NE-SW 走向褶皱,规模从山岭大小到厘米级(Nelson,1966a,1966b,1971; Bateman,1965a; Ross, 1965; Morgan and Law,1998)。该区域的沉积岩被许多中侏罗世和晚白垩世的深成岩体所切割(Krauskopf,1968; Crowder et al. 1973),但褶皱作用与深成岩作用的关系仍存在争议,一些学者认为 侏罗纪深成岩体晚于褶皱(Morgan and Law, 1998),另外一些学者认为晚白垩世深成岩体晚于褶皱,并在侵入期间膨胀,向旁边挤压或者把它们的围岩顶起(Nelson and Sylvester, 1971; Sylvester et al., 1978b; Morgan et al., 1998, 2000; de Saint Blanquat et al., 2001),还有人认为侵位之后白垩纪深成



# 图 9.10 拉马克 (Lamarck) 侵入杂岩体的详细剖面 (据 Davis et al. (2012))

地质简图展示了一个横穿拉马克(Lamarck)侵入杂岩体的详细剖面,说明了在 Dusy 盆地中岩体的褶皱性质。注意横穿单 元的对称性被认为是褶皱轴。这是一个最好的区域,可看到解理不是轴向平面的,而是一个整体的、初始水平的组构。这 个区域刚好位于上图(图9.8)的东南角,是在拉马克的突出向斜的延续



图 9.11 简图展示图奥勒米 (Tuolumne) 侵入系列中两组不同的横切解理 (据 Zak 等 (2007) 修改) 这些解理显然是岩浆期后的,在这里被认为是在两期褶皱形成过程中产生的, 这两者都比深成岩体的侵位年龄年轻。

岩体被变形,但不一定是褶皱作用导致(Paterson et al., 1991)。此外, Glazner 和 Miller(1997)提 出某些侏罗纪深成岩体形成了具有环边缘背斜的巨大载荷。Vines 和 Law (2000) 认为 164 Ma 的圣丽 塔平台(Santa Rita Flat)深成岩体最初是由沿着向斜枢纽带打开的一个空腔口,随着深成岩顶板向 上拱隆起而扩展增大。Stein 和 Paterson (1996)认为侏罗纪的复式环状深成岩体在该区域的就位是通 过围岩的向下位移和侧向韧性流动造成的。图 9.12显示了部分白印优山脉的多期变形,其两组褶皱 一起构成一个复杂的类型1干涉图案。几个褶皱伏卧翻转(Ross, 1967)。请注意, 大多数但不是全 部的侏罗纪深成岩体具有协调一致的接触关系。也就是说,只有一个或者两个地层层位沿着深成岩体 接触面延伸数千米。几个侵入体在盆地中就位,如克里克溪-约书亚平台(Beer Creek-Joshua Flat) 和大理岩峡谷(Marble Canyon)岩体,类似于在其他地方见到的,沉积单元通过对较老的褶皱重新 褶皱而形成。考虑到该地区的褶皱程度,很难想象这些深成岩体能够利用这样褶皱之后的地体,形成 很长的协调一致接触关系。另外,通过对约书亚平台岩体(图 9.12)东北端的仔细地调查,清晰地 显示了被包裹着的褶皱鼻。大理岩峡谷岩体由很多向内倾的板状岩体组成(图 9.12),因此被合理地 解释为一个被再次褶皱的褶皱。83 Ma 的 Papoose 平台深成岩体(图 9.12)占据着背斜的核, Sylvester 等(1978b)认为这个背斜代表着一个膨胀的气泡。然而,对东端仔细地调查显示有两个 E-SE 向伸展构造,背斜核部被一个向斜分开,表明这个岩体也被褶皱过。这与 Paterson 等(1991)和 Morgan 等(1998)的详细研究结果是一致的, Paterson 等认为该区在白垩纪深成岩体侵入过程中或者 之后曾经历了显著的区域变形, Morgan 等(1998)利用顶板岩石中先存的斑岩变晶包裹体行迹, 推 断深成岩体最初是一个一致的席状岩体。此外,年龄为白垩纪(Nelson, 1971)的石英闪长岩岩床



图 9.12 加利福尼亚州白印优山脉中央的地质简图

揭示了两期褶皱形成的干涉模式。注意侏罗纪和白垩纪深成岩体明显都被褶皱,因为岩体接触几乎在所有地方都与围岩岩石相整 合,这是非常困难的,但也不是不可能发生,如果深成岩体侵入到已经褶皱的岩石中的话。大理石峡谷 (Marble Canyon) 深成岩体 明显占据了一个构造盆地,表明它在两期褶皱之前侵位。Papoose Flat 侵入岩体的东部边缘已经被解释为是褶皱化的,因为它的接 触面几乎沿着褶皱的层理。地图根据 Nelson (1966a, 1966b, 1971)的原始填图;年龄据 Miller (1996), Chen (1977)和 Coleman

等(2003)。BC-桦树溪(Birch Creek)深成岩体; RC-雷丁峡谷(Redding Canyon)深成岩体(Coleman et al., 2003)

(图 9.21) 被认为明显发生过褶皱,并与它们的围岩一起被错断。

所有这些观察结果说明了这种可能,尽管该区白垩纪深成岩体都被褶皱过,但它们的顶板和底板 现在和岩基及岩基之间的隔板在走向上是一致的。事实上,根据其它地方的观察(Hildebrand et al., 2010b),深成岩体及与围岩的接触通常可能是水平的,这取决于侵位的深度,所以现在多数深成岩 体和围岩的接触代表着板状深成岩体倾斜的顶板和底板(Hamilton, 1988a)。根据前文所述的图形样 式和面理特征,以及刚好位于白印优山脉以东地区之间的关系,很可能100 Ma之后内华达山脊侵入 体也被褶皱过。

内华达山脉和白印优山脉深成岩体褶皱模型之所以值得相信,是因为有记录良好的和双褶皱的斯 图尔特山(Mount Stuart)复式岩基,该岩基位于华盛顿州的高喀斯喀特山脉之内,含有年龄少于91 Ma 的深成岩体(Matzel, 2006; Paterson and Miller, 1998)。类似地,在不列颠哥伦比亚省的海岸深 成杂岩中,年龄少于84 Ma(图 7.3 和图 7.4)的深成岩体明显被褶皱了,同时发生褶皱的还有年龄 范围从 104→94 Ma 的较老岩体(Brown et al., 2000; Brown and McClelland, 2000)。

内华达山脉褶皱假说的意义是:沿东部边缘出露的许多深成岩岩和原来认为是复合岩墙充填的岩体,如 McDoogle (Mahan et al., 2003)和公驴湖 (Jackass) (McNulty et al., 1996),可能是作为岩床 杂岩体就位的,而不是岩墙。假说也暗示,实际上大的岩墙群可能曾经是以岩床的形式侵位,而不是 以前认为的岩墙 (Moore and Hopson, 1961; Moore, 1963)。该假设获得三分峰地区的野外填图关系 的支持 (Moore and Sisson, 1987; Moore, 1981),那里独立岩脉明显是与侵入岩体和围岩框架一起发 生了褶皱 (图 9.13),



图 9.13 美国加州红杉国家公园内华达山脉岩基地质示意图 该图说明被褶皱岩席状深成岩体的堆叠特征。注意三分峰(Triple Divide Peak)区域相对褶皱翼的岩床走向的变化。 该图根据 Moore 和 Sisson (1987)以及 Sisson 和 Moore (1994)地质填图绘制

通常而言,在整个内华达山脉岩基中,只有少量不整合在更老的中生代和古生代岩石上的白垩纪 盖层得以保留,这使我想到大多数深成岩体的顶板靠近火山盖层与其侏罗纪基底的接触面。在随后的 隆起期间,抗风化能力较弱的火山岩被剥蚀掉,保留了抗风化剥蚀能力较强的深成岩体及其变质围 岩。里特山脉(Ritter Range)顶垂体的 Minarets 火山口能够得以保存,是因为其核心沿着它的环形破 碎断层下降了(Fiske and Tobisch, 1994)。

如果深成岩体发生褶皱,那么变形主要阶段的年龄必定比这个区域最年轻的被褶皱岩石更年轻, 在白印优山脉中是 83 Ma。如前面所讨论的,如 95~84 Ma 的图奥勒米侵入系列的杂岩体中两个最年 轻的面理切过接触面,暗示变形在侵位之后发生(Bateman et al., 1983b;Žák and Patterson, 2005; Žák et al., 2007),这与来自更年轻的深成岩体(如那些与 92 Ma 的拉马克花岗闪长岩有关侵入体) 已变形的岩墙一样(Coleman et al., 2005)。如果内华达山脊岩浆作用事件的深成岩体(内华达山脉 岩基最年轻的岩浆作用)发生了褶皱,那么变形将明显比约 85~83 Ma 更年轻,时间上很可能是属于 拉勒米造山期。

来自特哈查比山脉(Tehachapi Mt.)东部的内华达岩基南部资料支持这一结论。那里,可能是 坎帕阶—麦斯里希特阶的Witnet 组沉积岩,不整合地覆盖在92 Ma的内华达花岗闪长岩之上,在中新 世之前,沿着携带了内华达山脉花岗质岩石的逆冲断层向北汇聚,该组岩石被逆掩上冲和褶皱,局部 甚至倒转(图 9.6)(Lechler and Niemi, 2011; Wood, 1997)。这表明那些褶皱和冲断层是在晚白垩 世拉勒米造山运动期间形成的,内华达山脉的横褶皱也是在那个时期形成的。

大多数情况下,内华达山脉中年龄从100→83 Ma的深成岩体岩石的古地磁极形成了一个紧密的阵列,如果从表面来看似乎会排除侵位期后的褶皱作用 (Frei et al., 1984; Frei, 1986; Hillhouse and Groomé, 2011),但是由于100 Ma的火山岩明显被褶皱,我怀疑古地磁结果反映了较年轻的重置和再平衡,因为持续的岩浆侵入升高的热流一直持续到80 Ma (Dumitru, 1990)。如果古地磁极在大约80 Ma 固定,它不会记录80 Ma 之后内华达山脉板块相对于大盆地地区的大规模纬度运动,但会有之前的。

#### 9.16 加拿大分区双科迪勒拉岩基

人们早已认识到(Monger et al., 1972, 1982),加拿大分区内有两条平行的科迪勒拉型岩基带: 位于加拿大东部的奥米尼卡带和沿太平洋分布的海岸深成杂岩带(图 2.5 和图 5.12)。Monger 等 (1982)认为这两条带形成与海岛的(Insular)和山间的(Intermontane)这两个超级地体与北美的对 接或碰撞有关,在这里我坚持认为这些岩浆作用是俯冲作用所导致的,而且这两条带是大陆弧的一部 分。在普遍接受的科迪勒拉模式中,奥米尼卡带的深成岩体侵入北美地壳,导致向东的逆冲作用,但 是我认为它们是北美向西俯冲到卢比亚带状大陆之下所导致的。不管选择哪种模式,在加拿大科迪勒 拉内某处一定有一个大断裂,沿此断裂两条岩基带成双。但是,如我们所见,在距今125~80 Ma 的 科迪勒拉岩基岩浆活动之前,加拿大分区内几乎每一个地体都与附近的地块相连。所以,这个断裂唯 一可能存在的位置必然在这两个岩基地体之间,它沿卡什溪地体东部边界分布,该地体以白垩纪和古 近纪一新近纪的右旋走滑断层及斜向滑动断层的边为界(Wheeler and McFeely, 1991; Gabrielse, 1985; Struik et al., 2001)。该断层至少有 1500 km 的错断间距,是整个科迪勒拉山系中最大的断层之 一。它可能在阿普弟期(塞维尔碰撞)到始新世期间活动,但其时代还不太确定。

就在所提议的走滑断层西边的 Sustut 盆地, Irving 等(1995, 1996)最早根据古地磁学数据提出 了断层的假说,并称之为 Intra-Quesnellia 断层。Sustut 盆地是线型盆地,位于 Stikinia 地体和卡什溪 地体之上,平行于卡什克里克地体东部接触带,所提议的断层就位于该处(图 2.5)。该盆地地层有 2000 多米厚,是介于阿普第阶—阿尔必阶到坎帕阶之间的粗粒陆源碎屑岩(Evenchick and Thorkelson, 2005; Evenchick et al., 2007),看起来与科迪勒拉中部拼合格格不入,但从逻辑上可以 解释为沿断层弯折或斜面形成。类似地,出露在卡什溪地体北带的 89±2 Ma 的平顶山(Table Mountain)火山套和 85±5 Ma 的 Surprise Lake 岩基(Mihalynuk et al., 1992)也与所处之地不协调。

在科迪勒拉山系南部,横跨不列颠哥伦比亚省-华盛顿州边界区域的是几个不同时代、不同成因 的小地体,例如 Chilliwack、桥河(Bridge River)、Easton、Cadwallader-Tyaughton-Methow,还有更多 的(Monger and Struik, 2006)只是以构造岩片或超覆层序(MacLaurin et al., 2011)的形式存在,这 里没有对它们进行描述,因为对于我们要讲的主要内容来说,它们只是很次要的部分,但是,由于夹 在兰格利亚南部与 Stikinia 之间, 它们确实会对最西边的地体聚合产生影响。对这些岩片有各种各样 的解释。如它们可能代表了年龄变化范围为古生代到中生代的小块洋壳、开阔大洋的燧石、弧前盆地 和蓝片岩 (Monger et al., 1982, 1994; Umhoefer et al., 2002)。再往北, 对于兰格利亚-Stikinia/育空 塔塔纳河的接触带人们的看法不一,有的人认为其属于侏罗纪,有的人认为属于白垩纪,但是大部分 学者都认为 Gravina-Dezadeash-Nutzotin 带(代表两个地块之间发育的一条狭窄盆地) 在约 100 Ma 时 关闭 (Haeussler, 1992; Rubin et al., 1990; Journeay and Friedman, 1993; McClelland and Mattinson, 2000)。此外,古地磁数据(图 9.14)显示兰格利亚和 Stikinia 在纬度上从来就没有离得太远(Kent and Irving, 2010)。因此,如果这个盆地很窄,就很难理解各种大洋地块和弧前地块是如何在海岛的 和山间的超级地体之间合并。Monger 等人(1994)针对这个难题给出了一个令人信服的解释,他认 为有一个弧内断层左旋位移时期,该弧内断层将北部海岸深成杂岩向南搬运,并使南部海岸深成杂岩 的宽度加倍,这就解释了其宽度异常的问题。这次迁移超覆和捕获了之前在该弧更靠南部外侧的弧前 岩石。Gehrels 等人(2009)随后用这个模型解释成双的海岸深成杂岩和 Gravina-Nuzotzin 盆地在约 110~100 Ma 的关闭。无论如何解决这个问题,很显然的是,海岛的和山间的超级地体是在 100 Ma 时 连在了一起,完成了卢比亚带状大陆的拼合。





大约 125 Ma 塞维尔 (Sevier) 造山运动最初碰撞之后,相对于北美克拉通来说,卢比亚向南移动,接着,在 拉勒米运动之后,卢比亚又向北移动。

AS—二叠纪 Asitka 组。三叠纪—休罗纪单元: KM—Karmutsen lavas (约 225 Ma); SM—Savage Mountain (约 210 Ma); HZ—Hazleton 组 (约 195 Ma); BZ—博南扎 (Bonanza) 弧 (195 Ma)。白垩纪单元: SPR—Silverquick-Powell Creek; NM—温哥华岛 Nanaimo 组。新生代单元: FL—温哥华岛 Flores 火山岩; OL—中不列颠哥伦比亚 Ootsa 火山岩; SBC—不列颠哥伦比亚南部新近纪玄武岩

# 9.17 拉勒米事件:约80~75 Ma 变形与变质作用

在整个北美西部约 80~75 Ma 时发生的变形和变质作用,通常都被归为拉勒米事件。包括我本人 在内的几个科学家(Armstrong, 1968; Perry and Schmidt, 1988; Miller et al., 1992; Hildebrand, 2009) 将拉勒米构造或拉勒米造山运动这个术语限定在科罗拉多高原地区的厚皮构造带,但是那一时期的变 形作用从南美洲到阿拉斯加都有出现,所以将它们视为同一次运动明显是合理的。Drewes (1978, 1991)使用科迪勒拉山系造山带一词,是指南美洲和北美洲侏罗纪到始新世变形的整个带,但是他 的工作主要关注的是从加利福尼亚州到德克萨斯州晚白垩世—古近纪时期的造山带区域。Hildebrand (2009)使用科迪勒拉造山运动一词,是指从阿拉斯加到墨西哥南部整个造山带,甚至越过该造山带 的晚白垩世到古近纪时期的变形作用。没有哪个名字可以让每个人都觉得满意,这些历史上曾经用过 的术语显然有它们各自的优点,所以拉勒米和科迪勒拉山系这样的词应该被保留下来。这里,我使用 拉勒米事件一词指晚白垩世—古近纪的所有变形和变质作用。总体上,它包括从晚白垩世到始新世的 各种变形特征,包括莫哈韦沙漠-索诺拉地区的推覆体,大盆地分区中有基底卷入的厚皮逆冲断层、 覆盖在更广阔的线性前渊上的局部盆地的演化,加拿大分区、墨西哥东部、中美洲和南美洲北部的薄 皮冲断作用,以及内华达山脉区域的一些褶皱。以这种方式使用这一术语,可以清晰地将该事件和大盆 地区域早先发生的塞维尔事件区分开来,而且,因为一部分位于加拿大和索诺兰沙漠分区内的岩石明显 属于北美大陆的,它们基本没有受到较老变形作用的影响,所以不应该混淆。但是,由于局部一些地区 一直在进行,所有白垩纪—古近纪—新近纪的变形都可以包括在科迪勒拉造山运动中。

拉勒米变形作用的经典地区位于造山带的大盆地分区,北部以路易斯和克拉克线性构造为界,南部以菲尼克斯断层(图4.1)为界。在那里,主要变形是北美大陆克拉通基底被卷入的厚皮逆冲断层(Grose,1974; Smithson et al.,1979; Brewer et al.,1982; Rodgers,1987; Hamilton,1988b)。一般认为内华达-白印优山脉地块没有明显的拉勒米变形,但是根据前面提出的分析,我认为该区域有 85 Ma之后的褶皱,所以在拉勒米事件期间被变形了。在加拿大分区内的路易斯和克拉克线性构造北部,拉勒米变形主要是落基山脉褶皱-逆冲断层带的薄皮冲断层,而在菲尼克斯断层以南的地区,正如之前所描述的那样,变形的变化性很大,由莫哈韦沙漠-索诺拉西部地区有基底卷入的厚皮变形、玛利亚(Maria)褶皱-逆冲带中有基底卷入的逆冲断层和达到角闪岩相的变质作用,以及中-南墨西哥的薄皮变形等构成。在中美洲和南美洲北部地区,变形作用包括加勒比海大弧(Great Arc)之间明显的碰撞和增生,以及在其之前增生的洋底高原。

自从科尼和雷诺兹(Coney and Reynolds, 1977)总结并解读了加利福尼亚州和亚利桑那州南部的放射性测量数据后,他们认为岛弧岩浆作用在 80 Ma 席卷了 1000 km 的内陆地区,然后在大约 45 Ma 又再次出现,这是由于板片以缓角度俯冲,随后变陡。此后大部分模型认为拉勒米事件在一定程度上和北美克拉通之下的缓倾角俯冲有关(Coney, 1976; Dickinson and Snyder, 1978; Bird, 1988; Hamilton, 1988a; Dumitru et al., 1991; Saleeby, 2003; Grove et al., 2003b; Jacobson et al., 2011)。但是,北美西南的拉勒米变形和岩浆作用的走向更接近于东西向明显而不是北南向(图 4.1),也就是说Coney 和 Reynolds(1977)使用的数据是在平行或几乎是沿着这条带的走向上收集的,而不是穿过它(Glazner and Supplee, 1982)。因此,在距离-年龄图上时,数据点呈近水平状排列也就不足为奇了。

在汇聚边缘,一些学者将俯冲过程本身与上盘板块压缩变形联系起来(Monger and Nokleberg, 1996; Hutton, 1997; DeCelles, 2004; DeCelles and Coogan, 2006; DeCelles et al., 2009),尽管有相当的证据表明是以未压缩的弧前盆地和地区的形式(Von Huene, 1984; Loveless et al., 2005),几乎每个弧都有低水位、中性到张性的特征(Hamilton, 1981, 1988a, 2007; Hildebrand and Bowring, 1984; Hildebrand, 2009),以及板片回撤占主导的过程——当致密的大洋板块垂直沉入地幔时(Elsasser, 1971; Karig, 1971; Molnar 和 Atwater, 1978; Garfunkel et al., 1986; Hamilton, 1985, 2007; Schellart et al., 2006; Clark et al., 2008)——不能产生可观测到的变形或变形程度。例如,下沉的洋壳怎么能够传递足够的力量来移动和抬升贝尔特-珀塞尔这样—个体积超过 450 km ×200 km ×25 km、覆盖超过整

个被动边缘,并且超出前渊的白垩纪岩石的呢?

此外,正如其他一些作者(Maxson and Tikoff, 1996; English et al., 2003)指出的那样,这种上盘 变形需要地幔岩石圈的彻底融蚀以传输压缩压力,而新生代地幔捕虏体同位素研究表明其一直存在 (Farmer et al., 1989; Livaccari and Perry, 1993; Lee et al., 2000, 2001)。Saleeby (2003)认识到这个 问题,于是提出了一个分段模型,在这个模型中只有俯冲板块中的一条狭窄区域是缓的。但是,实际 的变形从南美洲一直延伸到阿拉斯加,比他的模型所预计的要宽得多。

再者,显示出 Gutscher 等(2000)认为的平板俯冲是非常关键的,假设拉勒米变形是由一个高 耸的洋底高原和北美大陆碰撞产生的(Livaccari et al., 1981; Henderson et al., 1984; Saleeby, 2003), 根据 Barth 和 Schneiderman (1996)的观点,最新模型包括了假设的结合物,与赫斯和沙莰基隆起 (Hess and Shatsky)的碰撞时代略有不同,赫斯和沙莰基隆起现在位于西北太平洋区域(Liu et al., 2010)。在最新模型中,高原大部分或完全俯冲,留下了很少或几乎没有留下,这就使得很难对模型 进行验证。但是现在,这类碰撞在其他一些地方仍然正在发生,所以有可能根据现在的实际观测来研 究高原碰撞的影响。

正在碰撞的洋底高原最壮观的例子发生在西南太平洋的所罗门群岛。在那里,所罗门弧 (Solomon arc)和厚达33km广阔的翁通-爪哇(Ontong-Java)高原——中生代形成位于太平洋的大 洋地壳上的一个玄武岩高原,至少被1000m的远洋沉积物所覆盖——在上新世的碰撞显然导致了沿 现今的森林云雀(Woodlark)盆地开始了一个新的向北东倾的俯冲带启动(Hamilton,1979及其参考 书目)。森林云雀盆地包含多个洋底高原和一个俯冲扩张脊。地球物理研究(Miura et al., 2004; Phinney et al., 2004)显示在碰撞过程中,只有高原上部靠前部分被拆离,形成了叠瓦状、并增生到 马莱塔岛(Malaita)增生杂岩内的所罗门弧前锋(Hughes, 2004),而高原的下部和主体部分继续俯 冲(Phinney et al., 2004; Mann and Taira, 2004; Taira et al., 2004)。铪钕铅同位素数据揭示古老太平 洋和翁通爪哇组分加入到了所罗门弧的弧岩浆中(Schuth et al., 2009),支持该模型。也许让那些与 高原俯冲相关的大规模、宽范围、上盘变形的支持者感到惊讶的是,对整个高原增生阶段处于活动状 态的中央所罗门弧内盆地的详细研究,只发现有限的褶皱和断层(Cowley et al., 2004)。只有板片回 撤才能解释这样厚实的高原俯冲。

Von Huene 和 Ranero (2009)回顾了沿美国西部发生的 6 次不同的碰撞所带来的影响,发现尽管 在阿拉斯加东南部,如 Yukutat 地块(洋底高原)(Perry et al., 2009; Worthington et al., 2012)那样 厚的地块到来亦可以在局部形成高山和强烈变形。但是在广阔的大洋内,与高耸的、有浮力的玄武质 地区碰撞相关的主要过程是①大洋板块完全俯冲,甚至在高地形处也是如此;②上盘局部和短期抬 升;③增生楔局部俯冲侵蚀。他们得出结论,高耸的大洋特征碰撞通常产生较小的板块间摩擦,形成 的上盘地势突起是暂时的,为了形成高的海岸山脉,某一地区必须有巨大突起体长时期的俯冲,就如 现在正在发生的椰子山脊(Cocos Ridge)向中美洲之下的俯冲一样。这一结论有效地排除了具有强 斜向汇聚的区域。

最近在跨墨西哥火山弧下的俯冲板块研究表明,俯冲板块向东逐渐变得更平坦,直到它看起来明 显直接位于格雷罗地体近 250 km 的内陆地壳下面,从而解释了来自中美洲海沟方向的岛弧岩浆作用 的巨大差异(Pardo and Sudrez, 1995; Pérez-Campos et al., 2008)。虽然正如全球定位系统探测到的 那样,俯冲板块和北美地壳之间显然存在着某种相互作用(Payero, 2008),但那里明显没有类似于 拉勒米事件的那种变形或特别厚的地壳。这种平板俯冲确实导致了宽广区域的第四纪岩浆作用 (Blatter et al., 2007),再结合对源自板片和俯冲水出现在软流圈楔及产生火山岩的观测(Jodicke et al., 2006; Johnson et al., 2009; Chen and Clayton, 2009),他们提出平板俯冲能够很好地将挥发分输送 到正常弧-海沟的内侧,且能产生大量的岛弧岩浆。类似地,Yukutat 洋底高原向阿拉斯加东南部的北 美洲之下的平板俯冲已经产生大量岩浆,形成了壮观的复式火山锥和盾状火山(Neal et al., 2001; Richter et al., 1995, 2006; Trop et al., 2012)。因此,即使大盆地(Great Basin)的平板模型是可行 的,好像并没有明显的理由可以说明为什么在拉勒米事件期间缺少俯冲相关的岩浆。

# 9.18 沿造山带的拉勒米碰撞

从逆冲作用和相关造山带前渊发育的时代推断,拉勒米事件几乎同时出现在超出整个造山带长度 上,尽管很难精确确定不同区域逆冲作用的时代、边缘形状轻微的不规则使得边界形态复杂的过程。 根据变形年龄的相似性,卢比亚和北美克拉通的碰撞肯定是几乎垂直的。这就强有力地说明那时卢比 亚带状大陆是一个整体。

在阿拉斯加,凡兰吟期-森诺曼期的卡希尔特纳盆地(图7.7)在约74 Ma时向北逆冲,与坎帕期—麦斯里希特期发育的坎特威尔(Cantwell)盆地一致,该盆地是一个逆冲过程中形成的冲断层顶部盆地(Ridgway et al., 2002; Trop and Ridgway, 2007)。晚白垩世—古近纪向北汇聚的逆冲和褶皱改变了阿拉斯加北部早白垩世的特征(Moore et al., 1997)。

在加拿大区段,在晚白垩世—古近纪期间,北美大陆台地的岩石沿位于寒武纪页岩内的一条拆离带与它们的基底岩石分离、褶皱并向东逆冲,形成落基山褶皱--逆冲断层带(图 2.5、图 7.1)(Price and Mountjoy, 1970; Price, 1981; Price and Fermor, 1985; Fermor and Moffat, 1992)。在这次变形中,在前陆盆地的东部形成了一个坎帕期-古新世的厚层碎屑楔(Larson et al., 2006; Ross et al., 2005)。

从塞维尔到拉勒米时期,大盆地分区的北美大陆边缘逆冲作用显然是持续进行的,但已减弱很多(例如,DeCelles and Coogan, 2006; Yonkee and Weill, 2011)。在85~75 Ma 期间,内陆带发生了强烈的变形和岩浆作用(Camilleri et al., 1997; McGrew et al., 2000),在麦斯里希特期(Maastrichtian),科罗拉多高原开始了经典的厚皮变形(Dickinson et al., 1988; Lawton, 2008)。

一个连续的晚白垩世—古新世前陆褶皱-逆冲断层带和相关的前渊就出现在墨西哥湾西部的整个 墨西哥中北部和东部地区(图 2.5)(Eguiliz de Antuñano et al., 2000)。它们形成于格雷罗超级地体 增生期间(Tardy et al., 1994; Centeno-Garcia et al., 2008, 2011),位于卢比亚的中南部。在晚白垩 世—古近纪期间,Oaxaquia 和密斯特克(Mixteca)的西部边缘发生了变形,该处位于一个总体上向 东汇聚的褶皱-逆冲断层带(Suter, 1984, 1987; Hennings, 1994; Fitz-Diaz et al., 2012)和在前展式 (背驮式)冲断层上发育的坦皮科-米桑特拉(Tampico-Misantla)前渊上(Busch和Gavela, 1978)。向 南,是 Zongolica 褶皱-逆冲断层带,它被卷入了桑托期—坎帕期(Santonian-Campanian)深水沉积岩, 向东逆冲覆盖在礁碳酸盐为主的科尔多瓦(Cordoba)地台之上(Nieto-Samaniego et al., 2006)。

在墨西哥南部的 Cuicateco 地体,麦斯里希特期(Maastrichtian)的片岩、绿岩、辉长岩和蛇纹岩在白垩纪末—古新世期间被向东逆冲覆盖在玛雅(Maya)地体的红层之上(Pérez-Gutiérrez et al., 2009)。在玛雅地体的南端,一个面朝西以碳酸盐岩为主的台地,位于玛雅(Maya)地体基底之上,并在坎帕期末被淹没,在麦斯里希特期—达宁期(Maastrichtian-Danian)期间被造山复理石掩埋,并被超基性推覆体逆掩。断层下盘结晶基底的岩石在 76 Ma 时变质为榴辉岩,说明北美边缘的这部分大约在那个时期俯冲了 60 km 以上的深度,并且在 1 Myr 后抬升剥露至角闪岩相,可能是由于板片断离(Martens et al., 2012)。

在已旋转的 Chorits 地块南面更远的地方, Rogers 等(2007)确认了一个向东南倾的晚白垩世叠 瓦式逆冲断层带,它们被认为代表了加勒比弧系统向 Chortis 地块的增生(也参见 Pindell et al., 2005; Pindell and Kennan, 2009; Ratschbacher et al., 2009)。这个含弧地块(图 9.15)继续延伸穿过 它与北美的巴哈马岸(Bahamian Bank)形成的穿时碰撞带,该碰撞带以古巴和伊斯帕尼奥拉岛 (Hispaniola)为代表,通过维尔京群岛(Virgin Islands)(Schrecengost, 2010),到达依旧活动的安的 列斯群岛(Antillian),然后到达南美洲北部,在那里,它沿海岸线自西向东穿时递进(Ostos et al., 2005)。安的列斯弧是大弧的一部分,被拉代西拉德岛(Antillean)存在侏罗纪大洋基底和燧石所支 持(Mattinson et al., 2008; Montgomery and Kerr, 2009)。在厄瓜多尔和哥伦比亚,在坎帕期-麦斯里 希特期时期,该弧与向西倾的位于俯冲带之上的南美洲西部边缘碰撞,之后才出现了太平洋岩石圈的 向东俯冲(Jaillard et al., 2004; Luzieux et al., 2006; Vallejo et al., 2009; Altamira Areyán, 2009)。因 此,大弧(Great Arc)也是卢比亚带状大陆的一部分(图 9.16)。

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图 9.15 加勒比地区地貌图 展示大弧 (Great Arc) 岩石 (灰色条带)大致分布和正文提到的其他特征

总体说来,从阿拉斯加到南美北部,拉勒米造山运动差不多是同时进行的,这就实际上排除了平板俯冲或高原俯冲作为其成因。而且,几乎所有的地方都可以证明与含岛弧地岩块的碰撞中,北美是下盘,这里称为卢比亚带状大陆。因此,孤立地对待大盆地区的厚皮变形,而不考虑同时代总体变形(Saleeby, 2003; Liu et al., 2010; Jones et al., 2011)的模型不太可能获得成功。

取而代之的是一个认为卢比亚带状大陆与美洲大陆西部边缘碰撞的碰撞模型,它明显能够最好地 契合这些数据资料。该带状大陆包括从阿拉斯加到南美所有之前拼合的地体,包括格雷罗复合地体和 加勒比海的大弧。板块运动重建显示,在这个时期,太平洋盆地内的板块运动有很强的向北运动分量 (Doubrovine and Tarduno, 2008)。

在碰撞过程中和碰撞之后,向弗朗西斯科和相关增生杂岩(例如楚加奇等)之下的俯冲停止, 板片断裂,一期与俯冲板片断离有关的岩浆活动上升至碰撞带,弗朗西斯科高级变质岩被迅速剥露至 地表,压扭推动卢比亚内大量先前汇聚的地体向北移动,撞击在北美克拉通上形成落基山褶皱-逆冲 断层带。这一设想在许多方面都与 Maxson 和 Tikoff (1996)以及 Johnston (2001)提出的相似。下面 我们来探究拉勒米碰撞的一些后果。

# 9.19 白垩纪—古近纪板片断离岩浆作用

Coney 和 Reynolds (1977) 仔细研究过西南地区晚白垩世—古近纪岩浆作用的分布情况,但是他 们推测大陆边缘走向大致为南北向,所以他们创建了一个模型,模型中源自一个向东倾斜的俯冲板片 的岩浆作用,在120 Ma 开始向大陆内侧挺进,然后在40 Ma 再次向外侧发育。然而,正如我们所看 到的,基于变形作用和岩浆活动,这一部分边缘朝向现今大致是东-西向的并且面朝南,这意味着在 80~60 Ma 期间,该地区的岩浆活动基本上是同时发生的。这条带包括亚利桑那州南部和新墨西哥州



#### 图 9.16 加勒比海演化的 3 种模型

在模型1中,向东俯冲到大弧之下导致哥伦比亚-加勒比大洋高原(CCOP)碰撞和俯冲逆转到弧之下(Burke, 1988; Rogers et al., 2007; Altamira-Areyún, 2009),然后在南北美洲之间一起移动进入大西洋。在模型2中,该弧和 CCOP 的碰撞发生在120和110 Ma之间,在那个时候俯冲快速翻动(Pindell et al., 2005)。这两个模型的主要问题 在于格雷罗(Guerrero)含岛弧超级地体下部的俯冲是向东倾斜的,随后与 CCOP 碰撞,但是在墨西哥,该超级地 体在拉勒米事件时期是向西倾斜的一个俯冲带的上盘,那样的话要求有一个俯冲极性反弹。模型3是本文推荐的模型,在该模型中,格雷罗超级地体和大弧下部的俯冲一直是向西倾斜,有一部分还建立在 CCOP 的东部边缘, 因此该弧无需反转它的俯冲极性

所谓的拉勒米岩浆活动,80 Ma 之后的莫哈韦沙漠地区的埃达克质-碱性深成岩体和76~55 Ma 深成岩体线性带继续向南延伸,遍及墨西哥西部大部分地区 (Damon et al., 1983; Zimmermann et al., 1988; Titley and Anthony, 1989; 1995; McDowell et al., 2001; Henry et al., 2003; Valencia-Moreno et al., 2006, 2007; Ramos-Velazquez et al., 2008)。

索诺兰沙漠地区有大量斑岩铜矿,并伴随着它们形成约 75~50 Ma 之间的深成岩源区 (Titley and Anthony, 1989; Titley, 1982; Damon et al., 1983; Barton et al., 1995; Barra et al., 2005; Valencia-Moreno et al., 2006, 2007)。许多侵入体尚未进行锆石精确定年,所以并不知道它们精确的年龄。尽管如此,根据 65±10 Ma 的年龄, Hildebrand (2009)认为索诺兰沙漠地区中的斑岩矿床可能与俯冲板片断离有关,这就像已出现在阿尔卑斯山带 (de Boorder et al., 1998)、巴布亚新几内亚的中央山脉造山带 (Cloos et al., 2005; McDowell et al., 1996)和西南太平洋其他地方的板片断离情况一样 (Solomon, 1990)。

在不列颠哥伦比亚省的海岸深成杂岩中, 侏罗纪—晚白垩世的英云闪长质-花岗闪长质深成岩体 在弧岩浆活动停止后被快速剥露, 英云闪长岩-花岗闪长岩在角闪岩相-麻粒岩条件下变形变质为片 麻岩,并且通常被认为其构成了科迪勒拉型岩浆弧的中下地壳(Armstrong, 1988; Hollister, 1982; Van der Heyden, 1992; Crawfor et al., 1999)。伸展作用最晚在 60 Ma 已发生(Armstrong, 1988), 且卷 入了至少 15 km 的构造剥露, 它伴随着大量晚白垩世—古近纪深成岩侵入岩的形成,包括多来源的埃 达克质和 A 型岩体(Hollister and Andronicos, 2000, 2006; Hollister et al., 2008; Andronicos et al., 2003; Gehrels et al., 2009; Mahoney et al., 2009), 这些是板片断离型岩浆活动的标志性特征。海岸 山脉岩浆活动的成分与青藏高原的岩浆岩相似。青藏高原的后碰撞岩浆作用是变化的,但占主导的是 埃达克质和碱性成分,这就说明在始新世新特提斯板片断离之后,发生了地壳的逐步抬升和热地幔向 北分流(Chung et al., 2003, 2005; Wang et al., 2010; Searle et al., 2011)。蒙大拿州碱性岩浆省的强 碱性和亚碱性岩石带与青藏高原北部年轻的岩石相似,位于板片断离侵入体主带的内侧,所以可能也 如 Chung等(2005)在青藏高原模拟时预想的那样,反映了碰撞加厚岩石圈的拆离和随后热软流圈 的上涌。上涌的热地幔可能是对该区域之下测量到的"快速"的地震波的合理解释(Schmandt and Humphreys, 2011)。

刚好在晚白垩世之前弧岩浆活动的停止,至少在 60~50 Ma 时片麻杂岩中心的快速剥露和大量晚 白垩世—古近纪岩浆活动,被认为是代表了从碰撞前大陆弧岩浆活动、向同碰撞、到后碰撞板片断离 岩浆作用和抬升剥露的转变。板片断离时,软流圈顺着狭窄撕裂处上涌形成岩浆岩,很好地解释了沿 西部边缘带伸长的晚白垩世英云闪长质岩浆岩高度集中的特征 (例如, Barker and Arth, 1990)。总体 来说,海岸山脉杂岩内约 1500 km 长的晚白垩世—古近纪深成岩线性条带可能是地球上所发现的板片 断离岩浆活动的最佳露头实例。

重建 80~70 Ma 海岸深成杂岩古纬度时,其南端和索诺兰沙漠分区能很好地匹配在一起,形成一条连续的晚白垩世—古近纪板片断离岩浆作用带(图 9.17)。经过这样的重建,如今位于海岸深成杂 岩南端的 Swakane 片麻岩与几乎完全相同的 Pelona 和 Orocopia 片岩处于同一位置,并且它也与半岛山脉、爱达荷岩基和海岸深成杂岩内部的走向滑动断层和 110~100 Ma 的盆地关闭相匹配,这说明它们属于同一系统的一部分。因此,有3条独立的地质证据可以支持这样的重建(图 9.18)。

在菲尼克斯断层和奥罗菲诺断层之间的岩浆活动间断期间,唯一的晚白垩世—古近纪岩浆活动出现在科罗拉多矿带(Colorado Mineral),这条带呈南西-北东走向穿过该区(Wilson and Sims, 2003)(图 4.1、图 5.12),这条带将拉勒米盆地分成两个不同的区域(图 4.1),明显有如 Hildebrand (2009)所说的由板片断离形成的岩浆,其成分可能混合了软流圈来源的玄武岩,以及熔融的元古界结晶基底,侵入在北美板块下盘(Stein and Crock, 1990; Bailley and Farmer, 2007),穿过一个北美大陆岩石圈裂隙处,该裂隙是沿长期存在的岩石圈边界(McCoy et al., 2005)发育(McCoy et al., 2005)。最近, Chapin (2012)对该区域的基础数据进行了非常出色的总结,表明该带岩浆活动源自向东倾斜的俯冲系统中的板片撕裂,这同那个时期北美明显缺少岛弧岩浆活动和这里提供的西向俯冲模型不一致。我认为存在另一种可能:当加拿大分区和索诺兰沙漠分区分离时,北美岩石圈破裂产生



图 9.17 加拿大-大盆地和索诺兰分区晚白垩世—古近纪期间相对位置构造模型 展示晚白垩世—古近纪期间加拿大-大盆地和索诺兰沙漠分区内各种岩浆要素可能的相对位置;奥米尼卡(Omineca)岩基和 内华达山脉岩基与其他科迪勒拉型岩基不同点在于它们没有后碰撞的板片断离岩浆活动,所以在这里一并进行考虑;然而, 内华达山脉的古地理位置有很大的不确定性。C—卡马克火山岩(Carmacks volcanics)

科罗拉多矿带的岩浆活动(图 5.12)。随着卢比亚已经与北美大陆发生碰撞,可以设想加拿大分区强 烈的向北运动牵引着大盆地分区也向北运动,直至足以打开一个裂缝,使软流圈岩浆进入地壳,并在 那里熔化和同化地壳,产生了所观测到的岩浆活动。

# 9.20 弗朗西斯科俯冲停止

弗朗西斯科俯冲在大约 80 Ma 停止,在拉勒米事件期间或紧随其后,虽然还没有得到大多数人的 认可,但是三条证据指示这一俯冲肯定是在那个时期停止的:①已知最年轻的蓝片岩是 84 Ma;②在 大约 80→53 Ma 期间存在一个突然的沉积间断;③连贯的蓝片岩在 84 Ma 之后剥露,甚至部分地区在 67 Ma 被剥露至地表(图 9.19)。因此,我将弗朗西斯科分成两部分:弗朗西斯科 1 和弗朗西斯科 2, 它们被一个没有俯冲的时期分开。在这个间歇期,运动大部分都沿边缘向北,弗朗西斯科和大谷地的 岩石都被剥露和抬升。这些想法和 Jayko (2009)的类似。

尽管海岸弗朗西斯科岩石还没得到很好地定年,它很可能比约52 Ma 年轻,这是 Snow 等 (2010) 测定的圣布鲁诺 (San Bruno) 带中最年轻的碎屑锆石年龄。Tagami 和 Dumitru (1996) 收集的低分辨





率数据支持这一结论。类似地,大谷地群内大部分沉积作用在约75和65Ma之间的坎帕期-麦斯里希 特期都明显已经减弱或停止(Ingersoll, 1979; DeGraaff-Surpless et al., 2002)。因此,现有的碎屑锆石 数据来自于增生楔,80Ma之后没有岩浆活动且缺乏大谷地弧前盆地的沉积,说明东向俯冲在大约80 Ma停止(图9.19)。弗朗西斯科俯冲的停止与内华达岩基内岩浆活动的停止、塞维尔薄皮逆冲作用 的结束,以及拉勒米变形作用的开始时间都是一致的,所以我将它和拉勒米带状大陆的碰撞联系起 来。楚加奇增生杂岩现在位于阿拉斯加南部,但是如前所述,以前应该位于更遥远的南部,它在84~ 68Ma期间也不活动,所以进一步让人确信拉勒米碰撞之后出现了广泛的俯冲停止。我们现在来探究 这次停止的影响以及可能的原因。

# 9.21 弗朗西斯科杂岩剥露

弗朗西斯科杂岩和海岸山脉蛇绿岩间的接触带长期以来被认为是一个断层,并且最初被视为一个 残余俯冲带(Bailey et al., 1970; Ernst, 1970)。最近,地质工作者认为由于相对未变质的大谷地群 岩石与蓝片岩相弗朗西斯科岩石的变质程度(等级)差异很大,所以有一个较厚的地壳断层缺失, 这个断层肯定是拉伸的张性断层,而且发生了正断层性质的移动(Suppe, 1973; Jayko and Blake, 1986; Platt, 1986; Jayko et al., 1987; Harms et al., 1992)。类似地,穿过中央带和海岸带(Central and Coastal belts)之间的断层存在 3kb 的压力差(Terabayashi et al., 1992),所以在某种程度上有一 个"三明治"构造,那里较高级的变质岩夹在低级变质岩之间(Suppe, 1973)。这不是一个新问题,





我推断两个弗朗西斯科(Franciscan)杂岩属于明显不同的发育时期,它们之间存在一个非俯冲时期。从这张图上可以很明显地看到131~95 Ma 沉积物的快速掩埋,约80 Ma 时弗朗西斯科俯冲的停止和蓝页岩的迅速剥露,至少其中的一些在约67 Ma 才露出地表(Berkland,1973)。新的俯冲开始于60 到 53 Ma 之间的某个时间,其中的代表就是沿海弗朗西斯科。碎屑锆石年龄(有颜色的小圈)来自 Dumitru 等(2010),其中的108 Ma 年龄来自 Unruh 等(2007)。代表估计深度和变质年龄的有色方块改编自 Dumitru(1989),与同变质带的碎屑锆石年龄的颜色相匹配。SFMS——南福克山(South Fork)片岩(McDowell et al.,1984; Jayko et al.,1986); VSF——瓦伦廷泉构造(Jayko et al.,1986); YBT1——约拉波利 (Yolla Bolly) 地体的硬玉辉石带;YBT2——约拉波利地体的硬柱石霰石带(Suppe,1973; Jayko et al.,1986); DIABLO 1——Diablo Range 的硬玉辉石带;DIABLO 2——Diablo Range 的硬柱石霰石带(Suppe and Armstrong,1972; Moore and Liou,1979; Cloos,1983)。Diablo Range 的裂变径迹冷却年龄和峰期剥露年代来自 Unruh 等(2007)。注意在弗朗西斯科已知的最年轻的蓝片岩变质时代是科尼亚斯期(Coniacian)(Blake et al.,1985; Wakabayashi and Unruh,1995; Jayko,2010,个人沟通)。据 Jayko (2009)的讲义修改。

甚至不是一个局部性的问题,因为世界范围内的相似变质带中,5~15 kb 的压力差很常见(Maruyama et al., 1996)。尽管如此,对于弗朗西斯科,Platt (1986, 1993)认为蓝片岩是在同俯冲伸展期间被 正断层剥露出来的;Ring和Brandon (1994)提出剥露是沿着一个层序混乱的向西逆冲断层发生的;Cloos (1982)主张一个通道流模型,即在一个狭窄通道内存在活跃的流动;Ring (2008)主张对一 个抬升的弧前高地的侵蚀剥露;Unruh 等 (2007)支持同沉积伸展;Krueger和Jones (1989)认为拉 勒米板片俯冲的角度变浅造成了这种抬升;Terabayashi 等 (1996)认为是一个地幔楔挤出过程;许 多学者 (Cloos and Shreve, 1988a, 1988b;Dumitru, 1989;Jayko et al., 1987)提出底侵作用作为造成 隆升的机制。所有这些模型都失败了,因为他们都试图在持续进行的俯冲过程中形成剥露,却没有考虑到俯冲可能在剥露作用启动之前就停止了。

由拉勒米碰撞引起的板片断离能够引起俯冲在大约 80 Ma 时停止的认识,可能为连贯的蓝片岩地体的抬升和剥露提供的一种机制。由于板片断离机制及其在造山带和岩浆作用方面的应用已在别处进

行了回顾 (Price and Audley-Charles, 1987; Sacks and Secor, 1990; Davies and von Blanckenburg, 1995; Hildebrand and Bowring, 1999; Davies, 2002; Levin et al., 2002; Haschke et al., 2002, Cloos et al., 2005; Hildebrand, 2009), 这里就不再重复。尽管完全正确地估计板片断离期间产生的物质强度和正负浮力是困难的,模拟 (Davies and von Blanckenburg, 1995) 清楚地显示在碰撞或者甚至是板片停滞时,致密的板片更下部分将会被撕下,这使得位于较浅位置的浮力更大的上层部分上升,不论它属于大洋物质还是大陆物质。如果是大洋物质,残余板片的上升可能主要是由热的软流圈上升流驱动。而且,这种断离和随后的剥露以及物质的上升可能会形成一个与剥露物质的表面对应的主要正断层,就如同 Chemenda 等 (1996) 实验所显示的那样。板片断离导致高压增生楔岩石剥露的观点并不是一个新观点,因为很多年前 Maruyama 等 (1996) 就提出了一个包含两个阶段的模型,即先由板片拆沉引起楔形挤出,然后发生穹隆状抬升。

总之,由停止-诱导的板片断离所造成的隆升可能是解释弗朗西斯科增生杂岩内部所观测到的各种关系的最有效的机制,这不仅是因为它解释了连贯的蓝片岩的快速剥露,还因为它解释了它们与海岸山脉蛇绿岩的关系(图 9.20)。这一概念和 Hildebrand (2009)的模型非常吻合,Hildebrand 给出了证据证明在拉勒米造山运动后,拼合碰撞带的西部边缘一直没有俯冲作用,直到约 53 Ma 才形成一个新的俯冲带,并且形成了从蒙大拿州到育空地区,以及从亚利桑那州向南穿过墨西哥西部的岛弧岩浆活动。



#### 9.22 地体向北迁移

不列颠哥伦比亚省、阿拉斯加州和华盛顿州北部向北迁移过很大距离的观点并不新颖,因为它已 经出现了几十年(Beck and Noson, 1972; Irving et al., 1980),被称为 Baja-BC(不列颠哥伦比亚) 假说(Irving, 1985),因为加拿大科迪勒拉山系的许多岩石相对于北美相似时代的岩石具有异常浅的 古地磁极,这意味着不列颠哥伦比亚的科迪勒拉山系大部分在约 90~60 Ma 之间向北移动了约 2~ 3000 km(Irving et al., 1995; Wynne et al., 1995; Kent and Irving, 2010)。如今人们认识到大陆地壳 的碎片可能会移动得非常迅速,比如在 10~15 Myr 移动 1000 km (Umhoefer, 2011)。

在大约 80~75 Ma 科迪勒拉山系开始了一个新的构造体制,那时①在前陆拉勒米厚皮变形开始; ②科迪勒拉型岩基岩浆作用停止;③弗朗西斯科杂岩内的高级变质作用停止;④大谷地群内的深水沉 积作用变浅;⑤板片断离岩浆作用在海岸深成杂岩和索诺兰沙漠地块开始活动。这些基本同时发生的 事件说明卢比亚两侧的俯冲大约在同一时间停止。

在今天发生相向俯冲的例子如菲律宾群岛,在那里,由转换边界连接的相向俯冲板块的相互作用 没有被充分地显示成像,所以人们对其知之甚少。另一个发生相向俯冲的地区是西南太平洋的所罗门 群岛,在那里,成像显示出较老的向西倾的太平洋板块带着厚厚的翁通-爪哇海底高原下沉到年轻得 多的森林云雀板块之下 (Mann and Taira, 2004)。这两个板块在深处究竟如何相互作用,还没有得到 解决,这主要是因为向东倾的板块太年轻了,不过我们可以设想出几种情形:①两个板块只是刚刚相 遇并陡峭地下降到地幔中;②较年轻可能也较弱的板块向自身弯曲,并与较老的板片一起被往下运 送;③两个板块都后撤,所以它们没有在一起接触太长时间;④澳大利亚克拉通前锋带进入这个海 沟,向东推动太平洋板块;⑤一个板块断离了,然后撞上了另一个板块,使地壳从地幔拆离,于是它 的板块断裂。我们可以很容易地想到这样的场景:大量的水从澳大利亚边缘的外缘和/或翁通-爪哇 高原中释放出来,以及板块被撕裂和掉落的变化。

另一个得到充分研究的例子出现在新西兰,在那里,相向俯冲带由阿尔卑斯(Alpine)转换断层相连。Liu和Bird(2006)模拟了澳大利亚与太平洋板块的碰撞,证明了一个板块冲撞另一个板块(也许发生在板片断离之后),使其拆沉的楔入模型最符合其地貌特征、隆升速率、地表侵蚀和深部地震活动。在菲律宾和新西兰,板块彼此相互作用和断离时,压扭机制占据了主导。

无论 80~75 Ma 北美那样的情形产生的准确原因是什么, 卢比亚之下向东和向西两个俯冲的停止, 加上东太平洋盆地内大洋板块强劲地向西北向移动和北美洲的向西南移动 (Doubrovine and Tarduno, 2008; Kent and Irving, 2010), 明显已引起部分卢比亚超级地体相对于北美大陆向北迁移。事实 上, 把地体组合的地质条件以及来自地体内各类岩石的古地磁资料结合起来, 表明大部分卢比亚在约 80 Ma 时开始向北迁移 (图 9.14), 并至少一直持续到约 58 Ma, 当落基山前陆缩短停止的时候, 以 冲断楔形体的有效剥露 (Price and Mountjoy, 1970; Ross et al., 2005) 和巨大的贝尔特--珀塞尔逆冲 岩席的剥蚀为标志 (Sears, 2001)。晚期压扭性质迁移的明显标志是沿着如 Tintina (Gabrielse et al., 2006) 那样的离散高角度断层运动, 以及沿着如 Denali (Fuis and Wald, 2003) 那样的断层持续运动 至今。

一个有趣的并发情况是在大约 85 Ma 可能出现了库拉-法拉龙(Kula-Farallon)扩张脊(Woods and Davies, 1982; Lonsdale, 1988)。尽管由于现在这条脊已经完全被毁了,它在当时的位置非常不确定,但是它的位置可能非常靠南,大约在今天的中美洲(Engebretson et al., 1985)。一些学者(Wallace and Engebretson, 1984)认为库拉板块强烈的北向迁移导致了阿留申(Aleutian)海沟和阿拉斯加-塔尔基特纳山脉和库斯科昆姆(Kuskokwim)山脉带中晚白垩世—古近纪线性岩浆带的产生。然而,正如我们将在本节稍后看到的那样,这些岩石可能在更远的南方,而且比目前更偏北。

如前所述,加拿大科迪勒拉(通常称为不列颠哥伦比亚省省会,或其缩写 Baja-BC; Irving, 1985)大部分地区的向北运动的看法一直都备受争议。几位作者(Umhoefer, 1987, 2000, 2003; Umhoefer and Blakey, 2006; Irving and Wynne, 1992; Cowan et al., 1997)曾简单回顾了这些争议历史,这

里就不必重复。Kent and Irving (2010) 重新计算的北美参考磁极提供了一个更新和更强大的框架, 用于评估相对于北美克拉通的地体移动。在我看来,地体大规模向北移动的最有力证据在于它们的一 致性 (表 9.1),也就是各类地体内的晚白垩世岩石都存在不一样的磁极,以一致的方式离开了北美 磁极,指示约 80 Ma之后沿边缘发生了右旋剪切 (Beck, 1991, 1992)。而且,许多现代研究没有利 用深成岩岩石,而只是关注有火山岩和沉积岩的层状区域,以减轻或至少合理评价压实和后沉积变形 的影响。在下文中,我利用古地磁数据中最强有力的古磁极,试图将它们放置在一个相关的和符合逻 辑的框架内,主要是为了测试整体模型,并评估这样巨大的移动在科迪勒拉造山运动的构造发展史上 是否能说通 (表 9.1)。

単元	年 龄	实测的古纬度	预计的古纬度	北向运输	资料来源
Silverado Formation	60±2Ma	$25^{\circ} \pm 7^{\circ} N$	37°±3°N	12°±6°	Morris et al. (1986)
Silverado Formation	62±2Ma	26°±6°N	37°±3°N	11°±5°	Lund and Bottjer (1991)
Carmacks volcanics	70±1Ma	54. 8°±4. 1°N	72. 1°±2. 7°N	17. 3°±5. 5°	Enkin et al. (2006a)
Punta Baja Formation	70±3Ma	29°±13°N	34°±4°N	5°±11°	Filmer and Kirschvink (1989)
Pigeon Point Formation	约 71±7Ma	21°±5°N	47±2°N	24°±5°	Champion et al. (1984)
Point Loma Formation (N)	72±2Ma	22°±4°N	37°±5°N	14°±5°	Bannon et al. (1989)
Point Loma Formation (R)	72±2Ma	$20^\circ \pm 12^\circ N$	37°±5°N	$17^{\circ} \pm 10^{\circ}$	Bannon et al. (1989)
Rosario Formation (P Baja)	74±6Ma	$26^{\circ} \pm 7^{\circ} N$	$34^{\circ}\pm5^{\circ}N$	8°±7°	Flynn et al. (1989)
Nanaimo Group	75±8Ma	35. 7°±2. 6°N	60. $7^{\circ} \pm 3^{\circ} N$	25°±3.7°	Enkin et al. (2001)
Rosario Formation (P San Jose)	77±3Ma	$25^{\circ} \pm 2^{\circ} N$	36°±5°N	11°±5°	Filmer and Kirschvink (1989)
MacColl Ridge Formation	80Ma	53°±8°N	68°±6°N	15°±8°	Stamatakos et al. (2001)
Ladd and Williams formations	82±8Ma	27°±5°N	38°±5°N	11°±6°	Morris et al. (1986)
Valle Formation 1 (Vizcaino)	85±1Ma	$22^{\circ}\pm8^{\circ}N$	$36^{\circ} \pm 4^{\circ} N$	13°±8°	Patterson (1984)
Valle Formation 2 (Vizcaino)	87±1Ma	$20^{\circ} \pm 5^{\circ} N$	$36^{\circ} \pm 4^{\circ} N$	16°±5°	Patterson (1984)
Valle Formation 3 (Vizcaino)	90±2Ma	$25^{\circ} \pm 4^{\circ} N$	$36^{\circ} \pm 4^{\circ} N$	11°±4°	Patterson (1984)
Valle Formation 4 (Cedros Island)	90±2Ma	$22^{\circ}\pm5^{\circ}$	$37^{\circ} \pm 4^{\circ} N$	15°±5°	Patterson (1984)
Valle Formation 5 (Vizcaino)	90±2Ma	$25^{\circ} \pm 2^{\circ} N$	$35^{\circ} \pm 4^{\circ} N$	9°±4°	Patterson (1984)
Valle Formation 6 (Vizcaino)	90±2Ma	$32^{\circ}\pm6^{\circ}N$	$36^{\circ} \pm 4^{\circ} N$	4°±6°	Patterson (1984)
Silverquick-Powell Creek Formations	90±5Ma	39. 5°±2. 2°N	59. 8°±3°N	20. 3°±2. 7°	Enkin et al. (2006b)
Blue Mountains terranes	约 93Ma	39. 2°±4. 5°	55. 1°±2. 7°N	15.9°±4.1°	Housen and Dorsey (2005)
Valle Formation 7 (Vizcaino)	94±2Ma	20°±1°N	35°±4°N	14°±4°	Patterson (1984)
Valle Formation 8 (Vizcaino)	94±8Ma	$24^{\circ} \pm 12^{\circ} N$	36°±4°N	$12^{\circ} \pm 10^{\circ}$	Hagstrum et al. (1985)
Valle Formation 9 (Cedros)	95±5Ma	21°±3°N	37°±4°N	16°±4°	Smith and Busby-Spera (1993)

表 9.1 所选科迪勒拉单元的古地磁学数据

注: P-半岛 (peninsula); N-正极性 (normal polarity); R-反转极性 (reverse polarity)。

来自 Stildnia 和兰格利亚的数据(图 9.14)表明它们是联系在一起或者至少彼此相当接近,根据 是晚三叠世—早侏罗世的纬度,在 225 Ma 的兰格利亚 Karmutsen 组(Irving and Yole, 1972, 1987; Schwarz et al., 1980; Yole and Irving, 1980)和 210 Ma 的 Stikinian Savage 山脉组(Monger and Irving, 1980)在那个磁极上彼此相当接近,就象 195 Ma 的兰格利亚博南扎组(Irving and Yole, 1987)和 195 Ma 的 Stikinian Hazelton 群(Monger and Irving, 1980)在磁极上非常接近。从温哥华岛采集的晚 普林斯巴期(Late Pliensbachian)菊石具有古地中海的亲缘性和与古地磁数据库一致的古生物纬度 (Smith et al., 2001)。Kent 和 Irving (2010)指出:①从 225→90 Ma,兰格利亚-Stikine 联合地块向北 运动了大约 20°,而北美克拉通向北运动了大约 35°,所以兰格利亚-Stikine 地块相对于北美存在一个 净向南的运动;②在 90 和 50 Ma 之间,兰格利亚-Stikine 地块向北移动了 20°,而北美克拉通在约 145 和 90 Ma 期间一直向南运动,在 90→50 Ma 又增加了 5°,所以从角度上讲,总的相对运动在右旋 方向为 25°。这些运动解释了早期左旋剪切,然后是右旋剪切。

位于 Stikinia 东边的卡什溪地体含有古地中海的动物群和具有 DUPAL 异常的玄武岩 (Monger and Ross, 1971; Johnston and Borel, 2007; Johnston, 2008),而古地中海的动物群也存在于温哥华岛的晚三叠世—早侏罗世的岩石中 (Tozer, 1982; Smith et al., 2001)。所以,除了晚期与卢比亚超级地体一道的向北迁移,那些地体在与卢比亚超级地体合并之前,从古地中海赤道地带向东移动了相当大的一段距离。

经过对 91 Ma 登上斯图尔特(Mount Stuar) 岩基的分析(Beck and Noson, 1972)、二次分析(Beck et al., 1981)、三次再分析(Ague and Brandon, 1996)以及其后更为细致的研究,发现它相对于北美向北移动了大约 24.5°±6.3°。同样地,在认识到它们的构造复杂性之后,再次研究了 109 Ma的杜克岛(Duke Island)层状超基性杂岩体,发现其相对于北美克拉通存在大约 2350 km (21°)的异常(Bogue and Grommé, 2004)。Enkin发现温哥华岛上 75 Ma 的纳奈莫(Nanaimo)群岩石向北移动了约 2750±400 km (Enkin et al., 2001),而对 Methow 盆地(图 4.1)东侧的阿尔必期(Albian)温思罗普(Winthrop)组的叶缘(leaf margin)研究,显示其北向位移了大约 2200 km (Miller et al., 2006)。

在更向东的一些地体中, Rees 等(1985)发现 Quesnellia 的 94 Ma 古磁极位置与更西边的地体类 (似,因为它们相对于北美都存在 23°±10°的异常。加拿大科迪勒拉山前陆带内超过 500 km 走向长度 的详细古地磁研究显示,弗兰特岭(Front Ranges)内的古生界碳酸盐岩有一个急剧的晚白垩世再磁 化,其磁极与北美克拉通的磁极一致(Enkin et al., 2000)。如果在主山脊西部已重新磁化碳酸盐岩 岩石中测量的缓倾磁角也在晚白垩世形成,那么这两条带不同的磁倾角说明卢比亚与北美之间的缝合 带在主山脊的踢马圈或就要其西边,之前穿过碳酸盐岩-页岩的相变相关性是错误的(Enkin, 2006)。

正如前面所讨论的,包括 95~85 Ma 的 Silverquick 和 Powell Creek 组岩石,位于不列颠哥伦比亚 省南部,它们代表了覆盖在海岸深成杂岩和山间的超级地体之上的层序;根据可靠的沉积岩和火山岩 的古地磁研究结果,自堆积以来它们往北迁移了约 2300 km (Wynne et al., 1995; Krijgsman and Tauxe, 2006; Enkin et al., 2006b; Enkin, 2006)。这与 70 Ma 的卡马克斯 (Carmacks) 群火山岩的结 果完全吻合,该群位于 62°N 的育空地区海岸深成岩带的北端(图 2.5 和图 5.12),当它们在 70 Ma 喷发时 (Enkin et al., 2006a; Wynne et al., 1998),位于现今美国西部俄勒冈州的纬度。在那时,北 美的位置比现在更往北一点,这样旧金山是位于约 45° (Kent and Irving, 2010)。由于杂岩长约 1500 千米 (从育空地区延伸到华盛顿州的高喀斯喀特),其南端将差不多位于现今加利福尼亚州南部-墨 西哥北部的纬度。目前,在华盛顿州的南端有一个奇怪的岩体,称为 Swakane 片麻岩,它几乎每一个 已知的属性包括年龄、成分、快速深埋和剥露,以及构造背景,与南加州和亚利桑那州 Pelona 和 Orocopia 片岩都很相似 (Matzel et al., 2004)。通过恢复海岸深成杂岩在 70 Ma 时的纬度,Swakane 片麻岩大约可恢复到 Pelona Orocopia 片岩的同一个带(图 9.17)。类似的结果也从阿拉斯加 78 Ma 的 MacColl Ridge 组层状岩石中获得,它属于兰格利亚复式地体的一部分,目前位于 61°N,将它们放到 重建的旧金山地区,它要比现在的古纬度低 15°±8° (Stamatakos et al., 2001),而与卡马克斯 (Carmacks)的结果一致。

95~85 Ma 鲍威尔河(Powell Creek)和 Silverquick 组位于海岛和山间(Insular and Intermontane)两个超级地体之上。

26 个火山熔岩流点和 54 沉积岩点,连同几个正向接触带、砾岩和倾角测试,所有这些汇总起来,提供了该地区地磁场倾角合理而可靠的记录,获得古纬度为 39.5 °±2.2 °,这比预计的那个时候的北美古纬度约 20.3°±2.7°偏南(Enkin et al., 2003; Enkin et al., 2006b; Kent and Irving, 2010)

在以前的研究者(Pope and Sears, 1997; Wernicke and Klepacki, 1988)提出的模型中,卡斯尔地 台和加拿大科迪勒拉山系的 Stikinia(图 2.5)从爱达荷-蒙大拿向北逃逸,但地体之间的联系和上面 所引用的古地磁数据表明,加拿大科迪勒拉山系地体在 70 Ma 已经拼合到一起,并位于相当远的南 方。事实上,如前所述,今天位于海岸深成杂岩北端的育空地区的卡马克斯火山岩,过去位于旧金山 附近,所以看起来更有可能的是,加拿大科迪勒拉山系的地体最初位于大盆地分区的南端,而不是北 端。重建的科迪勒拉山系南部也将 80~75 Ma 索诺拉-莫哈韦沙漠地区的板片断离深成岩和那些海岸 深成杂岩连接起来 (图 9.17),如前所述 (Hildebrand, 2009)。

除了内华达岩基,美国西部的其它地体显示的位移与加拿大地体的运动一致。在加州海岸中部的 盐碱地块两处采集了岩石用于古地磁分析,一处是圣佩德罗(Point San Pedro)(图 5.18),那里古新 世沉积岩位于与白垩纪花岗闪长岩陡壁不整合接触,另一处是鸽子岬(Pigeon Point),该处有厚层的 坎佩尼阶-马斯特里赫特阶浊积岩露头(Champion et al., 1984)。研究结果表明,两个组的岩石最初 都在比它们现在各自 25°N 和 21°N 的纬度往南约 2500 km 处。

在古地理重建中,内华达岩基存在一个悬而未决的问题,因为已有的古地磁数据显示约100 Ma 以来,它相对于北美大陆运动距离很小(约1000 km)(Frei et al.,1984; Frei,1986)。我们从102→ 97 Ma 的深成岩体中采集了样品,正如我们所知,100 Ma 的火山岩已被褶皱,很难理解这些深成岩体 是如何避免被褶皱的。因此,磁化强度的测量可能是第二级的和褶皱后的。不管如何解释,内华达岩 基需要更多的古地磁研究。人们很容易将内华达山脉岩基与奥米尼卡岩基带联系起来(图9.17),因 为它们都没有海岸深成杂岩和索诺拉-莫哈韦沙漠地区晚白垩世—古近纪深成岩体的特征。

就位于东面的白印优山脉是一个孤立的地块,包含7km厚的早古生代岩石,往东北延续,但只进入内华达的埃斯梅拉达县,并且它和内华达山区地块可能已从墨西哥北部向北迁移了,这从内华达地区古地磁数据和其侵位之后北美大陆向南迁移来看是可能的 (Kent and Irving, 2010),整体发生了约1000km的相对位移。

在下加利福尼亚半岛南部,一些古地磁的研究得出结论:巴列斯(Valles)弧前岩石沉积的位置 比它们现在的位置更往南10°~20°(Hagstrum et al., 1985; Smith and Busby-Spera, 1993; Sedlock, 1993; Hagstrum and Sedlock, 1990, 1992)。一些固定论者(Butler et al., 1989, 1991; Butler and Dickinson, 1995; Dickinson and Butler, 1998)认为,浅成的古地磁倾角结果不是可信的古纬度指标,因为 那些取自沉积层序的岩石有可能压实变浅,而深成样品在侵位后是倾斜的和/或受到年轻得多的再磁 化的影响。Beck(1991)曾检测各种可能性,得出数据一致性显著的结论,并支持远距离迁移的 结论。

Smith and Busby-Spera (1993)也测试了承压的影响,并分析了 olistostromal 单元中的崩滑块体 (slump blocks)以检验再磁化假说。他们得出的结论是,岩石没有被重新磁化,并且向北位移 18°±7°是对这些数据的最好解释。最近,Sedlock (2003)提供了关于古地磁利与弊的很好综述,得出结论,包括巴列斯 (Valles)组在内的西部块体向北迁移是很有可能的。

沿墨西哥西岸的阿卡普尔科(Acapulco)南北两侧是一个明显在 80 Ma 后被缩短的边缘,在那里,前寒武纪—中生代变质岩和 Xolapa 地体的白垩纪深成岩岩石(图 2.5)与中美洲海沟的异常靠近,所以,有可能代表了如今在遥远的北方发现的那些岩片最初的"家"(Karig et al., 1978)。根据碎屑锆石总体, Wright 和 Wyld (2007)认为,西部大谷地群的岩石最初在墨西哥瓦哈卡(Oaxaca)附近沉积,这可能是一个合理的匹配。

跳到北方,稀少的阿拉斯加北部岩石的古地磁数据显示,许多之前位于加拿大西部的地体向北迁 移构成了阿拉斯加马蹄形山系(Johnston, 2001)。以前引用的数据清楚地表明,更多的外侧地体,如 楚加奇和威廉王子(Prince William),最初在沿海岸深成杂岩西侧分布,后来向北迁移(Farmer et al.,1993; Sample and Reid, 2003; Roeske et al.,2003,2009; Housen et al.,2008)。北极阿拉斯加、 育空-塔纳纳河、塞尔温盆地等地体的岩石彼此相似(如碎屑锆石组合)(Beranek et al.,2010a, 2010b),而且它们与内华达州西部的岩石相似,如罗伯茨山外来体,它们均在晚泥盆世—早密西西 比世变形和变质,这些特征在北美克拉通西部还未见。北极阿拉斯加在130 Ma 时似乎位于当前位置 再往南12°±5°(Halgedahl Jarrard, 1987);因此有理由得出这样的结论:这个地体群和那些后来增 生添加上去的部分,以前位于内华达地区以北的加拿大西部的纬度。

到 50 Ma (图 53)时,科迪勒拉地体向加拿大分区的向北迁移明显大部分已完成,因为不列颠哥 伦比亚省中部的始新世 Ootsa 湖火山岩大致就位 (Vandall and Palmer, 1990),50 Ma 温哥华岛的弗洛 雷斯 (Flores)火山岩 (Irving and Brandon, 1990)以及不列颠哥伦比亚省新近纪晚期的玄武岩也均已就位 (Mejia et al., 2002)。50 Ma 以后,北美及其附属的地块向南移动了一小点 (Kent and Irving, 2010)。请注意,这些地体的整体移动很好,在 25~40 Myr 中达 1000 km,完全符合现代地体平移估算量范围 (DeMets et al., 2010; Umhoefer, 2011)。

卢比亚向北迁移的主要影响是在北美地台大陆架内岩石形成了落基山褶皱-逆冲带(Price and Mountjoy, 1970; Price, 1981)。最近 Larson 等(2006)的研究表明,贝尔特--珀塞尔-温德米尔地体的 早古生代岩石位于珀塞尔复背斜的东翼,在侵入体侵位前被逆冲变形,侵入体<sup>40</sup> Ar/<sup>39</sup> Ar 定年为108 Ma;而东边的被动边缘岩石,在坎帕阶前没有被卷入逆冲,这个时间与前渊中最厚沉积楔的沉积和 含有变质岩和火成岩碎屑的沉积相符(Leckie and Smith, 1992)。Larson 等(2006)认为褶皱--逆冲带 可能代表了沿北美边缘的右旋转换挤压作用。因此,在加拿大分区,卢比亚超级地体和北美边缘之间 的缝合带有相对较晚的特点,它恰好位于北美西部大陆架边缘,在不同位置表现为不同的逆冲。在国 境线两边,缝合带是携带着贝尔特--珀塞尔超群的逆冲断层带,在某些地方(如 Lewis-Eldorado-Hoadley)越过前渊的白垩纪沉积岩石。沿着缝合带,在很多地方有半地堑(Price,出版中; Constenius, 1996),这可能是由于北美地层被拉拽到卢比亚前缘之下,由增厚地壳的垮塌所 造成。

转换挤压模型的一个有趣推论是,如果莫纳希(Monashee)杂岩(Sevigny et al., 1990; Parrish, 1995)的时间-温度曲线是正确的,不仅是卢比亚岩石要向北迁移,而且当地一些早些时候(125~108 Ma)曾被卢比亚地体逆掩和变质的北美岩石也要向北迁移。因此,莫纳希核杂岩的基底可能是北美的一部分,该地岩从位于更遥远的南方克拉通中分离出来连同它的外来体——外来的卢比亚盖层一起向北运移。因此它是有些孤立的地壳块体,类似于分隔玛丽亚褶皱-逆冲带北部和菲尼克斯断层南部的北美微地块(图 4.1)。

#### 9.23 科罗拉多高原和拉勒米大峡谷

科罗拉多高原西南角明显受拉勒米碰撞和随后的板片断离的影响很大,在晚白垩世—古近纪时明 显变得高耸并发育了向北的水系(Flowers et al., 2008; Hill and Ranney, 2008; Wernicke, 2011; Flowers and Farley, 2012)。转换带与科罗拉多高原边线附近,古新世—始新世圆状砾石层表明在高地 貌的表面存在北东向的古水流(Elston and Young, 1991; Potochnik, 2001),因为局部古河道中填充的 砾石厚达1200 m (Young, 1979),砾石中含有 80~64 Ma 的火山岩碎屑(Elston et al., 1989)。通过 峡谷和碎屑物源相结合可以得出一个令人信服的分析,即科罗拉多大峡谷(Grand Canyon)主要是在 坎帕期被加利福尼亚河在向北-北东方向流过地表切蚀而成(Wernicke, 2011)。

犹他州(Utah)南部上白垩统 Wahweap 和 Kaiparowits 组河流相地层从底到顶的碎屑锆石的最小 年龄峰值为 82 Ma、77 Ma 和 73 Ma,表明被北东流向的加利福尼亚河搬运的岩屑从南→南西逐渐变得 年轻(Larsen, 2007; Jinnah et al., 2009; Larsen et al., 2010)。这些年龄峰值太年轻,不可能来自科 迪勒拉岩基,但其与南部地区板片断离所引发的喷出和侵入岩石吻合得很好。较年轻的麦斯里希特期 (Maastrichtian) 沉积岩中含有 105~100 Ma 的碎屑锆石,表明科迪勒拉岩基或者可能的 Delfonte 火山 岩区的岩石受到了剥蚀(Link et al., 2007b; Larsen et al., 2010)。科罗拉多高原隆起但缺乏折叠和渗 透变形,恰好在玛丽亚(Maria)褶皱-逆冲断层带北部,表明两者在以后并列,最有可能是沿着菲 尼克斯断层(Hildebrand, 2009)后或者在褶皱带增厚期间的中新世崩溃。我认为科罗拉多高原是在 拉勒米事件期间形成,当时卢比亚与该区的西部和南部两边同时碰撞,在北部,高原则以奥洛菲诺断 层和/或路易斯和克拉克断层系统为界。

#### 9.24 变质核杂岩和内陆带双重构造

在北美西部,似乎有四个年龄组的变质核杂岩,它们对应于造山带的四个主要部分:①始新世段,位于加拿大分区内的路易斯和克拉克区域线性构造北部;②古新世为主的段,位于路易斯和克拉 克线性构造到死亡谷的大盆地分区;③中新世段,在索诺兰沙漠分区,这个弧向东从莫哈韦沙漠向南 一直延伸到墨西哥索诺拉;④早白垩世群,主要位于沿阿拉斯加布鲁克斯山脉的南部边缘分布。 Coney和Harms (1984)提出了阿拉斯加之外的一般性分布图,但是由于那时他们没有足够的年龄数 据,并没有给出其时间关系。

科罗拉多河-亚利桑那州南部地区的中新世杂岩是最著名的(Davis, 1980; Davis et al., 1980; Rehrig and Reynolds, 1980; Anderson, 1988; Dokka, 1989; Spencer and Reynolds, 1990; Foster and John, 1999)。它们向南延伸到索诺拉(Sonora)(Anderson et al., 1980; Nourse et al., 1994)并向西进入到 莫哈韦沙漠中部(Dokka, 1989; Glazner et al., 1989, 2002; Walker et al., 1990a)。由于变质核杂岩 的分布紧随更老的拉勒米变形带并和板片断离岩浆岩相关,笔者和 Coney(1987)认为从莫哈韦向东 延伸到亚利桑那州,并向南延伸进入墨西哥的变质核杂岩伸展,与拉勒米带的这部分在中新世的垮塌 有关。Walker 等(1990a)也将莫哈韦地区的伸展作用与中新世岩浆活动的脉动联系起来,但他并没 有考虑到这些杂岩体的整体分布。

Hildebrand (2009)认为大盆地和加拿大分区的内陆带,在古新世—始新世时期的垮塌与板片断 离有关,并引起了卢比亚之下的北美大陆隆升,但这可能并不完全正确,因为这个内陆带与其它碰撞 造山带的内陆带具有很多共性,其中一些在垮塌初并没有发生板片断离(Mattauer et al., 1983)。最 明显和相关的特点是在碰撞造山带中由于风化剥蚀和底侵作用所形成的断层下盘双重背斜 (Malavieille, 2010)。这些双重构造通常会在造山带的山前一侧孤立地形成外来岩石的飞来峰(图 9.21)。

在阿尔卑斯山脉,一个含有欧洲基底出露在缝合线之下的陶恩(Tauem)构造窗的大型地壳双重构造,已经分离出了北部钙质阿尔卑斯山,它们包括有阿普利亚地区(Apulian)物源的低级变质沉积岩,构造接触覆盖在欧洲岩石之上,在双重构造核前的外侧(Schmid et al., 2004)。类似的,最近在台湾的研究(Beyssac et al., 2007)表明,在变质程度较高的中央山脉(Central Range)的大部分变形和剥蚀都是通过底侵作用来维持的,而不只是受关键楔形体的正面积累。同样,在低喜马拉雅有一个由底侵作用形成的类似双重构造,孤立了主中央逆冲带之上的岩石,使它在北部从其主体地块分离,成为阿莫拉(Almora)飞来峰(Célérier et al., 2009; Bollinger et al., 2004)。

在阿曼造山带(Oman), Saih Hatah 山被加厚山根、双重背斜下伏,在其前陆一侧形成了孤立的 塞迈尔(Semail)蛇绿岩大型飞来峰(Hanna, 1990; Al-Lazki et al., 2002; Gray and Gregory, 2003; Searle et al., 2004; Searle, 2007)。在北美科迪勒拉内,在内陆带残余背斜构造里可能出现北美结晶 基底,如出现在红宝石、东洪堡特、Raft河-Albion、Pioneer 山、华盛顿州的 Priest 河杂岩以及不列 颠哥伦比亚的莫纳西杂岩内的背斜构造(Howard et al., 1979; Journeay, 1992; Parrish, 1995; Doughty et al., 1998; Snoke and Miller, 1988; Link et al., 2007a; Gervais et al., 2010),这表明内陆带最初形 成于板底垫托双重构造地区,并且外来岩石的孤立飞来峰位于如今北美地台岩石的东部。故正式将出 露在犹他州西部房屋山脉(House Range)的飞来峰定义为房屋山脉飞来峰。在阿拉斯加的布鲁克斯 山脉中也存在一个类似的内陆带,该区 Doonerak 构造窗(图 2.5)覆盖在沿 Doonerak 山脊分布的背 斜之上(Moore et al., 1997; Fuis et al., 2008),被认为是双重基底的构造窗(Oldow et al., 1987, 1989)。

强烈的侵蚀和剥露相结合,为同步挤压和后压缩性拉伸崩塌提供了理想的场所。因此,这些地区可以在局部同时保存逆断层和正断层以及两者之间复杂的相互作用。例如,已知白垩纪正断层沿布鲁克斯山脉南侧发育的所谓的片岩带与隆升一致 (Gottschalk, 1990; Gottschalk and Oldow, 1988; Gottschalk et al., 1998; Miller and Hudson, 1991; Law et al., 1994; Little et al., 1994)。白云母氩年龄向南



图 9.21 双重构造剖面对比图

(A) 引自 Schmid et al. (2004); (B) 引自 modified from Célérler et al. (2009); (C)、(D) 引自 Malavleille (2010)
显示位于拆离基底之上发育的变形楔状体受侵蚀剥露作用而形成的双重构造,属典型的飞来峰,位于内陆带中发育成熟的板块下盘前陆一侧: (A) 穿过东阿尔卑斯山的剖面(据 Schmid et al.,2004); (B) 穿过喜马拉雅山的剖面(据 Célérier et al.,2009); (C)
穿过台湾弧-陆碰撞带的剖面; (D) 按比例模拟的剖面; (E) 穿过大盆地的前盆岭省模拟,显示位于内陆带东侧的孤立 Ruian 飞来峰。其他例子还有阿曼山脉(Oman Mt.)的 Saih Hatat,那里塞迈尔(Semail)蛇绿岩成为了一个孤立飞来峰;阿拉斯加布鲁克斯山脉,那里 Doonerak Fenster 位于双重侵蚀岩柱的顶部,并出露了下盘岩石; Wopmay 造山带,那里 Hottah 外来地体的一个大型飞来峰,孤立地位于由太古代奴隶克拉通(Slave craton)板片形成的下盘双重构造东侧(据 Hidebrand et al.,2010a)。

增加,从约90→100 Ma (Vogl et al., 2002),而 Toro 等 (2002)收集的数据表明伸展峰期发生在112 Ma 之前,并且,在98→90 Ma 时期有个快速冷却事件。总的来说,出现含逆断层和正断层的侵蚀双 重构造,使外来岩石飞来峰孤立地出现在造山带山前一侧可能是共同的特征,或许可以作为碰撞造山 带的判别标志。

#### 9.25 盆岭省伸展

我曾提出(Hildebrand, 2009)盆岭伸展省发育在覆盖在北美克拉通之上的卢比亚超级地体的区域。在大盆地内,一套从犹他州西部到加利福尼亚州东部的大陆反射剖面合作项目(COCORP)的深部地震线显示,从内华达中部向西到加利福尼亚州东部的深部地壳中出现强烈的水平反射层(All-mendinger et al., 1987),这可能代表了缝合带以及卢比亚之下的一个可能的古生代原地硅质碎屑变沉积岩薄层。由于大多数大盆地地区都具有"正常"厚度(30~35 km)的地壳(Heimgartner et al., 2006)并且经历了差不多100%的伸展(Gans and Miller, 1983; Wernicke, 1992),在发育正断层之前地壳可能曾被加厚了近一倍。由于逆掩断层的分布与延伸区和跨塌区的分布一致,我认为由于在碰撞事件中形成卢比亚-北美双倍厚度的地壳,该区的垮塌直接导致了早—中第三世的伸展作用。而地貌可能是伸展作用的主要控制因素,可能在碰撞之后,当夹在冷的地壳下盘与其自身冷的上部岩层之间时,热的、可能已融熔的上盘板块的下地壳可能更容易侧向流动(Wernicke, 1992; Burov and Watts, 2006)。

青藏高原是一个非常好的参照物,青藏高原是由于印度和欧亚大陆之间的汇聚作用形成的具有双 倍地壳厚度的地区(Molnar and Tapponnier, 1975)。至少有一部分,也可能是全部具有加厚地壳的区 域,都与印度岩石圈向欧亚大陆之下的俯冲作用直接相关(Searle et al., 1987)。各种地球物理数据 (如地震反射、高导和低速带、高热流以及地震波的强衰减等)综合解析均表明在高原之下 15~20 km 的深度存在一个部分熔融区(Nelson et al., 1996; Schilling and Partzsch, 2001)。

在高原上,受加厚地壳重力势能的驱动,加厚区可能会沿部分熔融层向外流动,导致其上面席状 区域的伸展(England and Houseman, 1988; Teyssier et al., 2005)。对于北美而言,我认为存在一个流 变梯度带,使卢比亚板块的下部能够在缝合带之上侧向流动,而最上部地壳则只是脆性断裂。墨西哥 盆岭省一直向南延伸,从菲尼克斯断层穿过墨西哥至少到跨墨西哥火山带(Trans-Mexican) (Stewart, 1978; Henry and Aranda-Gomez, 1992; Henry et al., 1991)。尽管没有墨西哥盆岭省的深部 地震数据,但它紧邻墨西哥东部褶皱-逆冲断层带西侧,具有与大盆地相同的构造背景,因此有理由 假设下盘板块的地壳在地下继续向西延续。

在加拿大地区没有出现盆岭型伸展作用,其原因可能是由于卢比亚超地体并不像大部分北美大陆 地壳那样几乎都位于顶部,因为它在更远的南方。落基山脉褶皱带是在卢比亚地体向北移动时的转换 挤压作用形成的,因此几乎不在北美。

# 第10章 科迪勒拉山系大陆拼合总结

本研究主要的主要成果之一是, 卢比亚带状大陆在东部和西部边缘的增生, 并且, 在它与北美西 部被动陆缘拼合的时候几乎已经完全发育成熟, 第一次增生发生在塞维尔事件期间的大盆地分区, 更 彻底的一次是在拉勒米事件期间。北美大陆没有通过渐进式加积方式向西逐渐生长(图 8.1)。

在晚泥盆世—早密西西比世安特勒造山运动期间,罗伯茨山外来体在卢比亚边缘就位,粗粒岩屑 向东流出形成一个碎屑楔形体,覆盖在碰撞前的安特勒大陆架之上。在早三叠世的索诺玛 (Sonoman)造山运动期间,宝山外来体中夹枕状玄武岩透镜体的上泥盆统到三叠纪最早期的燧石-泥 质板岩沉积序列就位,覆盖在改造过的罗伯茨山外来体西部边缘之上。如果罗伯茨山外来体西部边缘 有弧,它会在宝山外来体到达之前被断层移去。它可能作为库特尼地体位于加拿大分区内。

在加拿大分区,兰格利亚和亚历山大地体被一个 309±5 Ma 深成岩体缝合在一起,那些地体在晚 古生代和中生代期间一起演化。沿着卢比亚东部边缘,向西倾斜的俯冲持续,一系列含弧地块和它们 的增生杂岩沿着向西倾的缝合带被不断添加到超级地体中。在 260 Ma 和 253 Ma 之间,卡斯尔地台-塞尔温盆地的西部边缘被白北鲑逆冲断层拉拽到育空-塔纳纳地块之下,而大洋性质的斯莱德山地体 的岩石被压缩,并在构造上位于这两个地体之间。

在 187 Ma 和 173 Ma 之间,库特尼地体西部边缘被拉到 Quesnellia 地体之下,形成了一个向东的 褶皱-逆冲带,随后不久,早期向南西的构造(如 Scrip 推覆体)在约 173 Ma 和 168 Ma 之间被向北东 的褶皱和逆冲叠覆,并被一群深成岩体侵入。第二阶段的变形对应于贝尔特-珀塞尔-温德米尔地体 173 Ma 时企图向西俯冲到库特尼地体之下。在阿拉斯加,兰格利亚在 170 Ma 时被拉到大洋性质的塔 尔基特纳弧之下,并沿其北部边缘产生了一个向北的褶皱-逆冲带和前渊。

科迪勒拉山系造山运动大概的时序关系见图 10.1,内华达山脉以西的事件见图 10.2。在这里所选的模型中,一个单独的带状大陆或复合弧地体,由新元古代—古生代小地块,如嘘飞、雷丁、崔尼蒂和怀里卡等组成,连同二叠纪—三叠纪的麦克劳德(McCloud)弧地体构成了一个晚三叠世——侏罗纪弧岩浆作用的基底,直到在 160 Ma 时与卢比亚碰撞(图 9.1)。在碰撞以前,在较老的内华达-克拉马斯地块两边都有向西倾的俯冲带。沿着它的西边缘,有向西俯冲到三叠纪——侏罗纪斯莱特溪-结合湖-海福克弧之下(图 10.2),直到 169~164 Ma 与内华达-克拉马斯地块西部的西边缘相撞(Wright and Fahan, 1988; Day and Bickford, 2004)。后碰撞深成岩体侵入到内华达(159~150 Ma)和克拉马斯(162~156 Ma)地块中,这归因于碰撞期间的俯冲板片断离。

往东,黑石-克拉马斯-内华达山脉-莫哈韦沙漠-索诺拉弧之下的向西俯冲及其主要是新元古代 一古生代基底导致三叠纪—侏罗纪大陆岩浆弧发展,从内华达州北部向东南延伸穿越莫哈韦沙漠-索 诺兰沙漠地区。这个弧与卢比亚西部边缘及其以碳酸盐岩为主的被动边缘在约160 Ma 时碰撞,形成 了向东的 Luning-Fencemaker 薄皮逆冲带和相关的向南东的逆冲断层(图9.1)。在这个碰撞过程中, 俯冲板片断离和热软流圈物质能够通过撕裂上涌进入地壳,导致地壳熔化和形成一个线性的岩浆带侵 位,包括年龄在约150 Ma 和145 Ma 之间的独立岩墙群和双峰式、碱性柯维亚深成岩套。

在约159 Ma, 斯马特维尔弧与内华达地块西部边缘碰撞形成另一个碰撞带(图9.1和10.2)。导致这次碰撞的俯冲极性是向西的,位于斯马特维尔弧之下,内华达地块的西部边缘只是部分被俯冲到该弧之下,所以,这个弧现在位于内华达地块的大陆地壳之上。

在阿拉斯加,布鲁克造山运动显然发生在晚侏罗世末—尼欧克姆期,当 Angayucham 洋关闭,科 尤库克地体与北极阿拉斯加相撞时,科尤库克地体被认为代表了位于上盘的弧地体。在布鲁克造山运



图 10.1 科迪勒拉地区不同区域时间-空间图

动前,阿拉斯加中部的许多较小的地体,如Farewell、红宝石和基尔巴克(Kilbuck)等可能就已经附属于北极阿拉斯加。关于布鲁克造山运动的确切年龄仍有一些不确定性,它可能在晚阿普第期就已经 开始。年龄相似的蛇绿岩,如Angayucham、英格尔斯、约瑟芬和海岸山脉等,分布的长度超过了造 山带,表明它们在170~160 Ma之间形成于同一边缘海。

弗朗西斯科俯冲杂岩在159 Ma之后发育,也许没有直到131 Ma之后才发育,最年轻的碎屑锆石 年龄在最老和最内侧的相关单元,即南福克山片岩(图10.2)。即使是大谷地群中最古老的岩相,石 溪含有的锆石年龄仅135 Ma。因此,这是可能的,即卢比亚西侧的向东俯冲一直没有开始,直到发 生塞维尔碰撞事件。只有外来体中显示一个较老的弗朗西斯科年龄,且它们是多旋回的,因为在并入 混杂岩之前通常已被蛇纹岩包裹。在科迪勒拉山基侵入之前,唯一可能与向东俯冲有关的岩浆活动是 产于克拉马斯和西内华达山区变质带的一小群140 Ma的深成岩体,但即使是那些深成岩体也可能是 与之前提及的稍老碰撞有关的板片断离岩浆作用。不过,这样大型的深成岩体的出现证实增生地体位 于大陆地壳之上。

俯冲杂岩还没有增生,直到 123 Ma 卢比亚的东边与北美大陆的西部在大盆地地区碰撞时(图 10.3),大谷地弧前盆地的岩石被破坏、断裂和扭曲,这发生在北美大陆以及其面朝西的被动陆缘和 克拉通台地上拱骑在外膨胀到向西倾的海沟东部、或古大洋、卢比亚超级地体侧边之后,向东沉积下 来的碎屑形成一个大片分布的砾石和卵石砾岩层。然后克拉通台地被拉拽进入海沟、变形并沿基底滑 脱面从其基底拆离,而砾石和砾岩被前渊的造山的碎片埋藏。

在大盆地地区,塞维尔褶皱-逆冲带的主要逆冲时期发生在约 123 Ma→约 108~105 Ma,导致大量的元古界—寒武系碎屑大型逆冲岩席的增生。再往西,在后来成为造山带的内陆带,卢比亚带状大陆被拆离,岩石相对于北美大陆向东、向南迁移(图 10.3)。



图 10.2 时间-空间分布图

板片断离岩浆作用开始于约96 Ma,导致了富含金属的碱性深成岩体小线性带侵入到卢比亚地壳中。深成岩体现在出露于加拿大北部和阿拉斯加东部,在那里它们被称为 livengood-Tombstone-Tungten-Fairbanks-Salcha 岩套。板片断离岩浆作用在约90 Ma 停止,并且因为碰撞只发生在大盆地分区内,所以板片断离也只发生在那里。在那个分区大约100 Ma 时向西倾的板片的断离,解释了拉勒米事件之后那里缺乏板片断离岩浆作用的现象。STEP 断层沿着该分区的南部和北部边缘在北美克拉通的下部发育,并且一直活动到拉勒米最终发生碰撞。

在大约 100 Ma 时,沿卢比亚西缘的洋盆穿时闭合,如加拿大的 Gravina-Nuzotzin 盆地,下加利福 尼亚阿尔图斯弧以东的无名盆地,和位于内华达岩基内并将其一分为二的神秘盆地等,导致了科迪勒 拉岩基带内主要的转换挤压缩短(图 9.4)。推测起来,在卢比亚西缘的斜向俯冲,就像今天苏门答 腊岛那样,被分解为一个正交俯冲组分和一个走滑组分,并伴随一个与岩浆前锋大致同时的主断层。 碰撞前俯冲与中间盆地可能是直接向西的,因为在西部地块中岩浆活动明显向东变年轻。后碰撞的深 成岩体侵位大多数在东部,但局部也穿过了缝合线,例如在下加利福尼亚 La Posta 岩套中的 98~92 Ma,和年龄为 98~85 Ma 的内华达山脊岩浆事件形成的那些深成岩体,它们可能是板片断离岩浆或者 可能是俯冲和板片断离岩浆的混合。

在大约 80~82 Ma, 拉勒米事件期间, 几乎整个长度的卢比亚带状大陆和北美发生了碰撞, 形成 一个北美大陆西缘上与前渊演化相关的、面向克拉通收敛的逆冲带, 除了在大盆地段外, 那里的碰撞 发生得最早。在碰撞过程中, 弧岩浆活动停止, 因为连接到北美克拉通的俯冲板片被撕开和折断, 并 携带裂谷边缘和大量冒地槽物质和它的克拉通基底一并进入地幔, 参与再循环。板片断离轻易地解释 了北美被动陆缘上明显缺乏裂谷沉积的现象。在北美板块和卢比亚超级地体之间的压缩作用导致厚皮 逆冲带在造山带大盆地分区发育。在卢比亚西缘上的向东俯冲作用的停止导致弗朗西斯科杂岩的连贯 蓝片岩地质体剥露 (图 10.1)。

往南到索诺兰沙漠地块,即加拿大地体当时所处的位置,板片断离导致了板片断离岩浆作用带的 形成,自菲尼克斯断层向南延伸,穿过海岸深成杂岩、Quesnellia 南端和贝尔特--珀塞尔外来体,进入 到莫哈韦--索诺兰沙漠地区,那些地体在当时显然连接在一起(图 9.17 和图 10.3)。在大盆地地区,

解释了主要构造地质、沉积和岩浆事件,从内华达山脉向西到达太平洋沿岸。部分数据来自 Dickinson (2008)。



图 10.3 卡通说明不同时间段卢比亚与北美大陆的相互作用

A. 约 125 Ma, 卢比亚到达北美附近; B. 卢比亚带状大陆与北美的大盆地分区(GBS)碰撞; C. 北美向西俯冲板块在100~96 Ma 断离,导致了板块断离岩浆涌动; D. 北美和卢比亚在这时一起向北移动,但是因为卢比亚带状大陆的移动更慢,它们之间的剪切运动是左旋的; E. 在大约 80 Ma,北美开始向南移动并且整个卢比亚与它发生碰撞,大洋俯冲板块断离产生板片断离岩浆作用,同时卢比亚地体开始和太平洋板块一起向北运动; F. 兰格利亚和 Stikinia 一起向北迁移,可能沿着 Quesnellia 边缘在一个断层,双倍加厚了加拿大科迪勒拉山系的科迪勒拉岩基

这个年龄段的岩浆活动只出现在科罗拉多成矿带。

沿着卢比亚边缘俯冲的停止,加上太平洋底强烈的向北运动,使一个巨大的、连贯的卢比亚(那时,位于索诺兰沙漠分区内,紧贴着大盆地分区的南边)(图 9.17 和图 10.3)被太平洋板块捕获,并且相对北美大陆向北移动,而北美大陆正向南移动(Kent and Irving, 2010)。到了 80~75 Ma,它的一部分撞上了加拿大分区内的北美克拉通,在那里形成了洛基山褶皱-逆冲断层带和东面厚碎屑楔形体。

到大约 58 Ma, 卢比亚超级地体的主要部分大规模向北迁移已经停止, 落基山脉褶皱-逆冲断层 带内的逆冲作用也已经结束, 尽管某些走滑运动持续进行到始新世, 离散断层如廷蒂纳和至今还活动 的麦金利山断层。加拿大分区的造山带剥露于约 58 Ma 开始, 正如贝尔特外来体的隆升和加拿大西部 盆地内厚层坎帕阶前渊的剥蚀所记录的那样。

造山带的局部重力崩溃发生的时间取决于碰撞的时间和增厚模式。在大盆地分区内,变质核杂岩 形成于古新世,这反映了塞维尔缩短阶段的底侵作用和加厚作用。然而,在加拿大分区内,垮塌发生 在始新世,是塞维尔和拉勒米事件期间加厚的结果。在索诺兰沙漠分区内的垮塌发生在中新世,是拉 勒米事件的加厚导致的。在阿拉斯加,垮陷发生在白垩纪时期,紧随着部分与布鲁克造山作用的加厚 同时发生。

区域重力坍塌导致盆岭省的形成也发生在中新世,并且似乎反映了该区的北美地壳被拖拽到卢比 亚超级地体之下使该区的地壳厚度增加了一倍。结果中地壳很可能是热的且呈塑性,这样,在它之上 的岩石侧向流动决定了该地区上部地壳脆性变形的特征。总的来说,卢比亚与北美的拼合及相互作用 在这里概要表明:造山带是很容易解释为一个以多个弧-陆和弧-弧碰撞为特征的典型碰撞带,没有 必要用科迪勒拉型模式来解释其成因。

# 第11章 造山带往南延续

在南美洲北部,加勒比大弧 (Great Arc)的重要部分连同它的洋底高原与南美克拉通发生碰撞, 委内瑞拉、哥伦比亚、厄瓜多尔在坎帕期的 73~70 Ma 之间增生到向西倾的俯冲带之上的大陆 (Luzieux et al., 2006; Vallejo et al., 2006; Altamira-Areyan, 2009)。显然,外来岩石继续向南进入 了秘鲁西北部 (Feininger, 1980, 1987)。

Moores 等人(2002)提出了一个推测模型,他们认为大部分的南美边缘变形发生在晚侏 罗世—白垩纪并由弧碰撞引起,可能包括秘鲁和玻利维亚的外来中元古界阿雷基帕(Arequipa)地块(Ramos,2008)和巴塔哥尼亚(Patagonia)和火地岛(Tierra del Fuego)的罗卡斯·韦尔德斯(Rocas Verdes)蛇绿岩盆地的关闭。尽管对斯科舍(Scotian)边缘盆地南部的大部分地区知之甚少,且它们可能代表了中生代弧的碎片和微陆地(Barker,2001),但 很清楚的是:如同安替列群岛弧、斯科舍弧(Barker et al.,1991)一样,代表了一个迁移到大西洋的太平洋领域(Moores,1970; Pugh and Convey,2000),沿着它与南美洲和南极洲的转换断层边缘留下星星点点的散乱痕迹。

对于一些人来说, Moores 模型可能有些离谱, 但我相信它的价值巨大, 因为它已经解释了地壳 加厚和晚白垩世-古近纪的前渊 (DeCelles and Horton, 2003; Arriagada et al., 2006)、逆冲断层带, 并提出了安第斯中部的基底大型逆冲构造, 所有这些, 在我看来, 在目前的所有模型中都未能得到很 好地解释, 因为它们需要把中-上地壳岩石的巨大岩片从下地壳岩石中剥离出来, 而下地壳没有明显 缩短, 通过俯冲板片作用在地壳底部在摩擦力作用下向内陆运移数百千米。这样的变形通过尝试俯冲 的克拉通边缘更容易完成, 人们可以参照晚白垩世碰撞期间作为上盘的阿雷基帕 (Arequipa) 地块 (Romas, 2008)。

在智利和阿根廷中-南部的安第斯山脉,向东的晚白垩世阿格里奥(Agrio)褶皱-逆冲断 层带和相关的内乌肯(Neuquen)群前渊岩石(Cobbold and Rossello, 2003; Ramos and Kay, 2006),以及与之碰撞,并企图俯冲到南美西部边缘位于现在以约75 Ma 巴塔哥尼亚岩基为代 表的弧之下(Maloney et al., 2011),该岩石为一个主要为晚白垩世碰撞造山带延伸遍及南美 西部提供了的更多证据。

尽管南美科迪勒拉山系还遗留了许多复杂的地质问题需要学习,对秘鲁海岸岩基岩浆作用的匆匆 一瞥显示出与北美惊人的相似,无论在演化过程还是在形成时间上(图11.1)。当科迪勒拉深成岩体 和拉勒米变形发生在南美西缘时,如同它们沿着北美西部发生一样,卢比亚带状大陆可能曾经沿整个 美洲海岸延伸。因此,这个长期存在的假说可能需要修改,假说认为主要山链完全能够通过大洋岩石 圈俯冲产生,而不必有碰撞增生。


解释了它们在时间上和空间展布上的整体相似性

# 第12章 存在问题和下一步研究方向

根据这里提出的总结,似乎可以合理地得出结论: 塞维尔事件期间,在与北美碰撞后大部分科迪 勒拉相对于北美发生了向南迁移,然后在 80~50 Ma 的拉勒米事件期间向北迁移。美国西部内华达山 脉和大盆地这两个地区都缺乏古地磁学方面的研究,因此很难把它们放到适当的背景中。正如前面所 讨论的,内华达山脉现有的 102→97 Ma 的深成岩体资料表明其相对于北美最大的位移大约有 1000 km (Frei, 1986; Kent and Irving, 2010),但是,深成岩体有可能被褶皱或者重新磁化,归因于后续更年 轻的深成岩体侵位时被再次加热。

大盆地地区的侏罗纪和白垩纪深成岩岩石的详细古地磁研究也很重要,因为目前的研究依然不 足。如在犹他州的房屋范围(House Range)西侧出露的具有侏罗纪年龄的席状、整合深成岩体是一 个很好的研究对象,因为岩石顶、底面成层很好、倾角平缓。

最初位于白垩纪科迪勒拉岩基之间的盆地性质知之甚少,如同盆地内的俯冲极性那样,和这些问题一并出现的是后碰撞岩浆作用和剥露,是否由于板片断离,俯冲,或两者兼而有之产生。

弗朗西斯科在哪里,相关的增生杂岩和大谷地弧前盆地在拉勒米事件时期的位置在什么地方?在 100 Ma 左右,剧烈的转换压扭变形影响了内华达山脉岩基的中部和西部,但是被认为紧邻西部的沉 积岩却没有显示这种变形的证据。Wright和 Wyld (2007)提出:大谷地群的岩石恰好在墨西哥南部 沉积,但是这个观点依然有待考证。

有关主要断层的位移时间也是一个问题,这些断层走向大多数垂直于造山带,围限了主要的分 区,它们包括:路易斯和克拉克、奥罗菲诺、蛇河平原和菲尼克斯断层,是否这三个北部的断层限定 于卢比亚超级地体?它们没有明显错断北美的岩石,但是明显已经影响了那里的变形作用和沉积作 用。例如,路易斯和克拉克线性构造的断层明显地影响了前渊的沉积作用(Wallace et al., 1990),并 且该带向东的突出部分似乎标志着拉勒米厚皮构造变形的北部边界。同样地,奥罗菲诺断层东南的突 出部分明显和科罗拉多高原北端一致,是否是现在的断层正在切割卢比亚超级地体下盘 STEP 断层的 表现?

假设的蛇河右旋断层活动是什么时候? 它明显将加利福尼亚州的内华达北部和大谷地 15.5 Ma 的 拉夫乔伊(Lovejoy) 玄武岩(Garrison et al., 2008) 和类似年龄的俄勒冈州东南的斯蒂恩的玄武岩 (Steen's basalt) 水平错断了。

我之前认为是转换断层(Hildebrand, 2009)的菲尼克斯断层,清晰地将许多显著地貌区分开来, 例如盆-岭、内华达渐新世熔结凝灰岩的爆发和西马德雷山脉等;因此,有些人可能会反驳说它其实 是一个年轻的平移断层,但这明显无法解释 18.7 Ma 的桃花泉(Peach Springs)凝灰岩,这些凝灰岩 被认为穿过了可能的断层踪迹(Glazner et al., 1986),但是这个断层应该是早中新世,搞清这些断层 的时间和空间关系是对通过断层轨迹约束重建晚白垩世—古近纪推测个体活跃断层的基本前提要求 (Wyld et al., 2006)。

为更好理解侏罗纪变形作用,很有必要在内陆带进行详细工作。正如本文提出的,库特尼地体的 一部分有可能位于这个区域内。 此文献给 Cliff Hopson,他最早向我揭示了加利福尼亚州中生代岩石的奇妙和其中的问题。本书的所有内容归功于当地专家,其中我要特别感谢 Andrew Barth, Kevin Burke, Alan Chapman, Geoff Christe, John Dilles, Trevor Dumitru, Rita Economos, Marty Grove, Warren Hamilton, Jack Hillhouse, Ray Ingersoll, Ted Irving, Carl Jacobson, Oliver Jagoutz, Angela Jayko, Steve Johnston, Dennis Kent, Dave Kimbrough, Bob Miller, Pete Palmer, Scott Paterson, Jim Pindell, Matt Rioux, Sarah Roeske, Jason Saleeby, Rich Schweickert, John Shervais, Tom Sisson, Kathy Surpless, Allison Till, Cees van Staal, and Jim Wright, 不仅是因为他们回答了我的问题,同我讨论了具体的地质观点,并应我的请求发送了 PDF 文件,而且因为他们愿意与我分享他们深刻的看法和对科迪勒拉山系地质学的理解。我还要感谢同 Ray Price和 Jim Monger 激烈地辩论,这使我对加拿大 Cordillera 的看法更加深入和集中。

与美国地质调查局的 Bob Powell 的讨论帮助我在对 Transverse Ranges 复杂地质问题的研究中找到 了正确的方向。保罗链接(Paul Link)是理解各种碎屑锆石套的宝贵资源。正如前面提到的本书是 自我资助,所以我的家人忍受了时间和金钱的牺牲,见证了它的完成。Eldridge Moores, Paul Hoffman 和 John Wakabayashi 对手稿进行了正式评审并提出了很多建议,改进了成稿。最后,我想感谢 Peter Schiffman 的仁慈和支持。

## 附录1: 有关术语、地名翻译对照

- Stikinia
   Stikinia 是一个构造地层地体的名字,该地体位于加拿大不列颠哥伦比亚省的科迪勒拉山脉,形成于古生代和中生代时期的火山弧环境。
- Slate Creek——石板溪 石板溪是位于不列颠哥伦比亚省奥米尼卡地区一条河,这条河是曼森河的一条支流,从西流入。
- Omineca——奥米尼卡
   奥米尼卡是加拿大不列颠哥伦比亚省的一个选区,1916年首次出现在大选中。
- Cache Creek——卡什溪
   卡什溪是加拿大不列颠哥伦比亚省的汤普森内部的波拿巴河的一条支流,在横贯加拿大不列颠哥
   伦比亚省和凯里布高速公路交界处的卡什溪镇加入。
- Rattlesnake Creek——响尾蛇溪
   响尾蛇溪是油漆溪(Paint Creek)的一条支流。
- Salinian Block——盐碱地块 盐碱块体或盐碱地体是位于加州圣安德烈亚斯断层系统主构造痕迹西部的一个地质单元,南以文 图拉县(Ventura County)的大松树断裂(Big Pine Fault)为界,西以纳西缅托(Nacimiento)断 裂为界。
- 宾夕法尼亚纪(Pennsylvanian)
   是根据美国的宾夕法尼亚州来命名的,因为宾夕法尼亚纪的岩层广泛分布在这里,年代大约位于 318.1±1.3至299±0.8 Ma,属于古生代石炭纪的一部分,相当于晚石炭世。
- 拆离断层(detachment fault) 最早由 Pierce 于 1963 年提出,当时是指叠瓦状逆冲断层的底板断层,即滑脱面。Davis 1980 年 将其应用于伸展构造,定义为"结晶变质基底杂岩与上覆沉积盖层之间的大型低角度正断层或 伸展断层"。
- \_\_\_\_\_
- Antler orogeny——安特勒造山运动
- Brookian orogeny——布鲁克造山运动
- Browns Fork orogeny——布朗斯福克造山运动
- Elko orogeny——艾尔克造山运动
- Ellesmerian orogeny——埃尔斯米尔造山运动
- Hudsonian orogeny——哈德逊造山运动
- Laramide orogeny——拉勒米造山运动
- Sevier orogeny——塞维尔造山运动
- Sonoman orogeny——索诺玛造山运动
- Laurentian——劳伦系
- Albian——阿尔必阶
- Antian——阿尔妃·所
   Aptian———阿普第阶
- Bajocian——巴柔阶
- Bajocian 已采例
  Barremian——巴列姆阶
- Berriasian———贝利亚斯阶

- Campanian——坎帕阶 (白垩纪倒数第二个时期)
- Cenomanian——森诺曼阶
- Coniacian——科尼亚斯阶 (晚白垩世的第三个时期)
- Kimmeridgian——基默里奇阶
- Maastrichtian——麦斯里希特阶
- Neocomian——尼奥科姆阶
- Oxfordian——牛津阶
- Santonian——桑托阶
- Tithonian——蒂托阶
- Valanginian——凡兰吟阶
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- Archeocyathids——古杯动物门
- Artesia sequence——阿蒂西亚序列
- Belt-Purcell-Windermere——贝尔特-珀塞尔-温德米尔
- Bisbee basin——比斯比盆地
- bowser basin——鲍尔斯盆地
- Boyden Cave——伯伊登洞穴
- Burgess——博格斯页岩
- Burgos Basin——布尔戈斯盆地
- capstone sequence——压顶石层序
- Carmacks volcanic——卡马克斯火山岩
- Cascadia——卡斯卡底古陆
- Columbia river basalt——哥伦比亚河玄武岩
- Colville Basin——科维尔盆地
- juan de fuca plate——胡安德富卡板块
- Lovejoy basalt——拉夫乔伊玄武岩
- McCloud——麦克劳德
- Megathrust sheet——大型逆冲岩席
- Nicola——尼古拉群
- Plutonic addition rate——深成岩增加率
- Rubian superterrane——卢比亚超级地体
- Salinian block——盐碱地块
- San Andreas F. ——圣安德列亚斯断层
- Scolithus——虫形石
- Selwyn——塞尔温盆地
- slab-failure magmatism——俯冲板片断离岩浆作用
- Smartville Complex——斯马特维尔杂岩
- Terranes with McCloud & Tethyan faunas——具有麦克劳德 & 古地中海生物群的地体
- Trans-Mexican Volcanic Belt——跨墨西哥火山带

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- Transpressional collision——转换挤压碰撞
- Wrangellia Terrane——兰格利亚地体
- Wyoming salient——怀俄明凸
- Yukon-Tanana——育空-塔纳纳地块

- Alabama Hills——阿拉巴马山(加利福尼亚州)
- Alberta-British Columbia——阿尔伯塔-不列颠哥伦比亚省
- Aztec——阿芝特克
- BARSTOW SE——巴斯托东南
- Bean Canyon——贝恩峡谷
- Bird Spring——乌泉
- Bourgeau——布尔若
- Cabin-Medicine Lake——卡宾麦迪逊湖
- Cache Creek——卡什溪
- Cadomin——卡多明组
- Calaveras——卡拉维拉斯
- Calgary——卡尔加里
- Canyon Range-Wah Wah-Pavant---峡谷山脉-哇哇-帕旺
- Cascade Mountain Range——卡斯克德山脉
- Cedros——塞德罗斯岛
- Charleston-Nebo---查尔斯顿尼波
- Chuckwalla——秋克华拉山
- Corn Springs——玉米泉
- Crawford——克劳福德
- Crowsnest——克罗斯内斯特
- Dakota——达科他
- Dale Lake——戴尔湖
- Death Valley——死亡谷
- Denali——麦金利山
- Dome Rock——岩石圆顶
- Don Pedro-----唐佩德罗
- Dry Butte——干孤峰
- E. Klamaths——东克拉马斯
- Fairview Valley——费尔围峡谷
- Fiddle Creek——小提琴溪
- Fort Jones——琼斯堡
- Garlock F. ——加洛克断层
- Giant Forest——巨林区
- Goddard 戈达德
- Goldhammer——-戈德哈默
- Grouse Creek——格劳斯溪
- Guerrero——格雷罗
- Hayfork——海福克
- Horsethief Creek——马贼溪
- Humboldt——洪堡特 (南达科他州)
- Isabella Lake——伊莎贝拉湖
- John Muir——约翰・缪尔
- Josephine——约瑟芬
- Kechika——科奇卡

- Kitt Peak——基特峰
- Klamath——克拉马斯
- Kodiak——科迪亚克
- Lake Tahoe——太浩湖
- Lewis-Eldorado-Hoadley-Steinbach——路易斯-埃尔多拉多-霍德利-斯坦巴克
- Mackenzie——麦肯齐河
- Mariposa——蝴蝶百合
- McCoy——麦考伊
- Meade——米德
- MEX Sonora——墨西哥索诺拉
- Mojave desert———莫哈韦沙漠
- Mount Givens——吉文斯山
- Mount Goddard——-戈达德山
- Mount Jura——汝拉山
- Mt. Whitney——惠特尼峰
- Nixon Fork——尼克松福克
- Nogales-Cananea-Nacozari——诺加莱斯-卡纳内-纳科札里
- North Fork——北福克
- Oak Creek——橡树溪 (科罗拉多州)
- Oak Creek——奥克里克
- Oaxaca——瓦哈卡
- Ogilvie——奥格尔维
- Old Woman-Piute Range——老妇-派尤特山脉
- Pahrump——帕朗
- Palen——帕伦
- Palmetto——蒲葵
- Phoenix F. ——菲尼克斯断层
- Pigeon Point——鸽子岬
- Pine Nut----松果
- Pine Nut---松子
- Piute——派尤特
- Portal Ridge——门户脊
- Prince William——威廉王子
- Pyramid Peak——金字塔峰
- Queen Charlotte——夏洛特王后
- Raft River——拉夫特河
- Redding——雷丁
- Ritter Range——里特山脉
- Rivera——里韦拉
- Rocky Mountain trench——落基山山谷
- Rodman——罗德曼
- Rogue Valley——罗格谷
- ruby——红宝石
- S. Central AZ——亚利桑那中南部

- Sacramento——萨克拉门托
- Saddlebag Lake——鞍袋湖
- Saddlebag——鞍袋
- Sailor Canyon——水手峡谷
- Salmon River——萨蒙河
- San Emigdio——圣埃米格迪奥
- San Gabriel——圣盖博
- San Miguel de Allende——圣米格尔德阿连德
- San Pitch——圣节
- San-Bernardino——圣贝纳迪诺
- Selwyn——塞尔温
- Sequoia——红杉
- Seward Peninsula——苏华德半岛
- Shamrock——三叶草
- Sierra de Salinas——萨利纳斯山脉
- Sierra Madre Occidental——西马德雷山脉
- Sierrita——西雅里塔
- Slate Creek-Lake Combie——斯莱特溪-结合湖
- Slide Mountain——斯莱德山
- Smith River——史密斯河
- Sonoman——斯诺曼
- Sonora——索诺拉省 (墨西哥北部)
- South Fork——南福克
- SW AZ & SE CALIF——亚利桑那西南 & 加利福尼亚东南部
- Sylvania——西尔瓦尼亚
- Palmetto——帕尔梅托
- Table Mountain——平顶山
- Talkeetna——塔尔基特纳
- Taylorsville——泰勒斯维尔
- Tehama-Colusa-----蒂黑马-科卢萨
- Trigo Peak——小麦高峰
- Triple Divide Peak——三分峰
- Tuolomne River——图奥勒米河
- Tuttle Creek——塔特尔溪
- Victorville——维克多维尔谷
- Vizcaino——比斯凯诺半岛
- Wah-Pavant——瓦赫-帕旺
- Western Hayfork——海福克西部
- White-Inyo Mountains——白印优山脉
- Whitney——惠特尼山 (位于美国加利福尼亚州东部)
- Wood Canyon----伍德峡谷
- Woodlark Basin——木百灵盆地
- Wooley Creek——邬里溪
- Yerington——耶灵顿

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- Yolla Bolly——约拉波利 Zihuatanejo——芝华塔尼欧 •

## 附录2:美国州府名(缩写)

- 美国由 50 个州和1 个直辖特区(首都所在地华盛顿哥伦比亚特区)组成
- 州——缩写——首府
- 阿肯色 Arkansas——AR——小石城 Little rock
- 阿拉斯加 Alaska——AK——朱诺 Juneau
- 艾奥瓦 Iowa------得梅因 Des Moines
- 爱达荷 Idaho------博伊西 Boise
- 北达科他 North dakota——ND——俾斯麦 Bismarck
- 北卡罗来 North carolina——NC——纳罗利 Raleigh
- 宾夕法尼亚 Pennsylvania——PA——哈里斯堡 Harrisburg
- 俄亥俄 Ohio-----OH-----哥伦布 Columbus

- 佛罗里达 Florida——FL——塔拉哈西 Tallahassee
- 佛蒙特 Vermont——-VT——蒙彼利埃 Montpelier
- 弗吉尼亚 Virginia——VA——里士满 Richmond
- 华盛顿 Washington-----WA-----奥林匹亚 Olympia
- 加利福尼亚 California——CA——萨克拉门托 Sacramento
- 堪萨斯 Kansas——KS——托皮卡 Topeka
- 康涅狄格 Connecticut——CT——哈特福德 Hartford
- 肯塔基 Kentucky——KY——法兰克福 Frankfort
- 路易斯安那 Louisiana——LA——巴吞鲁日 Baton Rouge
- 罗得岛 Rhode island-----------------------普罗维登斯 Providence
- 马里兰 Maryland——MD——安纳波利斯 Annapolis
- 蒙大拿 Montana——MT——海伦娜 Helena
- 密苏里 Missouri——MO——杰斐逊城 Jefferson City
- 密西西比 Mississippi——MS——杰克逊 Jackson
- 密歇根 Michigan——MI——兰辛 Lansing
- 缅因 Maine——ME——奥古斯塔 Augusta
- 明尼苏达 Minnesota——MN——圣保罗 St. Paul
- 南达科他 South dakota——SD——皮尔 Pierre
- 南卡罗来纳 South carolina——SC——哥伦比亚 Columbia
- 内布拉斯加 Nebraska——NE——林肯 Lincoln
- 纽约 New york——NY——奥尔巴尼 Albany
- 特拉华 Delaware——DE——多佛 Dover
- 田纳西 Tennessee———TN———纳什维尔 Nashville

- 威斯康星 Wisconsin-------------------------麦迪逊 Madison

- 新罕布什尔 New hampshire——NH 康科德 Concord
- 新泽西 New jersey——NJ——特伦顿 Trenton

- 伊利诺伊 Illinois——IL——斯普林菲尔德 Springfield
- 印第安纳 Indiana——IN——印第安纳波利斯 Indianapolis
- 犹他 Utah——UT——盐湖城 Salt Lake City
- 佐治亚 Georgia——GA——亚特兰大 Atlanta

附录3:美国西部重要地质事件

时间/Ma	事件	标志
700~360	北美西缘为被动边缘,整个寒武纪是碎屑沉积(厚达6km),然后整个二叠纪是 碳酸盐岩沉积(>3km)	稳定的北美板块
360~350	罗伯特(Robert)山地体增生(洋底沉积物,火山岩,黑色页岩;+碎屑物)	安特勒 (ANTLER) 造山运动
260~250	宝山 (Golconda) 地体增生 (洋底沉积物+火山岩)	索诺玛 (SONOMA) 造山运动
280~140	黑石 (Black Rock) 地体堆积物 (洋盆和岛弧)	岛弧
250~180	Jungo 地体泥质页岩堆积物 (>7 km)	沉积
250~135	松果 (Pine Nut) 地体岩石堆积物 (火山岩、碳酸盐、碎屑物); 深成岩体年龄 230~130 Ma	岛弧,斑岩铜矿
245~140	伊甸园(Paradise)地体的岩石堆积(火山岩);深成岩体年龄150 Ma	岛弧,斑岩铜矿
180~140	西部所有岩石变形	内华达造山运动
180~140	广泛的变质作用	变质作用
160~140	侏罗纪大陆弧	大陆弧,斑岩铜矿
135~65	广泛的变质作用持续发生	变质作用
130~80	白垩纪大陆弧	大陆弧,斑岩铜矿
140~50	塞维尔 (SEVIER) 逆冲推覆	塞维尔 (SEVIER) 造山作用
ca. 130?	Jungo 地体向东逆冲推覆	变形
105?	松子 (Pine Nut) 地体侵位 (走滑断层)	变形
75~65	过铝质的、与碰撞相关的花岗岩	变形
45~20	大陆弧向南侧伏,由俯冲板片崩塌导致的后碰撞伸展火山作用	斑岩钼矿,斑岩铜矿
40~36	赋存 Au 的卡林型(Carlin)沉积岩形成	Au 矿床
36~26	硅质熔结凝灰岩爆发	火山作用
16~14	黄石公园大量热点玄武岩+流纹岩出现	火山作用
16~14	沿着科罗拉多河延伸走廊被极大地延长	变形
16~14	与双峰组合有关的浅成低硫化 Au 矿床侵位	Au 矿床

(据 Stephen D. Ludingten 资料整理)

### 附录4:北美地壳运动

- 【凯诺拉运动】Kenoran orogeny 又译基诺尔运动。史托克维尔(CH Stockwell, 1964)创名。加拿 大地盾前寒武纪早期的一次造山运动,是根据北美加拿大苏必利尔湖、大奴湖地区太古宇上部提 米斯卡明群与古元古界下部布鲁斯群之间的角度不整合确定的。这次运动有深成岩的侵入、变质 和变形作用。同位素年龄值为23.9亿~26亿年以前,原认为是地球最早的前寒武纪造山运动之 一,也是划分太古宇和元古宇的依据。它与阿尔戈马运动、非洲南部的罗得西亚褶皱和欧亚大陆 的贝罗摩尔褶皱的造山运动大致相当。相当于中国的五台运动(25亿年)和新太古代的阜平运动 (26亿年)。
- 2.【阿尔戈马运动】Algoman orogeny 曾译阿尔冈曼运动。发生于北美洲太古宇休伦系与元古宇阿尔 冈曼系之间的造山运动。根据美国明尼苏达及其邻区安大略等地前寒武纪的褶皱运动确定的。伴 随有花岗岩侵位,同位素年龄值为24亿年以前。与加拿大地盾新太古代末期的凯诺拉运动相当。
- 3.【基诺拉运动】Kenoran orogeny 发生在加拿大太古宙的一次造山运动,影响范围波及加拿大苏必利 尔湖地区。其后是哈德孙运动。此术语已不再使用。
- 4.【哈德孙运动】Hudsonian orogeny 在加拿大地盾发生在古元古界末的一次构造运动。根据北美加拿 大苏必利尔区古元古界上部阿尼米基群与新元古界下基维诺(Keweenaw)群之间的角度不整合确 定的,发生时限距今17.5亿~18亿年左右。在基诺拉造山运动之后,格林威尔造山运动之前。大 致相当于中国的吕梁运动。
- 5.【彭诺克运动】Penokean orogeny 北美中元古界与古元古界之间的造山运动。是根据北美明尼苏达和密歇根州的阿尼米基群与基维诺群之间的角度不整合确定的。此运动使前寒武纪地层变形并伴随花岗岩侵入,运动发生时限约在17亿年。与加拿大地盾的哈德孙运动相当。与中国的吕梁运动第一幕或主幕(19亿~18亿年)相当。
- 6. 【凯蒂利德运动】Ketilidian orogeny 北美洲发生在 18 亿~16 亿年前太古宙的造山运动幕。此运动 是在格陵兰岛确定的。
- 7.【劳伦运动】Laurentian orogeny 又译劳伦构造作用幕。太古宙末的造山运动,系据加拿大地盾太古 宙与元古宙地层之间的显著不整合而确定。这一运动使太古宇库契钦格组和基瓦丁统发生强烈褶 皱和变质,并伴随劳伦花岗岩的侵入。纳利夫金(Hаливкин, 1932)认为是太古宙末的造山运 动。它造成了太古宙与元古宙地层之间的显著不整合。劳伦一词使用较混乱,定义也不严格,既 指加拿大地盾的前寒武纪花岗岩,又指那里的前寒武纪造山作用。得名的地点为加拿大东部劳伦 斯河西北的劳伦高原,而1863年洛甘(Logan)在那里调查的劳伦花岗岩已确定其同位素年龄值 仅为1000 Ma,后来劳桑(Lawson, 1885)又把接近美国和加拿大边界苏必利尔湖西北部的最古 老花岗岩误称为劳伦花岗岩。舒克特(Schuchert)又据此得出他的劳伦运动或劳伦造山运动,认 为这一运动结束了太古宙。20世纪70年代证明劳桑的劳伦运动早于距今2400 Ma的老阿尔冈曼造 山运动或阿尔冈曼花岗岩。有人建议,为避免混乱,把劳伦一词恢复为洛甘原来的含义,即相当 于格林威尔运动。
- 8.【休伦运动】Huronian orogeny 北美休伦纪(24 亿~21 亿年)地层沉积之后、元古宙发生的褶皱运动,是根据北美休伦族的名字命名。有人还认为此运动发生在太古宙末。前苏联称为休伦构造作用幕或休伦褶皱。
- 9.【格伦维尔运动】Grenvillian orogeny根据北美加拿大苏必利尔区新元古界中基维诺(M. Keweenaw)群与新元古界上基维诺(O. Keweenaw)群之间角度不整合确定的。发生在距今10亿~8.8年前的前寒武纪一次重要的深成侵入、变质、变形事件,波及加拿大地盾东南部边缘的广大地区。地球上已经确定在许多地区曾有与格伦维尔运动时限相近的构造运动,如在现今的哥伦比亚、墨西哥、北美东部、格陵兰岛东部以及斯堪的纳维亚半岛(这里把它称为达尔斯兰造山运动)

或哥特造山运动)。在加拿大地盾东南部沿一条从休伦湖北岸向东北延伸的直线露头清晰可见。 这一造山运动是因大西洋发生板块运动使岩石向西北位移引起的。其时限与中国的四堡(燕辽) 运动大致相当,略早于晋宁运动,属前震旦纪晚期构造运动。

- 10. 【基拉尔尼运动】Killarneyan orogeny 加拿大地盾南部元古宙后半期发生的造山运动。基拉尔尼花 岗岩贯入休伦(Huronian)系中,且不整合覆于元古宙后期的基维诺(Keweenawan)系之上。
- 11. 【加达裂谷作用】Gardar rifting 在北美发生在即凯蒂利德造山运动之后,大约在 14 亿~10 亿年前,属裂谷作用的一个幕。
- 12.【阿瓦朗运动】Avalonian orogeny 发生于现今北美洲东部边缘约 650~500Ma 前(寒武纪到奥陶 纪)的一个造山事件,与大西洋形成时的裂谷作用和火山活动有关。阿瓦朗造山事件的遗迹沿 阿巴拉契亚造山带东南侧的山脊从加拿大的纽芬兰省到美国的佐治亚州断续出现,在大西洋沿岸 平原沉积层之下的岩层中也有阿瓦朗造山事件遗迹发现。在阿瓦朗造山事件遗迹表现明显的典型 地区,有年龄为 575Ma 的花岗岩,与其下有化石证据的下寒武统之间有前寒武纪晚期岩系存在。 阿瓦朗一名源自加拿大的纽芬兰省阿瓦朗半岛(Avalon Peninsula)。
- 13.【塔科尼运动】Taconic (Taconian) orogeny 曾译塔康运动。发生在奥陶纪晚期的造山运动。 "Taconic"一词源自纽约市哈得孙河东岸近南北向的塔科尼克岭 (Taconic Range)。塔科尼克运动遗迹在阿巴拉契亚山脉北段许多地区保存得都很好,包括在加拿大东部和美国东北的部分,在 不少地方都可以用含化石地层将运动时限严格限定为晚奥陶世,或者,把许多脉动期包括进去, 将运动时限延伸至自奥陶纪早期到志留纪早期。在加拿大纽芬兰,此运动被称为亨伯造山运动 (Humberian orogeny)。在阿巴拉契亚山脉南部,这一运动是由艾佩特斯洋 (Iapetus Ocean) 西缘 俯冲引起的。它导致了艾佩特斯洋西部盆地闭合和皮德蒙特 (Piedment) 微陆块与北美洲碰撞以 及更北部大陆架的沉陷。
- 14.【萨尔运动】Saalic orogeny 曾译萨阿尔运动。史蒂勒(H. Stille) 1924 年创名。华力西构造运动的第四幕,发生时代在二叠纪早期欧坦(Autunian)期和萨克森(Saxonian)期(现在欧洲的早二叠世欧坦期和中二叠世萨克森期)之间。运动遗迹见于下列地区:波兰和捷克边界上的苏台德山、西班牙与法国边界上的比利牛斯山、阿尔卑斯山东端的卡尔尼克山(Kalnik,在克罗地亚境内)。俄罗斯的乌拉尔山脉、北美洲东部和智利境内的西科迪勒拉山脉也有同期构造运动。
- 15.【普法尔茨运动】Pfalzian orogeny 曾译法尔琴运动、法尔兹运动。史蒂勒(H. Stille) 1924 年创 名。华力西构造运动的第五幕,发生时代在二叠纪末。在德国萨尔州(Saarland)和莱茵兰德-普法尔茨(Rhinland Pfalz)州(莱茵河通过本州的东部)黑森林山(Hessen Wald)地区,下三 叠统的斑砂岩统(Buntsandstein)不整合于二叠系赤底统(Rotliegende)之上。运动遗迹见于下 列地区:法国的孚日山脉(Vosges)、西班牙与法国边界上的比利牛斯山、英国、阿尔卑斯山西 段。在俄罗斯的乌拉尔山脉和阿根廷境内的科迪勒拉山也有同期构造运动。李四光(1931)认 为与中国的苏皖运动相当。
- 16.【阿卡迪运动】Acadian orogeny 曾译阿卡迪亚运动、阿卡德运动。北美洲阿巴拉契亚山地区第三 期造山活动。是自美国的纽约州到加拿大芬迪湾的阿巴拉契亚山脉北段表现比较明显的古生代中 期造山运动,在纽约市哈德孙河东岸近南北向的塔科尼克岭(Taconic Range)以东构造活动最剧 烈,阿瓦朗地体(Avalon terrane)(在阿瓦朗运动遗迹的典型地区,花岗岩的同位素年龄为575 Ma)往北西方向逆冲。在加拿大加斯佩(Gaspé,位于加拿大东南部圣劳伦斯河入海口处)及其 邻近地区,根据地层确定的运动主期时限为晚泥盆世早期,但变形作用、深成岩体的侵入和变质 事件持续了更长的时间,放射性年龄测定确定岩体的侵入和变质事件发生在360~330 Ma以前。 有人认为此运动相当于欧洲华力西造山运动的布雷顿幕和美国西部的安特勒(Antler)运动。 Acadian 一词源自于加拿大东南沿海过去法国殖民地时期的地名 Acadie (英语用 Acadia)。
- 17.【安特勒运动】Antler orogeny 罗伯茨(R. J. Roberts, 1951)创名。发生于泥盆纪晚期至密西西比 亚纪早期的造山运动,导致美国西部内华达州中北部大盆地(Great Basin)地区的古生代地层广

泛变形,西部的优地槽岩层沿着现在作为外来推覆体的罗伯茨山脉(Roberts Mountains)逆冲于 东部的冒地槽岩层之上。在主幕之后,还有较小的造山脉动,一直持续到二叠纪。此运动大致与 北美洲东部的阿卡迪(Acadian)运动相当。"Antler"一词源自内华达州巴特尔芒廷(Battle Mountain)附近的安特勒四方形山峰(Antler Peak Quatrangle)。

- 18.【阿勒格尼运动】Alleghany orogeny, Alleghanian orogeny 发生于美国阿巴拉契亚山脉南段和中段阿勒格尼山区(Alleghany Mountains,弗吉尼亚州-西弗吉尼亚州)及其附近阿勒格尼高原岩石变形的晚古生代造山运动。因为缺少可靠的能限定年代的上覆地层,所以运动的时限不能准确确定。宾夕法尼亚亚纪岩层在许多地方受到运动波及,而宾夕法尼亚亚系最上部-下二叠统的敦卡德岩系(Dunkardian series)只在几个地方发现受到影响。由伍德沃德(Woodward, 1957, 1958)创名。始于早石炭世,结束于二叠纪末的造山运动。由于北美洲和非洲发生碰撞,导致阿巴拉契亚山脉南段和中段(从美国亚拉巴马州到宾夕法尼亚州)西缘的下古生代基底和下二叠世岩层变形,构成大体上呈南西西-北东东向延伸的海西造山带的一部分。运动的影响向北远至加拿大的新不伦瑞克省和纽芬兰省。
- 19. 【阿巴拉契亚运动】Appalachian revolution 此术语是用来表述结束古生代的重大地壳活动事件,导致了北美洲现今阿巴拉契亚山脉南段和中段的岩层发生变形。现在多用术语阿勒格尼造山运动来说明这一地质事件。
- 20. 【塞维尔运动】Sevier orogeny 于晚白垩世早期发生在现今美国犹他州大盆地(Great Basin)东缘的一次造山运动。运动内容包括岩层褶皱和向东逆冲。
- 21.【拉勒米运动】Laramide orogeny, Laramian orogeny 曾译拉拉米运动。又称拉勒米革命(Laramide Revolution)。史蒂勒(H. Stille) 1924年创名。发生于白垩纪末的构造运动。典型地区在美国落基山脉东部。术语'Laramide orogeny'源自美国怀俄明州和科罗拉多州这一时代的沉积称拉勒米建造(Laramie Formation)。不同的学者对拉勒米运动有不同的理解。在美国,认为这是从白垩纪晚期至古新世末发生在现今落基山脉东部有深成岩侵入并伴随有矿床生成的一次地壳变动,当地一般称这些侵入岩和矿床为'Laramide'。加拿大所称的拉勒米运动还包括了美国的塞维尔运动。还有人认为拉勒米运动时限从白垩纪晚期至始新世末或更晚,涉及现在北美洲的整个科迪勒拉造山带。李四光(1939)认为相当于中国的闽浙运动,李春昱(1943)等提及与四川运动相当。
- 22.【施蒂里亚运动】Styrian orogeny, Styrian phase 曾译斯提里运动。史蒂勒(1924)创名。阿尔卑斯旋回第七幕,发生于中新世阿基坦期与布尔迪加尔期(Burdigalian)之间的构造运动。创名地点为奥地利中南部施蒂里亚(Styria)山、斯提里阿尔卑斯山区。出现在阿尔卑斯山、喀尔巴阡山、利古里亚-亚平宁山、塔夫利山、美洲西部及其他地区。前苏联文献认为发生于中新世中期,由两个亚幕组成:一在海尔微期的早期,一在托尔顿(Tortonian)期的早期。前苏联有在如高加索,这一构造作用幕发生在早中新世末期和中新世初期。
- 23.【瓦拉几亚运动】Wallachian orogeny 又译瓦拉赤运动。史蒂勒(1924)创名。阿尔卑斯山旋回末期一幕,发生于上新世末的构造运动。此运动遗迹见于罗马尼亚境内南喀尔巴阡山和瓦拉几亚(Wallachia)地区。同期运动在阿尔卑斯山、喀尔巴阡山、利古里亚-亚平宁山、狄拿里克、高加索、印度及美洲西部地区都有发育。
- 24.【内华达运动】Nevadan orogeny, Nevadian orogeny, Nevadic orogeny 侏罗纪和早白垩世发生在现 今北美洲科迪勒拉山系西部的造山运动,表现为构造变形、岩石变质和深成岩浆入侵。典型地区 在美国加利福尼亚州境内的内华达山脉,根据那里含化石地层判断,基底以上地层的变形可能结 束于侏罗纪晚期(基末利阶(Kimmeridgian)与波特兰阶(Portlandian)之间),但在其他地方 有比这里时间更早和更晚的内华达期构造变形。花岗岩的侵位和其他深成岩浆活动持续的时间比 构造变形持续的时间要长,深成岩的同位素年龄值在180~80 Ma,即从早侏罗世至白垩纪中晚 期。地质学家们对'Nevadan orogeny'的理解、使用有些不同,有的对其在时间和空间方面加以

严格限定,有的则将其用得更广泛,不怎么加以限制。最好还是按史蒂勒(Wilhelm Hans Stille, 1876~1976)原意将其作为一个造山时期对待。

- 25. 【拉勒米-哥伦布运动】Laramide Columbian orogeny 发生于白垩纪晚期至始新世的造山运动。影响 范围从美国西南部延伸到南美北部的地区,其成因可能是法拉隆(Farallón) 板块向北美板块之 下俯冲时伴生的逆冲断层活动所致。
- 26. 【帕萨迪纳运动】Pasadenan orogeny 史蒂勒(1936)创名。显生宙最年轻的一次造山运动,发生 在更新世中期的构造运动。根据加利福尼亚南部上新统与下更新统的关系确定,上更新统不整合 覆于其上。帕萨迪纳为加利福尼亚州南部洛杉矶附近的一个城镇(Pasadena)。
- 27.【喀斯喀特运动】Cascadian revolution 又译卡斯卡底运动。又称安的列斯运动。舒克特(C. Schuchert)等20世纪20年代所使用的一个术语,指北美西部新近纪末期的造山运动。他们认为这一次地壳活动结束了新近系沉积。但在典型地区(喀斯喀特山脉)未发现此期间所发生的明显地壳运动,故此词应予废弃。

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## **MESOZOIC ASSEMBLY OF THE NORTH AMERICAN CORDILLERA**

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